THE USE OF A CHARACTERISTIC BASED SCHEME FOR THE 2-D SHALLOW WATER EQUATIONS.

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1. <u>INTRODUCTION</u>

In recent years much effort in the field of computational fluid dynamics has been invested in approximately solving evolutionary conservation laws where discontinuities in the solution constitute a very important, if not the most important, part of the problem. One of the more successful strategies employed on these problems has been that of the family of methods based upon a characteristic decomposition of the system of equations known as Godunov methods. This is done by assuming that the grid values represent average values of the solution over the surrounding cell. A Riemann problem is then solved, approximately or exactly, at each cell interface. This problem has a wavelike solution and hence we may apply upwind difference schemes based upon the characteristic speeds and directions.

Of these methods one of the most popular, and the one we shall consider here, is that due to Roe. As with the similar methods that fall into this category Roe's method has been designed to provide accurate solutions away from jumps while being able to avoid the oscillations

generated by classical second order accurate schemes, Lax-Wendroff for example, at discontinuities. In common with most numerical schemes for conservation laws these schemes are conservative but other desirable properties, such as entropy satisfaction and total variation diminishing (TVD), can also be ensured with the scheme.

As has been said these schemes have primarily been used and developed for discontinuous flows. Why then should we want to try applying these schemes to atmospheric and oceanographic flows governed by the shallow water equations and not normally associated with violent changes? Clearly second order schemes are desirable to achieve accuracy on reasonably sized grids but more importantly many of the design features of these schemes are as relevant to the shallow water equations as they are to the Euler equations, conservation and the existence of wavelike solutions in particular. Discontinuities may arise in atmospheric and oceanographic flows as fronts or tidal surges and bores and so the shock handling capabilities of these schemes will not go to waste. papers, which include Roe et al. (1981, 1984), Sells (1980), Sweby (1984), Glaister (1985,1987) and Priestley (1987) have demonstrated the success of Roe type schemes in dealing with discontinuities. Here we concentrate on their ability to compete in areas which exhibit smooth solutions, for flows where jumps in the solution are only a possibility rather than a vital part of the calculation.

In Section 2 the philosophy behind the Roe approach will be given and B functions and flux limiters will be introduced as a way of generating a family of methods.

In Section 3 the shallow water equations are stated and the Roe decomposition performed in some detail to obtain the necessary building blocks of the method.

In Section 4 the various methods that can be obtained by this approach are applied to Grammeltvedt's problem (1969) and it is demonstrated that correct treatment of the source/forcing term is essential. It is shown that schemes capable of handling discontinuities can achieve the same, if not better, accuracy than the more common second second order schemes. (Indeed other numerical experimentation has shown the basic first order Roe scheme to be as accurate for the shallow water equations as the Lax-Wendroff method when the latter requires some artificial viscosity to remove oscillations.) The experience gained on

this problem is then applied to the problem of coastal flow. Here the main source of difficulty is the computationally open boundaries, and their treatment will be discussed.

2. ROE'S SCHEME

In this section we give a brief account of the philosophy behind the Roe type schemes, largely following Roe's original paper (1981), and refer the reader to that paper for more details.

Consider the initial-value problem for a hyperbolic system of conservation laws

$$\underline{\mathbf{q}}_{t} + \underline{\mathbf{F}}_{x} = 0 \tag{2.1}$$

with initial conditions

$$\underline{\mathbf{q}}(\mathbf{x},\mathbf{o}) = \underline{\mathbf{q}}_{\mathbf{o}}(\mathbf{x}), \tag{2.2}$$

where $\underline{F} = \underline{F}(\underline{q})$ and the Jacobian matrix $A = \partial \underline{F}/\partial \underline{q}$ has real eigenvalues.

Introduce now the discrete representation $x_i = x_0 + i\Delta x$, $t_n = t_0 + n\Delta t$ and suppose that \underline{q}_i^n approximates $\underline{q}(x_i, t_n)$. If the initial data (2.2) is specialised to

$$\underline{q}(x,o) = \underline{q}_{L}(x(o)) ; \underline{q}(x,o) = \underline{q}_{R}(x(o))$$
 (2.3)

then we have a so-called Riemann problem which has wavelike solutions. Godunov (1959) produced a numerical scheme for the solution of (2.1) that treats the data as a set of constant states separated by discontinuities at the points $\mathbf{x}_{\mathbf{i}+\mathbf{k}}$, say, a Riemann problem then being solved in each interval, with a time-step restriction such that the waves from one jump do not interfere with those from neighbouring jumps.

With the Roe type methods a slightly different approach is taken by considering the approximate problem

$$\underline{\mathbf{g}}_{\mathsf{t}} + \widetilde{\mathbf{A}}(\underline{\mathbf{g}}_{\mathsf{R}} - \underline{\mathbf{g}}_{\mathsf{L}}) = 0. \tag{2.4}$$

where $\widetilde{A} = \widetilde{A}(\underline{q}_L,\underline{q}_R)$ is a constant matrix that is chosen to be representative of the local conditions. A further restriction is that \widetilde{A} must satisfy the following properties collectively called Property U:

- (i) It constitutes a linear mapping from the vector space \underline{q} to the vector space \underline{F} .
- (ii) As $\underline{q} \to \underline{q}_R \to \underline{q}$, $\tilde{A}(\underline{q}_L,\underline{q}_R) \to A(\underline{q})$ where $A = \frac{\partial \underline{F}}{\partial \underline{q}}$.
- (iii) For any $\underline{q}_{L},\underline{q}_{R}$. $\widetilde{A}(\underline{q}_{L},\underline{q}_{R}).(\underline{q}_{L}-\underline{q}_{R}) = \underline{F}_{L} \underline{F}_{R}$.
- (iv) The eigenvectors of A are linearly independent.

Finding an \widetilde{A} that satisfies Property U is not a trivial problem. Neither of the two 'obvious' choices $\widetilde{A} = \frac{1}{2}(A_L + A_R)$ or $\widetilde{A} = A(\frac{1}{2}(\underline{q}_L + \underline{q}_R))$ will in general satisfy (iii), but in the case of (2.4) Roe has shown that one can be found.

This leads to a basic first order upwind scheme for which at each time-level, at each jump, we calculate the eigenvalues λ_i , the eigenvectors \underline{e}_i , and associated strengths α_i of $\widetilde{A}(\underline{q}_L,\underline{q}_R)$. Then the scheme is:

if
$$\begin{cases} \lambda_{i} > 0 \\ \lambda_{i} < 0 \end{cases}$$
 then add $-\frac{\Delta t}{\Delta x} \lambda_{i} \alpha_{i} \underline{e}_{i}$ to
$$\begin{cases} q_{R} \\ q_{L} \end{cases}$$
 (2.5)

[For a two-dimensional problem

$$\underline{\mathbf{g}}_{t} + \underline{\mathbf{F}}_{x} + \underline{\mathbf{G}}_{v} = 0 \tag{2.6}$$

we can calculate $\widetilde{A}(\underline{q}_L,\underline{q}_R)$ & $\widetilde{B}(\underline{q}_B,\underline{q}_T)$ and solve separately in the x-direction using \widetilde{A} and in the y-direction using \widetilde{B} . These solutions may then be combined in various ways, [see Strang (1968), Barley (1987) for discussions on operator splitting].

The basic scheme, (2.5) is only first order and takes discontinuities in its stride, by smoothing them out! It is so diffusive as to be totally inadequate for most problems, and in particular for our

applications. The problem then is to find a second order scheme, based on the philosophy previously illustrated, without losing the shock handling properties.

This is achieved by flux limiting, which is essentially done by limiting the anti-diffusive terms in the second order scheme at discontinuities, to avoid the oscillations of classical second order schemes, while retaining second order accuracy in smooth regions of the flow, see Sweby (1984,1985).

Define $\phi_{j'i+1/2}$ to be the signal from the j^{-+} eigenvalue at the jump at i+1/2, i.e.

$$\phi_{j},_{i+1/2} = \frac{-\Delta t}{\Delta x} \lambda_{j} \alpha_{j} \underline{e}_{j}. \tag{2.7}$$

The algorithm (2.5) then becomes:

$$\text{if} \quad \begin{cases} \lambda_{\mathbf{j}} > 0 \\ \lambda_{\mathbf{j}} < 0 \end{cases} \quad \text{then add} \quad \phi_{\mathbf{j'i+1}} \quad \text{to} \quad \begin{cases} q_{\mathbf{i+1}} \\ q_{\mathbf{i}} \end{cases}.$$

If we now transfer an amount $a_{j'i+1/2}$ against the direction of the flow we can achieve second order accuracy in smooth regions by choosing

 $a_j = \frac{1}{2}(1-|v_j|)$ where v_j is the CFL number of the j^{th} wave.

Define a transfer function, Baines (1983), by

$$B(a_{j,i+1/2}, \phi_{j,i+1/2}, a_{j,i+1/2-\sigma_j}, \phi_{j,i+1/2-\sigma_j}) = B(b_1, b_2), \text{ say,}$$
 (2.8)

where $\sigma_j = \text{sign } (\lambda_j)$. Choosing $B(b_1,b_2) = b_1$ gives Lax-Wendroff, while $B(b_1,b_2) = \frac{1}{2}(b_1+b_2)$ corresponds to Fromm's algorithm. (A more comprehensive list of the different B functions that can be used is given in the Appendix.) The two schemes above are classical second order methods and suffer from oscillations at discontinuities. We can, however, overcome this problem by not restricting ourselves to linear functions of b_1 & b_2 . Sweby (1984,1985) introduced a non-linear limiter function $\phi(r)$ where $r = b_1/b_2$. The region in $(\phi(r), r)$ space that ensures a TVD

(Total Variation Diminishing), i.e. oscillation free scheme, can then be plotted. This region can then be further constrained to ensure oscillation free second order schemes (away from extrema). Sweby (1984) plotted the minmod, Van Leer and Superbee limiters (see Roe & Baines (1982), Van Leer (1974), Roe (1985)) and showed these to lie in this region. (Two-dimensional algorithms can be constructed that are based on these limiters but also take into account the dimensionality of the problem by including cross terms to eliminate \mathbf{q}_{yx} type terms from the truncation error, see Baines (1983)). These ideas will now be used in the following sections.

3. SHALLOW WATER EQUATIONS

In primitive variables the 2-D shallow water equations relevant to atmospheric flows are given by

$$\phi_t + (\phi u)_x + (\phi v)_y = 0$$
 (3.1a)

$$u_t + uu_x + vu_y + \phi_x = \Omega v \tag{3.1b}$$

$$v_t + uv_x + vv_y + \phi_y = -\Omega u,$$
 (3.1c)

where u & v are the x,y velocities respectively and ϕ is the geopotential (ϕ =gh) where h is the height of the free surface. Ω is the Coriolis parameter which for the atmospheric flow will be given by the usual β -plane approximation and for the coastal flow problem will be taken to be a constant.

To apply Roe's scheme to these equations we need to rewrite them in conservative form. Replacing x & y velocities by x & y momenta $(m = u\phi, n = v\phi)$ we find

$$\phi_{t} + m_{x} + n_{y} = 0 \tag{3.2a}$$

$$m_{t} + \left[\frac{m^{2}}{\phi} + \frac{\phi^{2}}{2}\right]_{x} + \left[\frac{mn}{\phi}\right]_{y} = \Omega n$$
(3.2b)

$$n_{t} + \left[\frac{mn}{\phi}\right]_{x} + \left[\frac{n^{2} + \phi^{2}}{\phi}\right]_{y} = -\Omega m. \tag{3.2c}$$

This can now be written in a vector flux form as

$$\underline{q}_t + \underline{F}_x + \underline{G}_y = \underline{b}$$

where

$$\underline{q} = \begin{bmatrix} \phi \\ m \\ n \end{bmatrix}, \quad \underline{F} = \begin{bmatrix} m \\ \frac{m^2}{\phi} + \phi^2 /_2 \\ \frac{mn}{\phi} \end{bmatrix}, \quad \underline{G} = \begin{bmatrix} n \\ \frac{m \cdot n}{\phi} \\ \frac{n^2}{\phi} + \frac{\phi^2}{2} \end{bmatrix}$$

and
$$\underline{\mathbf{b}} = \begin{bmatrix} \mathbf{0} \\ \mathbf{\Omega} & \mathbf{n} \\ -\mathbf{\Omega} \mathbf{m} \end{bmatrix}$$
.

Following Roe (1981) we define an intermediate, or parameter, vector

$$\underline{\mathbf{w}} = \begin{bmatrix} \mathbf{w}_1 \\ \mathbf{w}_2 \\ \mathbf{w}_3 \end{bmatrix} = \begin{bmatrix} \phi^{1/2} \\ \phi^{1/2} \mathbf{u} \\ \phi^{1/2} \mathbf{v} \end{bmatrix}$$

and now, expressing the vectors \underline{q} , \underline{F} & \underline{G} in terms of the intermediate vector we obtain

$$\underline{\mathbf{q}} = \begin{bmatrix} \mathbf{w}_{1}^{2} \\ \mathbf{w}_{1} \mathbf{w}_{2} \\ \mathbf{w}_{1} \mathbf{w}_{3} \end{bmatrix}, \quad \underline{\mathbf{F}} = \begin{bmatrix} \mathbf{w}_{1} \mathbf{w}_{2} \\ \mathbf{w}_{2}^{2} + \frac{\mathbf{w}_{1}}{2} \\ \mathbf{w}_{2} \mathbf{w}_{3} \end{bmatrix}, \quad \underline{\mathbf{G}} = \begin{bmatrix} \mathbf{w}_{1} \mathbf{w}_{3} \\ \mathbf{w}_{2} \mathbf{w}_{3} \\ \mathbf{w}_{3}^{2} + \frac{\mathbf{w}_{1}}{2} \end{bmatrix}.$$

Using the standard notation of $\Delta \underline{x} = \underline{x}_R - \underline{x}_L$ and $\underline{x} = \frac{1}{2}(\underline{x}_R + \underline{x}_L)$ we proceed to calculate matrices $B(\underline{w})$ & $C(\underline{w})$ such that

$$\Delta \underline{\mathbf{q}} = \mathbf{B}(\underline{\mathbf{w}}) \Delta \underline{\mathbf{w}}$$

$$\Delta \underline{\mathbf{F}} = \mathbf{C}(\underline{\mathbf{w}}) \Delta \underline{\mathbf{w}}.$$

This leads to

$$B = \begin{bmatrix} 2\bar{w}_{1} & 0 & 0 \\ -\bar{w}_{2} & -\bar{w}_{1} & 0 \\ -\bar{w}_{3} & 0 & -\bar{w}_{1} \end{bmatrix},$$

$$C = \begin{bmatrix} \overline{w}_{2} & \overline{w}_{1} & 0 \\ 2\overline{w}_{1}\overline{w}_{1}^{2} & 2\overline{w}_{2} & 0 \\ 0 & \overline{w}_{3} & \overline{w}_{2} \end{bmatrix}.$$

We now find λ such that

$$\det (\lambda B - C) = 0,$$

giving

$$\lambda_{1\cdot 2\cdot 3}^{F} = \frac{\bar{w}_{2}}{\bar{w}_{1}} - \sqrt{\bar{w}_{1}^{2}} , \frac{\bar{w}_{2}}{\bar{w}_{1}} , \frac{\bar{w}_{2}}{\bar{w}_{1}} + \sqrt{\bar{w}_{1}^{2}}$$

which are seen to be the $u,u\pm\sqrt{\phi}$ characteristic speeds we would have expected.

These eigenvalues give three eigenvectors which, after multiplication by B, are given by: -

$$\underline{e}_{1},_{2},_{3} = \left\{ \begin{bmatrix} \overline{w}_{1} \\ \overline{w}_{2} - \overline{w}_{1} \sqrt{\overline{w}_{1}^{2}} \\ \overline{w}_{3} \end{bmatrix}, \begin{bmatrix} 0 \\ 0 \\ \overline{w}_{1} \end{bmatrix}, \begin{bmatrix} \overline{w}_{1} \\ \overline{w}_{2} + \overline{w}_{1} \sqrt{\overline{w}_{1}^{2}} \\ \overline{w}_{3} \end{bmatrix} \right\}.$$

It is noted here for future reference that, if regarded as a 3x3 matrix, the inverse of the above is given by:-

$$\frac{1}{2\bar{w}_{1}^{2}\sqrt{\bar{w}_{1}^{2}}} \begin{bmatrix} \bar{w}_{2} + \bar{w}_{1}\sqrt{\bar{w}_{1}^{2}} & -\bar{w}_{1} & 0 \\ -2\bar{w}_{3}\sqrt{\bar{w}_{1}^{2}} & 0 & 2\bar{w}_{1}\sqrt{\bar{w}_{1}^{2}} \\ -(\bar{w}_{2} - \bar{w}_{1}\sqrt{\bar{w}_{1}^{2}}) & \bar{w}_{1} & 0 \end{bmatrix}.$$
(3.3)

Three α 's are now found such that

$$\sum_{i} \alpha_{i} \underline{e}_{i} = \underline{A}\underline{q} \quad \text{(and by construction } \sum_{i} \alpha_{i} \lambda_{i} \underline{e}_{i} = \underline{A}\underline{F} \text{), i.e.}$$

$$\left\{ \begin{array}{c} \alpha_{1}^{F} \\ \alpha_{2}^{F} \\ \alpha_{3}^{F} \end{array} \right\} = \left\{ \begin{array}{c} \Delta w_{1} - \frac{\overline{w}_{1} \Delta w_{2} - \overline{w}_{2} \Delta w_{1}}{2\overline{w}_{1} \sqrt{\overline{w}_{1}^{2}}} \\ -\frac{\overline{w}_{1} \Delta w_{3} - \overline{w}_{3} \Delta w_{1}}{\overline{w}_{1}} \\ -\frac{\overline{w}_{1} \Delta w_{2} - \overline{w}_{2} \Delta w_{1}}{\overline{w}_{1}} \\ \Delta w_{1} + \frac{\overline{w}_{1} \Delta w_{2} - \overline{w}_{2} \Delta w_{1}}{2\overline{w}_{1} \sqrt{\overline{w}_{1}^{2}}} \right\}.$$

Performing similar operations in the y-direction gives

$$\lambda_{1,2,3}^{G} = \frac{\overline{w}_{3}}{\overline{w}_{1}} - \sqrt{\overline{w}_{1}^{2}}, \frac{\overline{w}_{3}}{\overline{w}_{1}}, \frac{\overline{w}_{3}}{\overline{w}_{1}} + \sqrt{\overline{w}_{1}^{2}} \quad \text{with}$$

$$\begin{bmatrix}
G \\
\underline{e}_{1,2,3} \\
\underline{e}_{1,2,3}
\end{bmatrix} = \begin{bmatrix}
\overline{w}_{1} \\
\overline{w}_{2} \\
\overline{w}_{3} - \overline{w}_{1} \sqrt{\overline{w}_{1}^{2}}
\end{bmatrix} \begin{bmatrix}
O \\
\overline{w}_{1} \\
\overline{w}_{2}
\end{bmatrix} \begin{bmatrix}
\overline{w}_{1} \\
\overline{w}_{2} \\
\overline{w}_{3} + \overline{w}_{1} \sqrt{\overline{w}_{1}^{2}}
\end{bmatrix}$$

and

$$\begin{cases} \alpha_{1}^{G} \\ \alpha_{2}^{G} \\ \alpha_{3}^{G} \end{cases} = \begin{cases} \Delta w_{1} - \frac{\overline{w}_{1} \Delta w_{3} - \overline{w}_{3} \Delta w_{1}}{2\overline{w}_{1} \sqrt{\overline{w}_{1}^{2}}} \\ \frac{\overline{w}_{1} \Delta w_{2} - \overline{w}_{2} \Delta w_{1}}{\overline{w}_{1}} \\ \Delta w_{1} + \frac{\overline{w}_{1} \Delta w_{3} - \overline{w}_{3} \Delta w_{1}}{2\overline{w}_{1} \sqrt{\overline{w}_{1}^{2}}} \end{cases}.$$

These eigenvalues, eigenvectors and wave strengths are used to decompose the problem into wave type components to which algorithm (2.5) can then be applied.

4. <u>MODEL PROBLEMS</u>

In this section two model problems will be discussed that will force us to select different aspects of the method to ensure a worthwhile solution.

(i) The first is an atmospheric flow using the equations in the form given in (3.2). It is known as Grammeltvedt's problem with initial conditions 1 (see Grammeltvedt (1969)) and concerns a flat bottomed channel 6000 km x 4400 km. The north/south boundaries are taken to be rigid walls and the flow is assumed peridic in the east/west direction.

The initial conditions for the height are shown in Figure 1. The velocities are obtained from the geostrophic approximation.

This problem is particularly useful because, although the exact solution is unknown, it is quite straightforward in that the geometry is rectangular, the boundary conditions are easy to write down, and owing to the nature of the problem the available energy, given by

AE =
$$\frac{1}{2g} \int \{ (u^2 + v^2 - \phi)\phi - \overline{\phi}^2 \} d\sigma$$
, (4.1)

where $\bar{\phi}$ is the average value of the geopotential of the free surface and σ is the domain, is conserved. Hence, by monitoring this quantity, we can measure the success of our efforts without knowing the analytic solution to the problem. It is worth just noting here that conserving 100% of the available energy is necessary for an exact solution but it is not a sufficient condition.

A numerical scheme is deemed to have become unstable when the available energy has risen by 10%. With some of the results the comments "going unstable" or "about to go unstable" have been added, which indicates that although we have not reached the 10% criterion the available energy for that particular method had started to increase, to a greater or lesser extent, and would, by our definition, have become unstable very rapidly thereafter.

A 200 km grid (30 x 22) is used and time-steps of 5 minutes are taken. The results in Table 1 correspond to the percentage of the available energy left after 5 days.

First Order 3.9% :--Lax-Wendroff 60 % (going unstable) 2nd order fully upwinded :-42 % 43.5% Fromm's algorithm 3rd order split : --45 % Minmod 23 % 66.3% (going unstable) Superbee :-Minmod (2-D) :-32.3%

Table 1

As can be seen these results do lack a certain something, namely rather a lot of energy. The schemes that come out best are the more classical second order methods, although as with all these schemes the time-step must be restricted because the CFL numbers are given by $u\Delta t/\Delta x & (u+\sqrt{\phi})\Delta t/\Delta x$, and the u<u>+</u>√φ values are an order of magnitude larger than the u characteristic. The methods, though, are explicit finite differences and so this time-step restriction is not too much of a Reducing the time-step to stabilize the Lax-Wendroff and Superbee methods still only leaves us with just over 50%. 'genuinely two-dimensional' minmod scheme does, in this case, do better than its 1-D version, although there are also examples where it does fractionally worse, but it still is a long way behind the other methods and is computationally very much more expensive to calculate.

So far the right hand side of the shallow water equations, $\underline{b} = (o, \Omega n, -\Omega m)^T$, has been ignored. For the shallow water equations, in this context, the Coriolis force is a significant term of equal magnitude to the $(\phi^2/_2)_X$ & $(\phi^2/_2)_Y$ terms, the geostrophic approximation. The results in Table 1 were all obtained by evaluating the Coriolis force pointwise and using this value to update at that particular gridpoint. In the light of the previous sentence we should perhaps pay more attention to these terms.

Roe (1986) shows how to deal with source terms by considering the simplest problem of any relevance to us, the scalar wave equation

$$u_t + au_x = b(x) \tag{4.2}$$

where a is a positive constant. The initial-value problem for (4.2) with initial data $u=u_0(x)$ has the general solution

$$u(x,t) = u_0(x-at) + \frac{1}{a} \int_{x-at}^{x} b(x)dx$$
 (4.3)

In practice the integrand in (4.3) will depend on the solution u, for which we only have information at the previous time-level, and hence we would expect to approximate the integral by

Solution, calculated by the Superbee scheme using

 $= 0.00 \text{ and } D_{x} = 200 \text{ km.}$

O time-steps with Dt

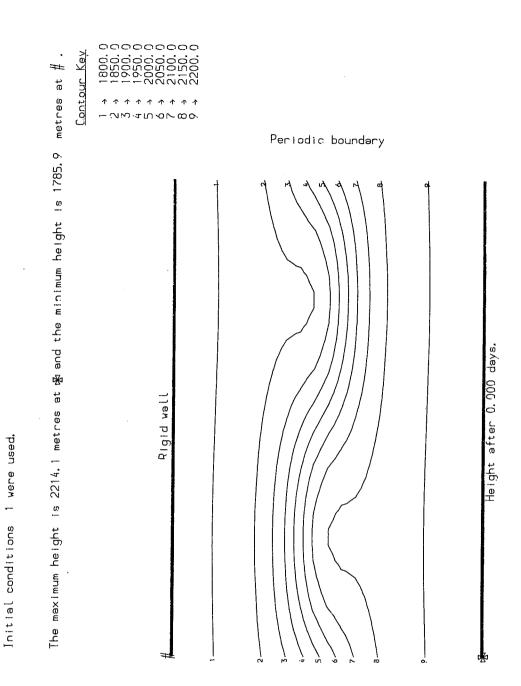


Figure 1: Initial height field.

$$\frac{1}{a}(x-(x-at)) b(x-at) = tb(x-at).$$

Hence the source term ought to be upwinded in the same way as the other terms.

Glaister (1987) has considered the 2-D Euler equations in r,z variables with cylindrical geometry. (Here the 'source' terms are due to the expansion of the grid in the radial direction.) These source terms were expanded in the eigenvectors associated with the radial direction and upwinded according to the corresponding eigenvalue. This procedure was found to significantly improve the results.

Our problem is slightly different in that in vector form the equations are

$$\underline{q}_t + \underline{F}_x + \underline{G}_v = \underline{b}$$

and we have one right hand side vector but two sets of eigenvectors in which to expand it, with no immediately clear way of splitting \underline{b} into $\underline{b}^A + \underline{b}^B$ prior to expansion in terms of the eigenvectors of A or B. One of the options is to take

$$\underline{b}^{A} = \begin{bmatrix} 0 \\ \Omega n \\ 0 \end{bmatrix} & \& \quad \underline{b}^{B} = \begin{bmatrix} 0 \\ 0 \\ -\Omega m \end{bmatrix}. \tag{4.4}$$

that is expand the source term associated with the x-momentum equation in the x-orientated eigenvectors and similarly for y. This sounds eminently reasonable but it is not at all clear from the vector equation why this should be done, and what treatment should be given to a source term in the ϕ equation which has no associated direction. Another option is to take

$$\underline{b}^{A} = \underline{b}^{B} = \frac{1}{2}\underline{b}. \tag{4.5}$$

In order to expand the forcing terms in the eigenvectors we need to know weights for each eigenvector. In the x-direction this is done by multiplying the vector to be projected, \underline{b}^{A} , by the matrix (3.3), a

similar operation being done to \underline{b}^B . This gives the weights associated with $\underline{e}_{1,2,3}^F$ & $\underline{e}_{1,2,3}^G$ as

$$\frac{\Omega \overline{w}_3}{2\sqrt{\overline{w}_1^2}} (1,0,-1) & \frac{-\Omega \overline{w}_2}{2\sqrt{\overline{w}_1^2}} (1,0,-1)$$
(4.6)

for the first option, (4.4), and

$$\frac{\Omega}{4\sqrt{\bar{w}_{1}^{2}}} \left(\bar{w}_{3}, 2\bar{w}_{2}\sqrt{\bar{w}_{1}^{2}, -\bar{w}_{3}} \right) & \frac{-\Omega}{4\sqrt{\bar{w}_{1}^{2}}} \left(\bar{w}_{2}, 2\bar{w}_{3}\sqrt{\bar{w}_{1}^{2}, -\bar{w}_{2}} \right)$$

$$(4.7)$$

for the second option (4.5).

For the first case we choose to average Ωn by $\Omega \overline{w_1} \overline{w_3}$, because it marginally simplifies the expressions in (4.6). However, the schemes do not seem very sensitive to the averaging used in the evaluation of \underline{b} , $\Omega w_1 w_3$ giving the same results. Both these options give marked increases in accuracy but it is the first that gives the best, and the results for this strategy are given below in Table 2 for some of the schemes, using 4 minute time-stepping.

First Order 72.4% L-W 85.9% (about to go unstable) 2nd order 87 % (about to go unstable) Fromm 85 % 3rd order 85.4% 77.3% Minmod 86.9% (about to go unstable) Superbee Minmod (2-D) 78 % Fromm limiter 82 % :-80 % Van Leer :-

Table 2

The effect of cross-terms becomes less significant as the time-step is reduced because they are $O(\Delta t^2)$.

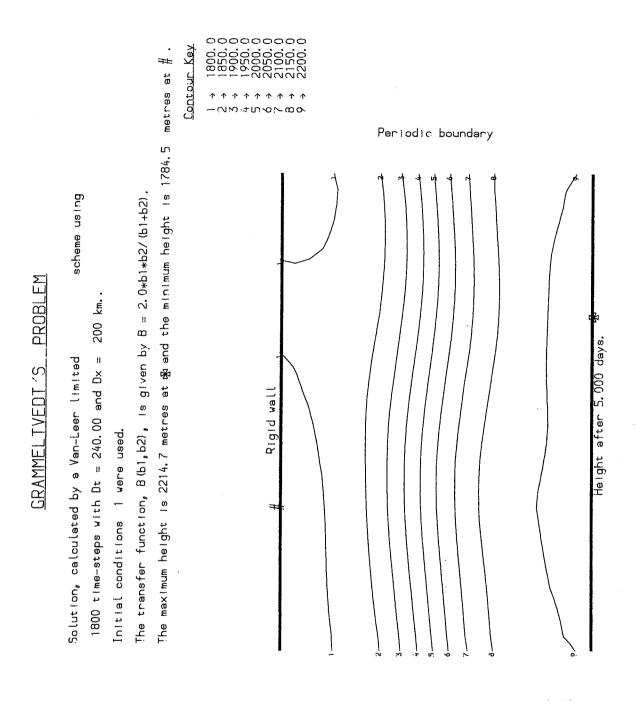


Figure 2: Predicted height field using the Van-Leer limiter.

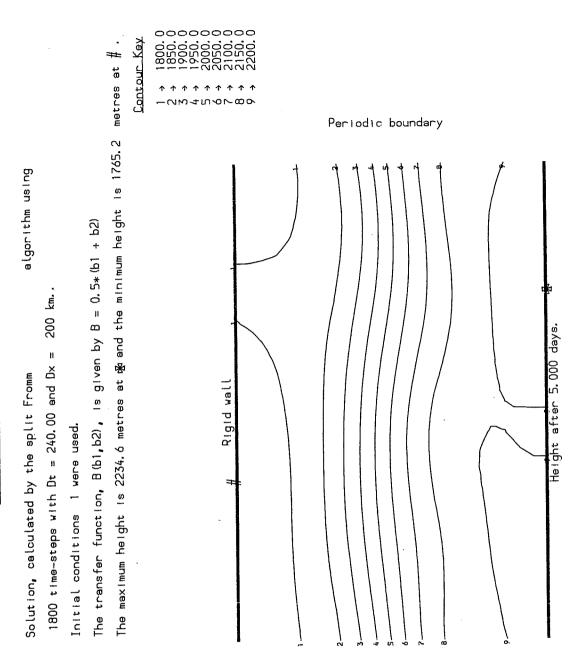


Figure 3. Predicted height field using Fromm's scheme.

Reducing the time-step to stabilize Superbee we obtain the following results

3 min :- 87.8% 2 min :- 90.4% 90 sec :- 92.8%.

2-D Superbee with the 3 min. time-stepping gives 88%, reinforcing the claim about the cross terms. These results are actually better than the classical methods give us with reduced time-stepping, but the Superbee scheme could still cope with any discontinuities that may arise. These results are a considerable improvement over our previous table, showing the importance of correctly dealing with source/forcing terms. Figure 2 shows the height for the Van Leer limiter, figure 3 shows the height for Fromm's scheme while figures 4,5,6 show the height, velocity field and available energy for Superbee with 2 minute time-stepping.

(ii) The second model problem is that of coastal flow. The equations used here are

$$u_t + uu_x + vu_y + gz_x + \frac{Fu\sqrt{u^2 + v^2}}{h + z} - \Omega v = 0$$
 (4.8a)

$$v_t + uv_x + vv_y + gz_y + \frac{Fv\sqrt{u^2 + v^2}}{h + z} + \Omega u = 0$$
 (4.8b)

$$z_t + (u(h + z))_x + (v(h + z))_y = 0,$$
 (4.8c)

where h is the depth of the sea bed, below some fixed level and z is the elevation above this level i.e. total depth is h + z. The new term is a friction term and F is the friction factor $(0(10^{-3}))$. Our conserved variables are then g(h + z), g(h + z)u, g(h + z)v, much as before. The only real differences from before are the addition of a potentially variable seabed and the addition of the friction terms to the right hand side. The Roe decomposition goes through as previously. The energy is no longer conserved, due to the presence of friction, and so

Solution, calculated by the Superbee scheme using 3600 time-steps with Dt = 120.00 and Dx = 200 km.

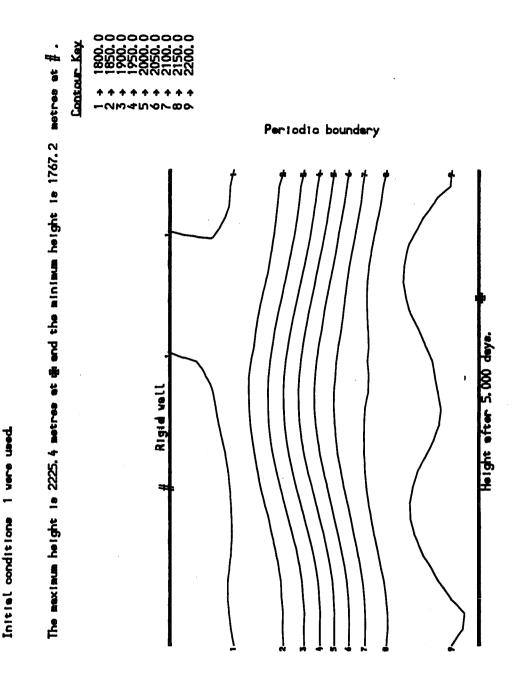


Figure 4: Predicted height field using the Superbee limiter.

GRAMMELTVEDT'S PROBLEM

Solution, calculated by the Superbee achese using 3600 time-steps with Dt = 120.00 and Dx = 200 km. Initial conditions 1 were used.

The meximum vind speed is 26.414 metres per second. (59.09 m.p.h..)

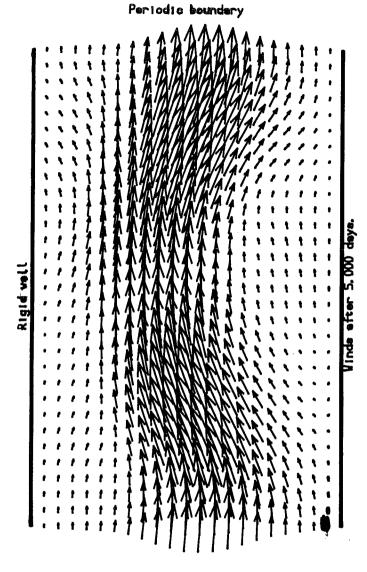


Figure 5: Predicted winds using the Superbee limiter.

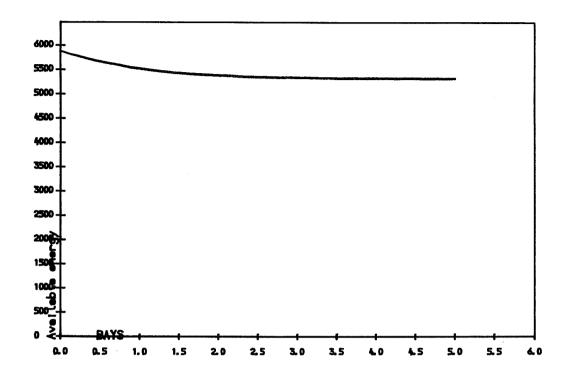


Figure 6: Conservation of energy using the Superbee limiter.

these equations possess no easy guide to the success of the approximation. We will therefore use our experience with Grammeltvedt's problem to choose the appropriate methods.

The real problem posed by this model problem is the treatment of the boundary conditions. We shall take as our problem a straight stretch of coastline 50 kms long. This presents no problems as it is treated as a rigid wall as in Grammeltvedt's problem. Although stretches of coastline 50 km long and perfectly straight are rare this is not a limit on the applicability of the scheme. Irregular domains can be transformed to a regular domain or attacked as in Priestley (1987). The real problem here is that to predict our coastal flow we do not want the expense of calculating the flow over the entire ocean and so we are forced to use computationally open boundaries.

As our model problem, then, we take a (10 x 50) km. region with the rigid wall to the west and the southern, eastern and northerly boundaries open. Assuming that the flow is sub-critical we shall need one boundary condition on outflow and two on inflow boundaries for the exact problem. Edwards, Please and Preston (1983) have discussed stable boundary conditions for the linearised shallow water equations. A problem with many finite difference techniques is that they require more boundary conditions than the mathematical problem and this can cause severe problems (see Burgess (1986). Using a scheme based on characteristics, though, means that we only need to use the correct number of boundary conditions, and those suggested by Edwards et al will be used. These are (i) elevation prescribed at outflow and (ii) elevation and tangential velocity given at inflow.

The initial conditions are taken from an exact solution to the 2-D wave equation, namely

$$u = 0 (4.9a)$$

$$v = \frac{Ag}{\sqrt{gh}} \sin \left[\frac{2\pi}{P} \left(\frac{y}{\sqrt{gh}} - t \right) \right]$$
 (4.9b)

$$z = A \sin \left[\frac{2\pi}{P} \left(\frac{y}{\sqrt{gh}} - t \right) \right]$$
 (4.9c)

where A = 1.0 m is the tidal amplitude and P = 12.42 hrs is the tidal

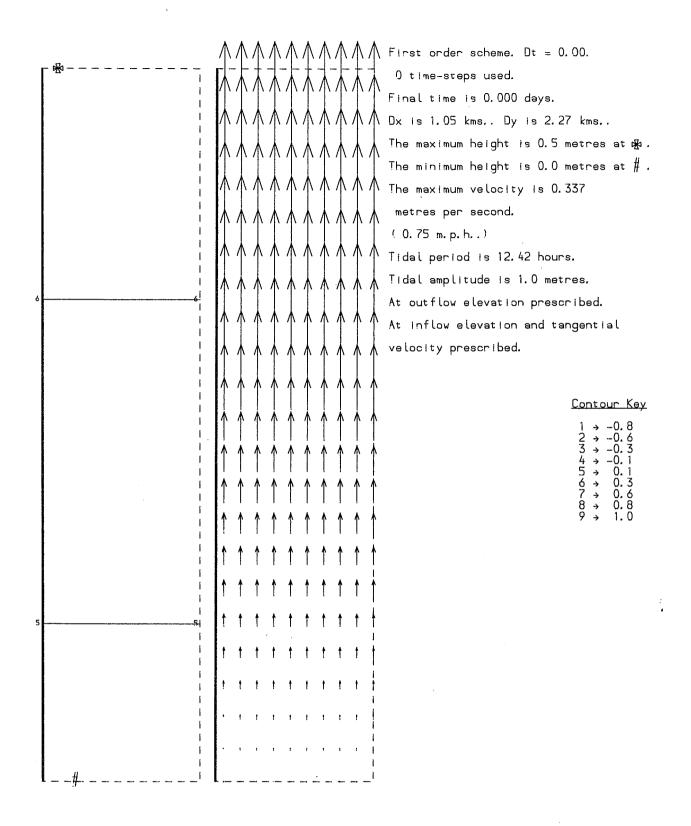


Figure 7: Initial conditions for coastal flow problem.

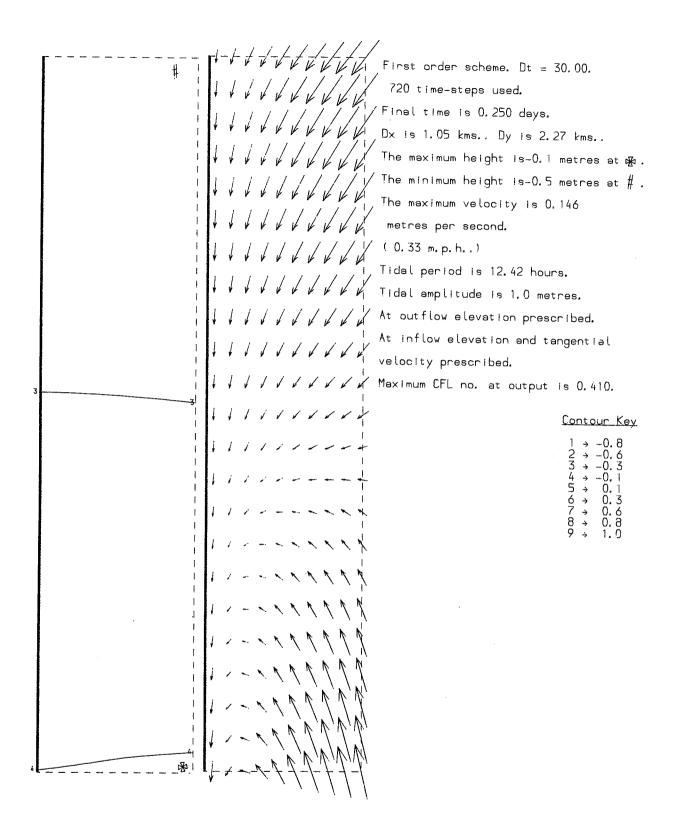


Figure 8: Predicted flow using first order upwind scheme.

E,

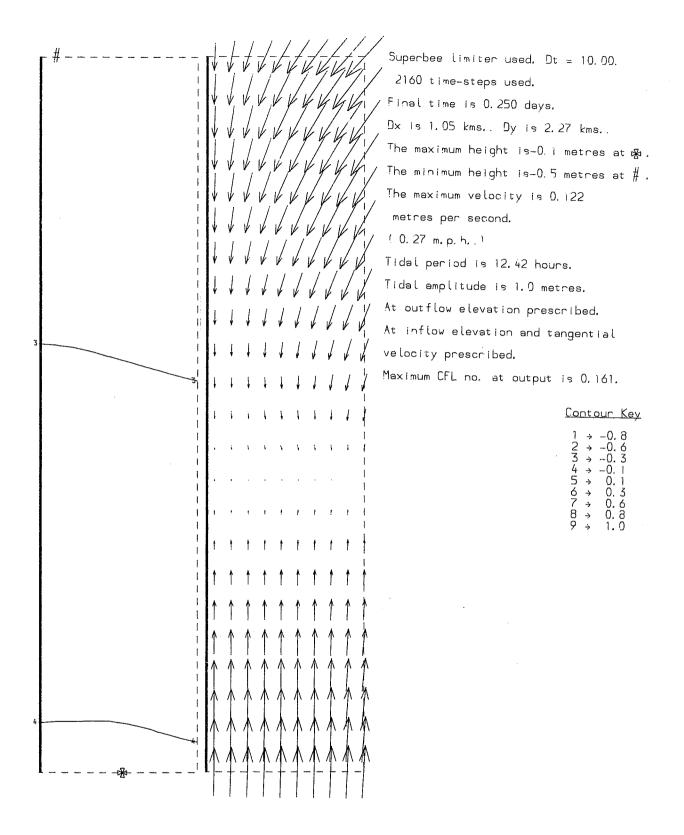


Figure 9: Predicted flow using Superbee limiter.

period. The initial values are plotted in Figure 7. Formulae 4.9 are not only used to give the initial conditions but are also used to prescribe the boundary conditions as well. Unfortunately the solution described by equations 4.9 rapidly lost touch with reality at the seaward boundary and so the region was extended to 0<x<20, -25<y<75km and the boundary conditions applied on these boundaries to try to minimize their effect on the region of interest.

Figure 8 shows the solution generated by using the first order method with 30 second time-stepping and Figure 9 shows the solution given by Superbee with 10 second time-stepping, plotted after six hours. These show quite similar results and demonstrate the stability of the Superbee algorithm in this situation.

5. CONCLUSION

It is well known that the Roe type flux limited schemes provide a very powerful tool for resolving flows, where discontinuities form a major part, without oscillations. In this paper smooth flows with substantial forcing terms have been considered and the methods have been shown to be, if anything, even more accurate than the more classical second order schemes provided that due attention is paid to the treatment of the Coriolis and friction components. It has been found in numerical experiments that for stability it is essential to include the forcing terms in the limiter function as well as just upwinding it. Duhamel's principle explains the need to decompose the forcing term within Roe's scheme, Sweby (1987), although we do not as yet have any insight into the dimensional splitting of the right hand side.

Another advantage illustrated here, of characteristic based schemes over non-characteristic based schemes is the ease with which boundary conditions may be applied and the stability of this application.

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7. <u>REFERENCES</u>

Baines, M.J., 1983: Numerical Algorithms for the Non-linear Scalar Wave Equation. University of Reading, Numerical Analysis Report 1/83.

Barley, J.J., 1987: A Survey of Operator Splitting Applied to Upwind Differencing. University of Reading Numerical Analysis Report (to appear).

Burgess, N.A., 1986: The Stability of an Approximation to the 1-D Shallow Water Equations" Oxford University Computing Laboratory Report no. 86/8.

Edwards, N.A., Please, C.P., Preston, R.W., 1983: Some Observations on Boundary Conditions for the Shallow-Water Equations in Two Space Dimensions. IMA Journal of Applied Mathematics, 30.

Glaister, P., 1985: Flux Difference Splitting Techniques for the Euler Equations in Non-Cartesian Geometry. University of Reading Numerical Analysis Report 8/85.

Glaister, P., 1987: Second Order Difference Schemes for Hyperbolic Conservation Laws with Source Terms. University of Reading Numerical Analysis Report 6/87.

Godunov, S.K., 1959: A Finite Difference Method for the Numerical Computation of Discontinuous Solutions of the Equations of Gas Dynamics, Mat Sb. 47.

Grammeltvedt, A., 1969: A Survey of Finite-Difference Schemes for the Primitive Equations for a Barotropic Fluid. Monthly Weather Review Vol. 97. No. 5.

Priestley, A., 1987: Roe's Scheme, Euler Equations, Cartesian Grids, Non-Cartesian Geometries, Rigid Walls and all that. University of Reading Numerical Analysis Report (to appear).

Roe, P.L., 1981: Approximate Riemann Solvers, Parameter Vectors, and Difference Schemes. Journal of Computational Physics, 43.

Roe, P.L., 1985: Some Contributions to the Modelling of Discontinuous Flows. Lectures in Applied Mathematics, 22, part 2. Am. Math. Soc..

Roe, P.L., 1986: Upwind Differencing Schemes for Hyperbolic Conservation Laws with Source Terms. Proc. 1st Int. Congress on Hyperbolic Problems, St. Etienne.

Roe, P.L. and Baines, M.J., 1982: Algorithms for Advection and Shock Problems. Proc. 4th GAMM Conf. on Numerical Methods in Fluid Mechanics, (Ed. H. Viviand), Vieweg.

Roe, P.L. and Pike, J., 1984: Efficient Construction and Utilisation of Approximate Riemann Solutions. Computing Methods in Applied Science and Engineering, VI, 499.

Sells, C.C.L., 1980: Solution of the Euler Equations for Transonic Flow Past a Lifting Aerofoil. R.A.E. Technical Report 80065.

Strang, G., 1968: On the Construction and Comparison of Difference Schemes. SIAM J. Numer. Anal., Vol. 5, No. 3.

Sweby, P.K., 1984: High Resolution Schemes using Flux Limiters for Hyperbolic Conservation Laws. SIAM Journal of Numerical Analysis, Vol. 21, No. 5.

Sweby, P.K., 1985: High Resolution TVD Schemes Using Flux Limiters. Lectures in Applied Mathematics, Vol. 22.

Sweby, P.K., 1987: High Resolution Oscillation free Schemes for Inhomogeneous Conservation Laws. University of Reading Numerical Analysis Report (to appear).

Van Leer, B., 1974: Towards the Ultimate Conservative Difference Scheme II. Monotonicity and Conservation combined in a Second Order Scheme. J. Computational Physics, 32.

APPENDIX

B function

)

First Order	$B(b_1,b_2) = 0$
Lax-Wendroff	$B(b_1,b_2) = b_1$
Second Order Fully upwinded	$B(b_1,b_2) = b_2$
Fromm's algorithm	$B(b_1,b_2) = \frac{1}{2}(b_1 + b_2)$
Third Order	$B(b_1,b_2) = \frac{1}{3}((2-\nu)b_1 + (1+\nu)b_2)$
	(v is the CFL no.)
Minmod	$B(b_{1},b_{2}) = \begin{cases} b_{1} & \text{if } b_{1} < b_{2} \\ b_{2} & \text{if } b_{1} > b_{2} \end{cases}$
Fromm based limiter	$B(b_1,b_2) = \begin{cases} \frac{1}{3} & \frac{b_1}{b_2} \\ 2 & \text{minmod } (b_1,b_2) \\ & \text{otherwise} \end{cases}$
	if $b_1b_2>0$ and zero otherwise
Van Leer limiter	$B(b_1,b_2) = \begin{cases} \frac{2b_1b_2}{(b_1+b_2)} & b_1b_2 > 0 \\ \end{cases}$

0

otherwise

Superbee
$$B(b_1, b_2) = \begin{cases} b_1 & 1 \le \frac{b_1}{b_2} \le 2 \\ b_2 & 1 \le \frac{b_2}{b_1} \le 2 \\ 2b_2 & \frac{b_1}{b_2} \ge 2 \\ 2b_1 & \frac{b_2}{b_1} \ge 2 \end{cases}$$

if
$$b_1b_2>0$$

0 otherwise

$$\text{Hyperbee} \qquad B(b_1b_2) = \begin{cases} b_1 & 1 \leq \frac{b_1}{b_2} \leq 2 + \frac{2\nu}{1-\nu} \\ b_2 & 1 \leq \frac{b_2}{b_1} \leq 2 + \frac{2(1-\nu)}{\nu} \\ 2 + \frac{2\nu}{1-\nu}b_2 & \frac{b_1}{b_2} \geq 2 + \frac{2\nu}{1-\nu} \\ 2 + \frac{2(1-\nu)}{\nu}b_1 & \frac{b_2}{b_1} \geq 2 + \frac{2(1-\nu)}{\nu} \end{cases}$$

if $b_1b_2 > 0$ 0 otherwise.

Ultrabee
$$\phi(r) = \frac{2r}{\nu(1-\nu)} \frac{\nu(1-r) + r(1-r^{-\nu})}{(1-r^2)}$$

$$B(b_1b_2) = b_2\phi(r) \text{ if } b_1b_2\rangle 0$$

O otherwise.

khb/10/87