



A revised ocean-atmosphere physical coupling interface

and about technical coupling software

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with numerous contributions from the community

Outline

Part I - On an revised ocean-atmosphere physical coupling interface

- Context and guidelines for the design of a new physical interface
- The physical exchanges
- Time sequence of exchanges

Part II - About technical coupling software

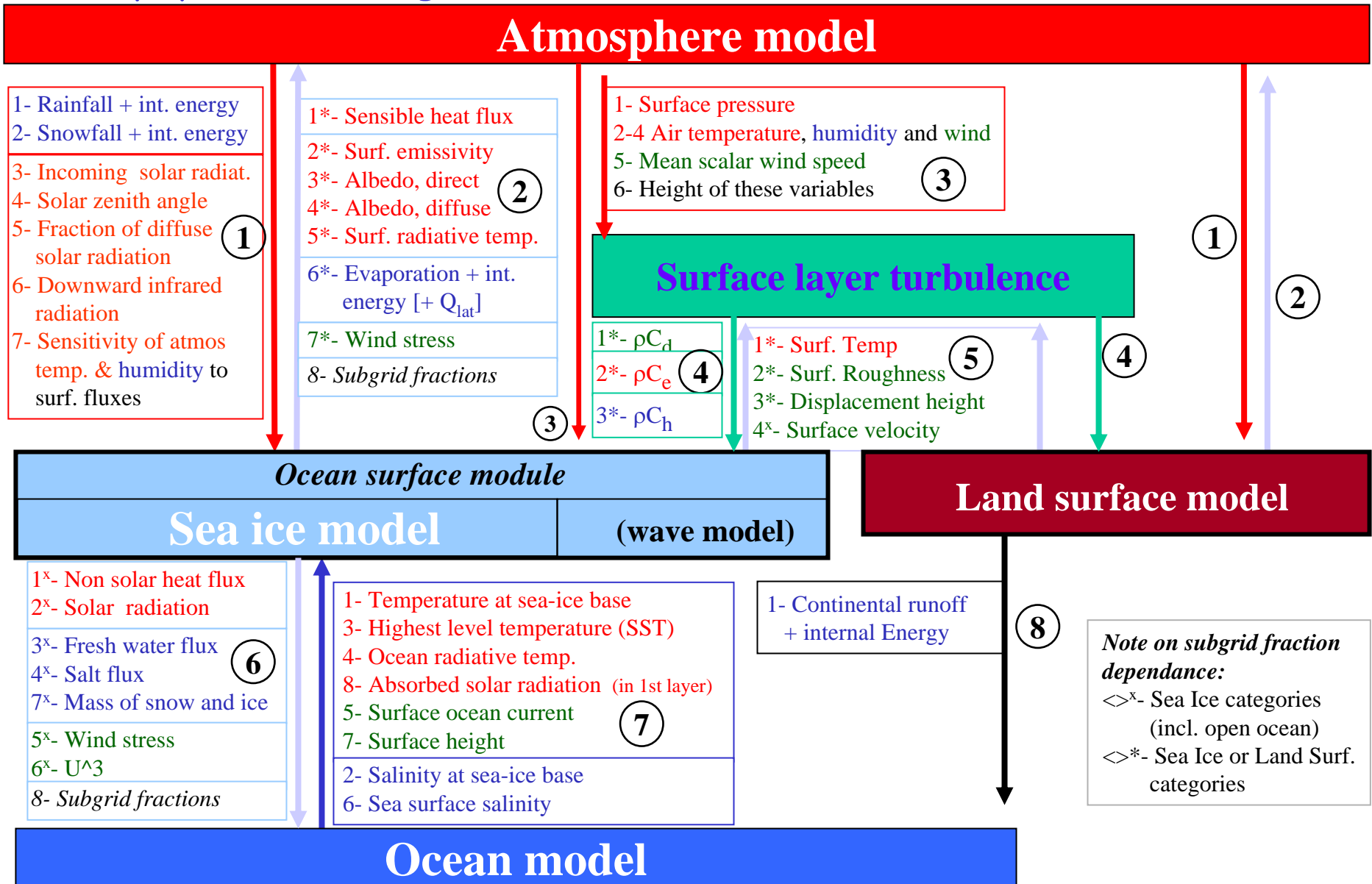
- Different technical solutions to assemble model codes
- The OASIS coupler (historic, community, ...)
- Regridding algorithms in OASIS
- 1st order conservative remapping (2nd order, SUBGRID)
- Non-matching sea-land mask
- Vector interpolation

I.1 Context and guidelines for the design of a revised interface

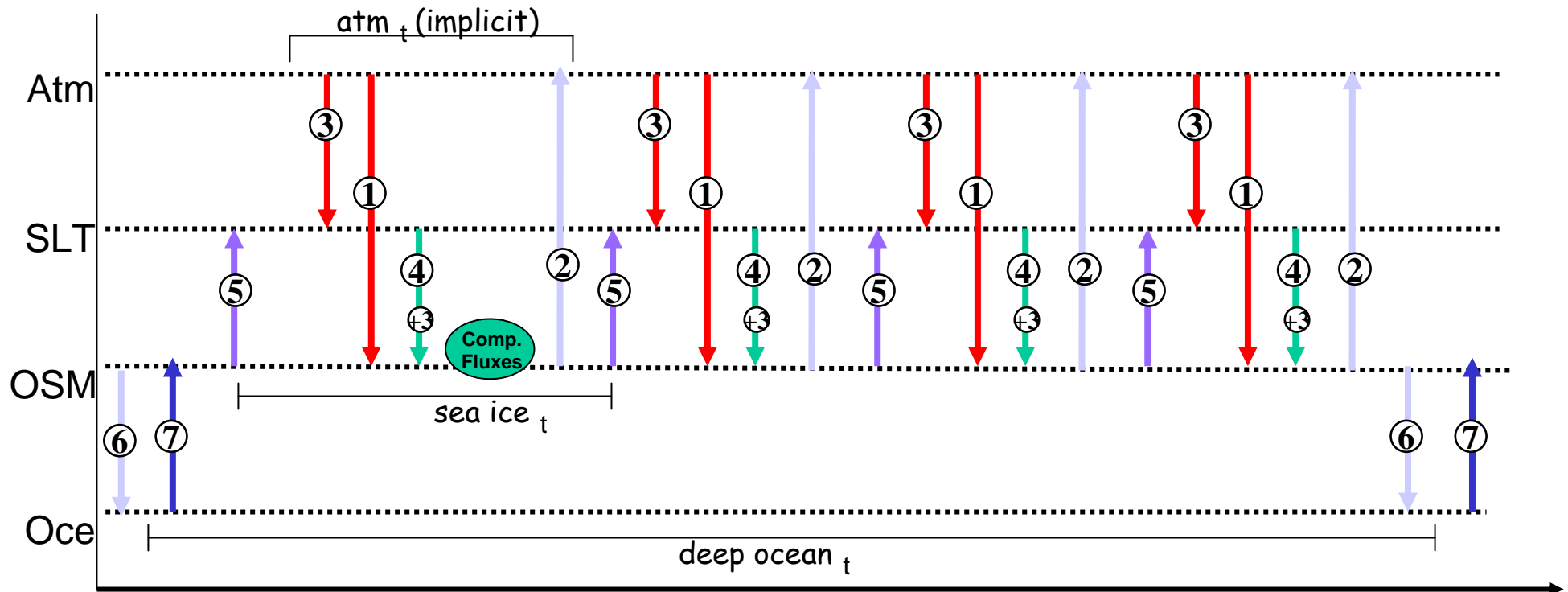
- Proposition discussed during the EU PRISM project (definition of "standard" physical interfaces), following the PILPS experience (Polcher et al 1998)
- J. Polcher (LMD), T. Fichefet (UCL), G. Madec (LOCEAN), O. Marti (LSCE), S. Planton (Meteo-France), E. Guilyardi (LOCEAN)
- Guidelines:
 - ❖ physically based interfaces across which conservation of mass, momentum and energy can be ensured
 - ❖ which process should be computed by which component/module
 - ❖ numerical constraints (stability, regridding, subgrid issues, local conservation,...)
 - ❖ historical and practical constraints

Part I - On a revised ocean-atmosphere physical coupling interface

I.2 The physical exchanges



I.3 Time sequence of exchanges



Frequency of coupling exchanges:

$$\underbrace{F_7 = F_6}_{\text{slow}} < \underbrace{F_5 = F_3 = F_1 = F_4 = F_2}_{\text{fast}}$$

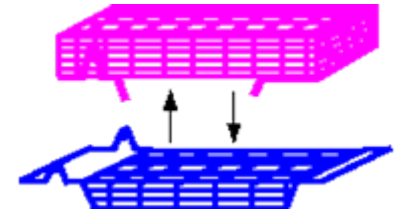
Comments and conclusions

- Increased modularity with SLT and OS modules.
- SLT runs on finer grid and computes surface turbulent coefficient.
- OS computes radiation and turbulent fluxes.
 - ✓ Separation of fast ocean + sea ice surface processes involving heat, water and momentum exchanges with the atmosphere from slower deeper ocean processes.
 - ✓ Calculation of fluxes at the resolution of the surface (would be non-physical to regrid the turbulent exchange coefficients C_d, C_e, C_h).
 - ✓ Implicit calculation of energy fluxes from the base of the sea-ice to the top of the atmosphere.

Part II - About technical coupling software

Why couple ocean and atmosphere (and sea-ice and land and ...) models?

- Of course, to treat the Earth System globally



What does “coupling of codes” imply?

- Exchange and transform information at the code interface
- Manage the execution of the codes

What are the constraints?

- ✓ The coupling should be easy to implement
- ✓ The coupling should be flexible
- ✓ The coupling should be efficient
- ✓ The coupling should be portable
- ✓ We start from independently developed existing codes

II.1 Different technical solutions to assemble model codes:

1. merge the codes:

```
program prog1  
...  
call sub_prog2(data)  
...  
end prog1
```



```
program prog2  
subroutine sub_prog2(data)  
...  
end prog2
```

- ☹ ~~easy~~
- ☹ ~~flexible~~
- ☺ efficient
- ☺ portable
- ☹ ~~existing codes~~

2. use existing communication protocols (MPI, CORBA, UNIX pipe, files, ...)

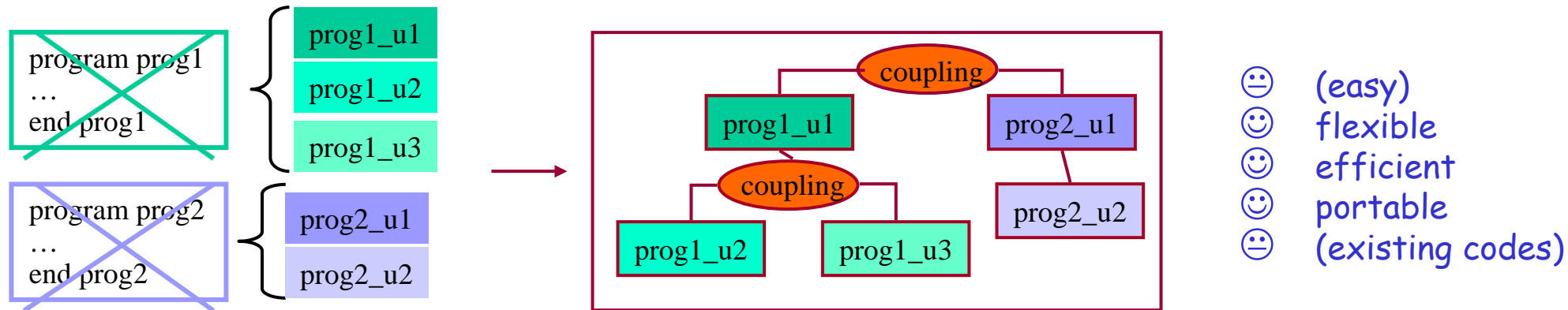
```
program prog1  
...  
call xxx_send (prog2, data, ...)  
end
```

```
program prog2  
...  
call xxx_recv (prog1, data)  
end
```

- ☹ ~~easy~~
- ☹ ~~flexible~~
- ☹ (efficient)
- ☹ (portable)
- ☺ existing codes

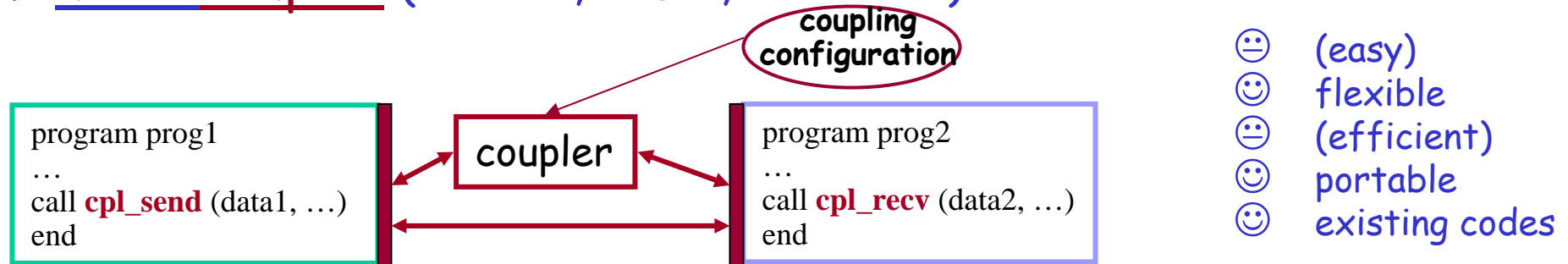
3. use coupling framework (ESMF, FMS, ...)

- Split code into elemental units
- Write or use coupling units
- Use the framework to build and control a **hierarchical merged code**
- Adapt code data structures



→ probably best solution in a controlled development environment

4. use a coupler (OASIS, PALM, MPCCI ...)

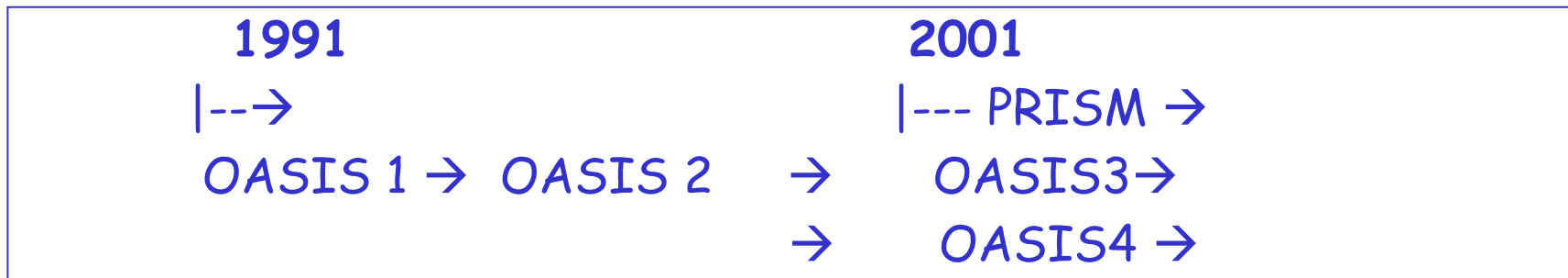


→ probably best solution to couple independently developed codes



II.2 The OASIS coupler

- ❖ developed by CERFACS since 1991 to couple existing GCMs
- ❖ currently an active collaboration between NLE-IT, CNRS and CERFACS



OASIS1, OASIS2, OASIS3:

- low resolution, low number of 2D fields, low coupling frequency:
 - flexibility very important, efficiency not so much!
 - ❖ New OASIS3_3 release in the next few weeks!

OASIS4:

- high resolution parallel models, massively parallel platforms, 3D fields
 - need to optimise and parallelise the coupler
 - ❖ OASIS4 beta version available

Part II - About technical coupling software

II.2.1 OASIS community today

•CERFACS (France)	ARPEGE3-ORCA2/LIM, ARPEGE4-NEMO/LIM-TRIP
•METEO-FRANCE (France)	ARPEGE4-ORCA2, ARPEGE3-OPamed ARPEGE3-OPA8.1/GELATO
•IPSL- LODYC, LMD, LSCE (France)	LMDz-ORCA2/LIM LMDz-ORCA4 ORCA4
•MPI-M&D (Germany)	ECHAM5-MPI-OM, ECHAM5-C-HOPE, PUMA-C-HOPE, EMAD-E-HOPE
•ECMWF	IFS - CTM (GEMS), IFS - ORCA2 (MERSEA)
•MET Office (UK)	MetOffice ATM - NEMO
•IFM-GEOMAR (Germany)	ECHAM5 - NEMO (OPA9-LIM)
•NCAS / U. Reading (UK)	ECHAM4 - ORCA2 HADAM3-ORCA2
•SMHI (Sweden)	RCA(region.) - RCO(region.)
•NERSC (Norway)	ARPEGE - MICOM, CAM - MICOM
•KNMI (Netherlands)	ECHAM5 - TM5/MPI-OM
•INGV (Italy)	ECHAM5 - MPI-OM
•ENEA (Italy)	MITgcm - REGgcm
•JAMSTEC (Japan)	ECHAM5(T106) - ORCA $\frac{1}{2}$ deg
•IAP-CAS (China)	AGCM - LSM
•KMA (Korea)	CAM3 - MOM4
•BMRC (Australia)	BAM3-MOM2, BAM5-MOM2, TCLAPS-MOM
•CSIRO (Australia)	Sea Ice code - MOM4
•RPN-Environment Canada (Canada)	MEC - GOM
•UQAM (Canada)	GEM - RCO
•U. Mississippi (USA)	MM5 - HYCOM
•IRI (USA)	ECHAM5 - MOM3
•JPL (USA)	UCLA-QTCM - Trident-Ind4-Atlantic

II.3 Regridding algorithms available in OASIS

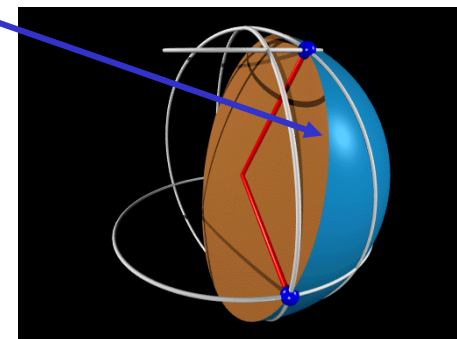
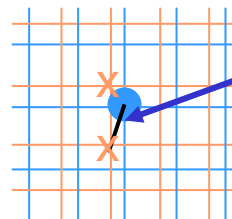
(Los Alamos SCRIP library, Jones 1999)

x: source grid point
● target grid point

- n-nearest-neighbours: $\text{weight}(x) \propto 1/d$

d: great circle distance on the sphere:

$$d = \arccos[\sin(\text{lat1}) * \sin(\text{lat2}) + \cos(\text{lat1}) * \cos(\text{lat2}) * \cos(\text{lon1} - \text{lon2})]$$

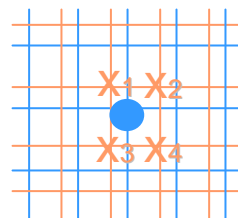


- gaussian weighted n-neighbours: $\text{weight}(x) \propto \exp(-1/2 d^2/\sigma^2)$

- bilinear interpolation

➤ general bilinear iteration in a continuous local coordinate system

using $f(x)$ at x_1, x_2, x_3, x_4



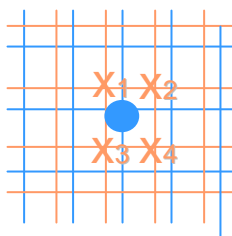
- bicubic interpolation: conserves 2nd order properties such as wind curl

➤ general bicubic iteration in a continuous local coordinate system:

$f(x), \delta f(x)/\delta i, \delta f(x)/\delta j, \delta^2 f/\delta i \delta j$ in

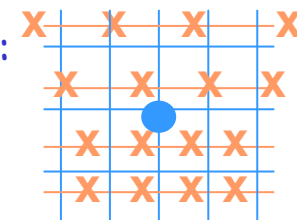
x_1, x_2, x_3, x_4

for logically-rectangular grids (i,j)



➤ standard bicubic algorithm:
16 neighbour points

for Gaussian Reduced grids

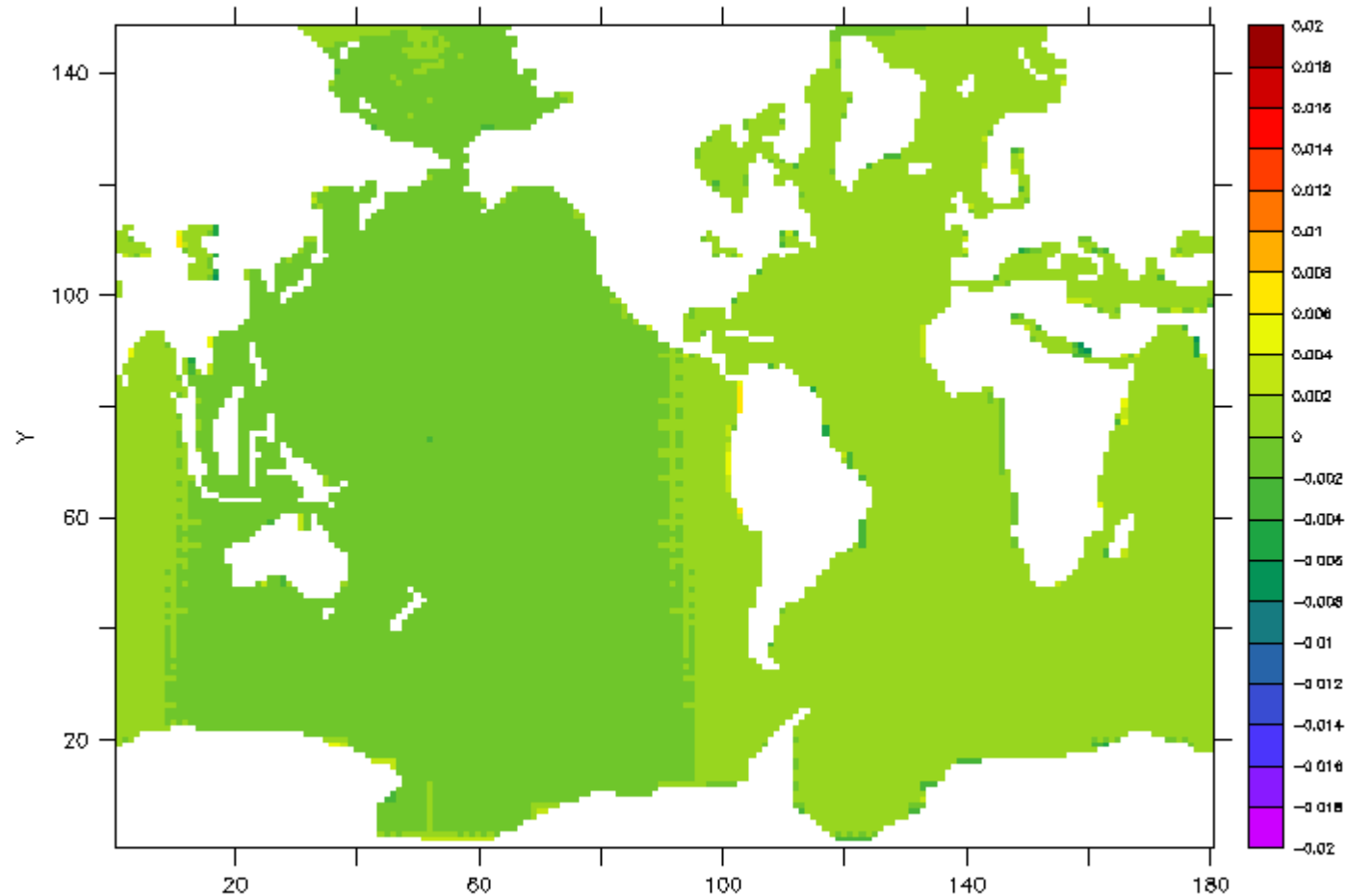
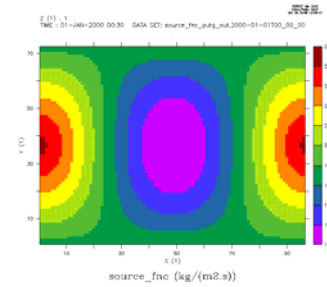


Part II - About technical coupling software

One example of bilinear interpolation error

$$F = 2 + \cos[\pi * \text{acos}(\cos(\text{lon})\cos(\text{lat}))]$$

LMDz grid (96 x 72) → ORCA2

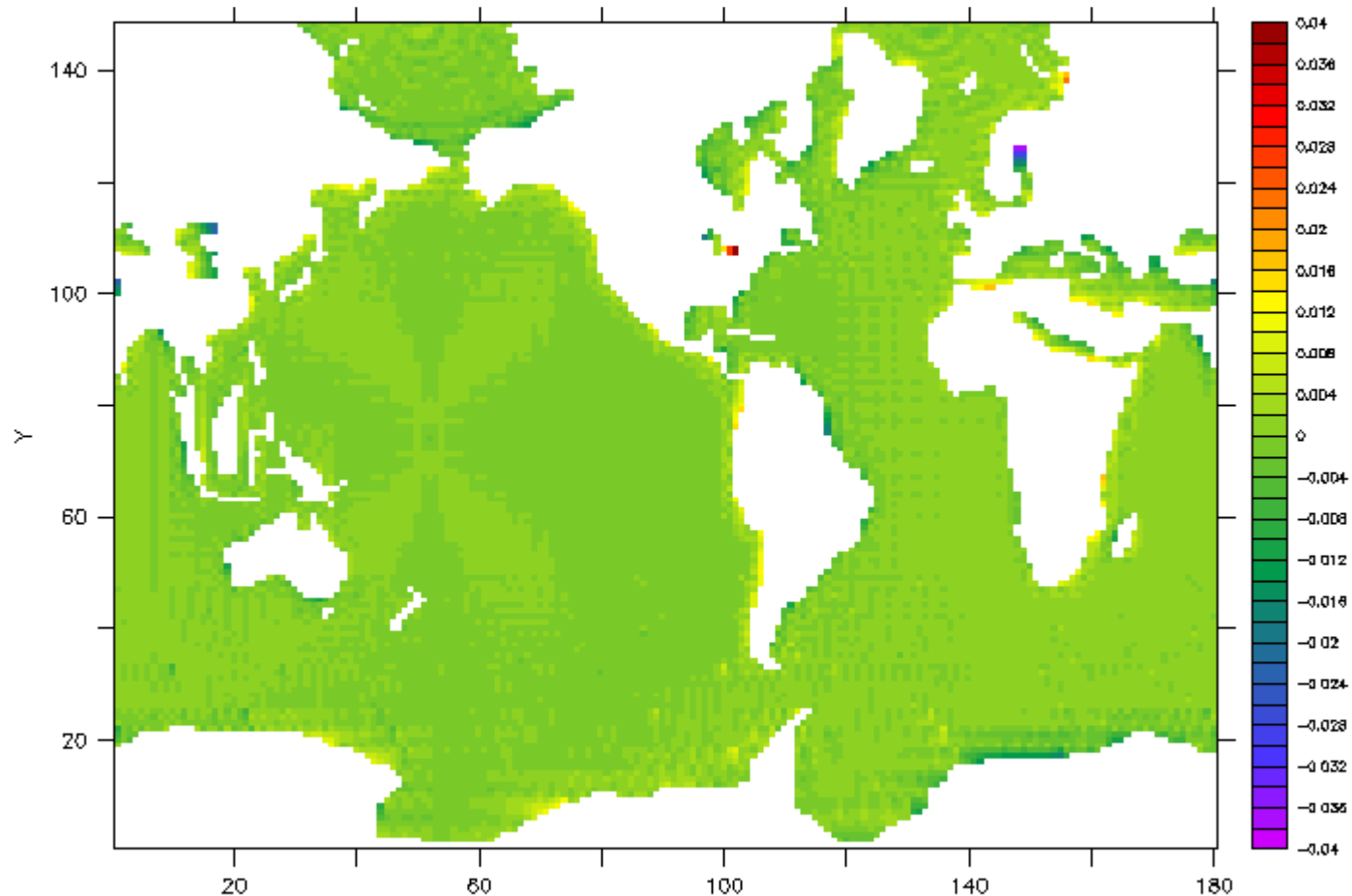


➤ < 0.2% whole domain; ~1% near the coastline

Part II - About technical coupling software

- One example of bicubic interpolation error

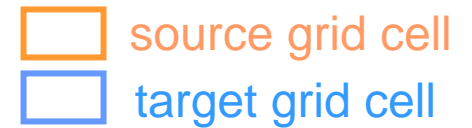
$F = 2 - \cos[\pi * \text{acos}(\cos(\text{lon})\cos(\text{lat}))]$
BT42 Gaussian red. -> ORCA2



- < 0.2% in equatorial and tropical regions,
< 0.4% at higher or lower latitudes (where the Gaussian grid is effectively reduced),
up to 4% near the coastline

II.3 Regridding algorithms available in OASIS

(Los Alamos SCRIP library, Jones 1999)



□ source grid cell

□ target grid cell

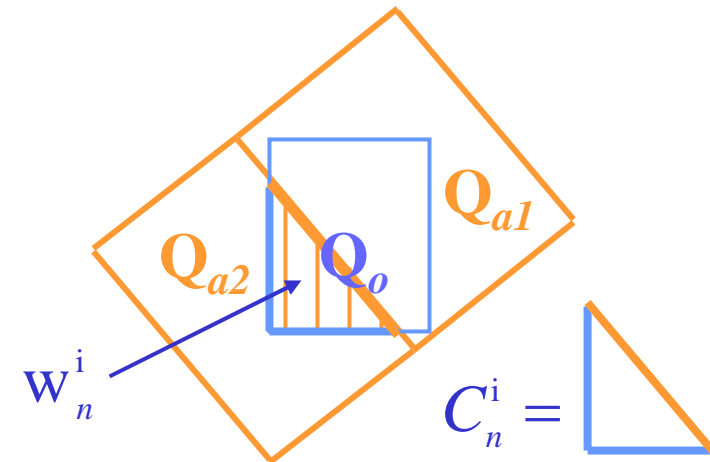
• 1st order conservative remapping:

- conserves integral of extensive properties
- weight of a source cell α to intersected area

$$Q_o^i = \frac{1}{A_o} \sum_{n=1}^N Q_{a_n} w_n^i \quad \text{with} \quad w_n^i = \oint_{C_n^i} -\sin(\text{lat}) d\text{lon}$$

❖ assumes borders are linear in (lat,lon)

- Lambert equivalent azimuthal projection near the pole for intersec. calc.



Actual limitations:

- assumes $\sin(\text{lat})$ linear function of lon (for line integral calculation)
 - need to use a projection near the pole (as done for intersect. calc.)
- exact calculation is not possible as "real shape" of the borders are not known
 - could use of border middle point
 - to ensure conservation, need to normalize by true area of the cells

Other methods e.g.:

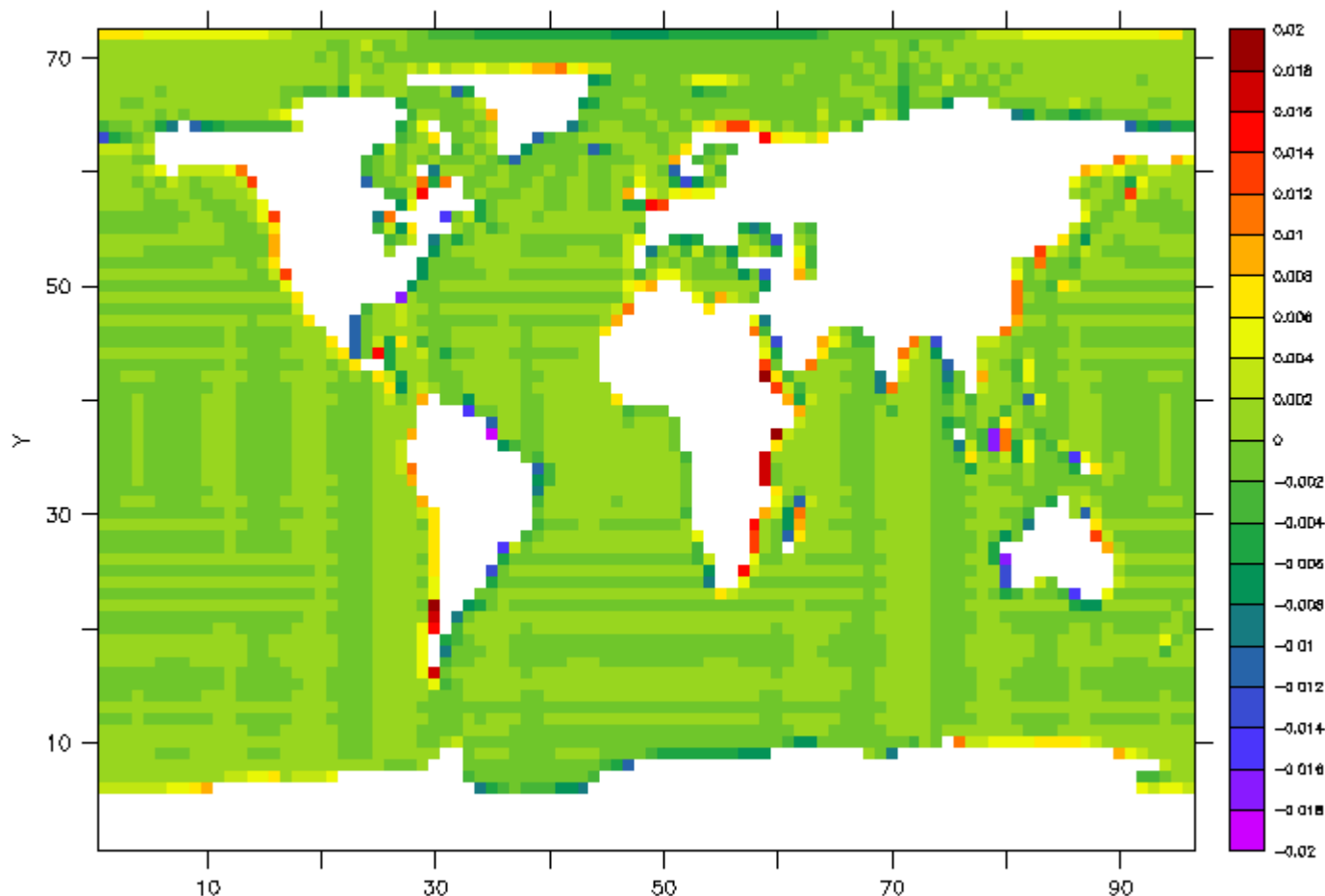
- Monte Carlo random walk
- Projection of the source and target polygons on a plane (IPSL)

Part II - About technical coupling software

- One example of conservative remapping error

$$F = 2 - \cos[\pi * \text{acos}(\cos(\text{lon})\cos(\text{lat}))]$$

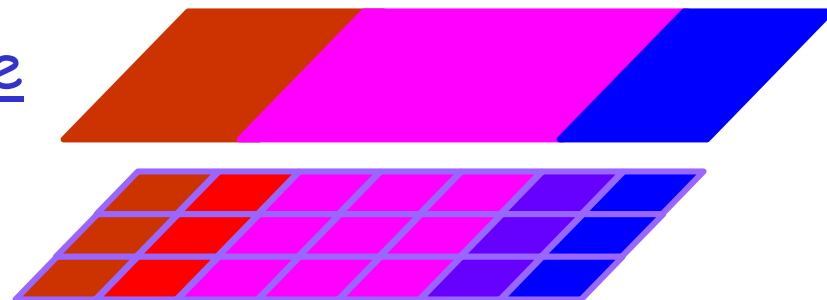
ORCA2 → LMDz (96x72)



- < 0.2% everywhere except
~ 0.8% for LMDz last row close to the North pole
~ 2% near the coastline

II.4 Problem with 1st order conservative remapping

(low to high resolution):



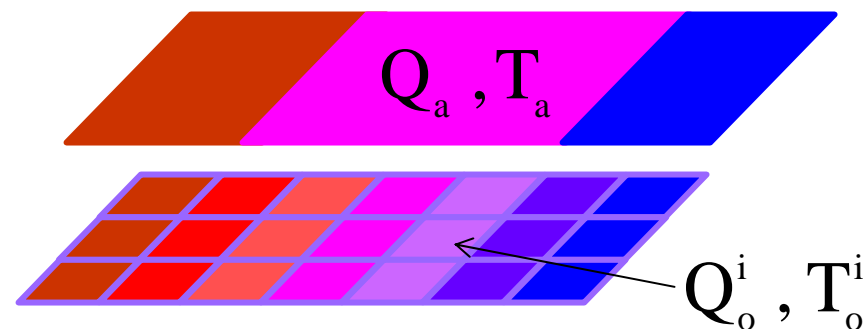
- Solution 1: use 2nd order conservative remapping:

$$Q_o^i = Q_a w_1^i + \frac{\partial Q_a}{\partial lat} w_2^i + \frac{1}{\cos(lat)} \frac{\partial Q_a}{\partial lon} w_3^i$$

- Solution 2: use SUBGRID transformation:

Solar type:
$$Q_o^i = \frac{(1 - \alpha_o^i)}{(1 - \alpha_a)} Q_a$$

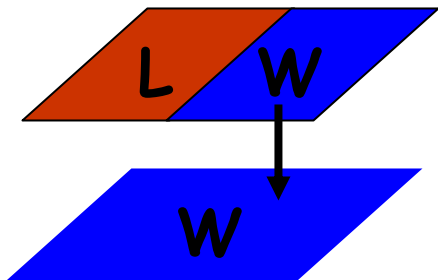
Non-solar type:
$$Q_o^i = Q_a + \left. \frac{\partial Q_a}{\partial T} \right|_{T=T_a} (T_o^i - T_a)$$



*conservative if α_a / T_a correspond to conservative remapping of α_o^i / T_o^i

II.5 Problem with non-matching sea-land masks $Q_o^i = \frac{1}{A_o} \sum_{n=1}^N Q_{a_n} W_n^i$

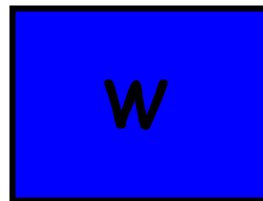
1- Ideally: Support subsurfaces in the atmosphere and use the ocean land-sea mask in the atmosphere to determine the fractional area of each type of surface



2- "DESTAREA" option

- local flux conservation
- possibly unrealistic flux values

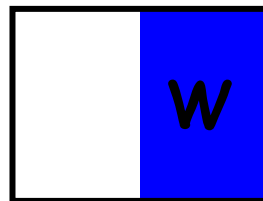
$$A_o =$$



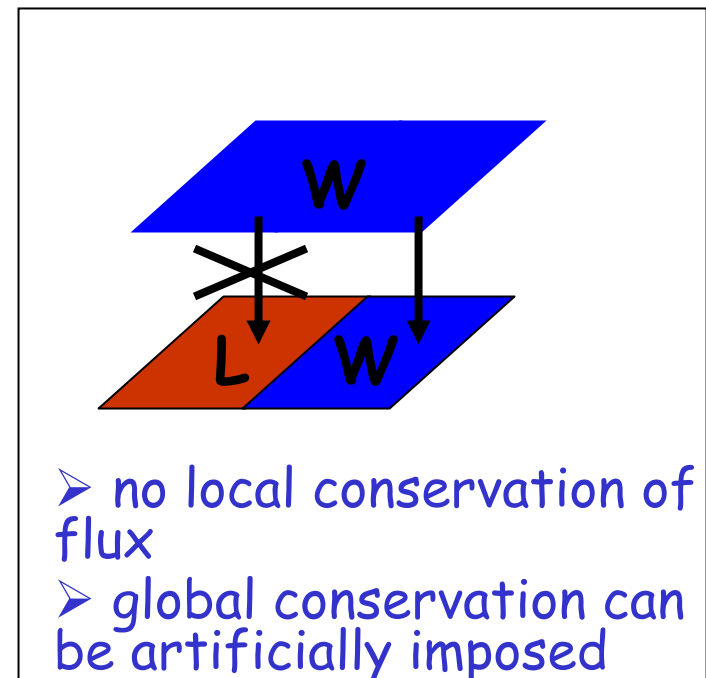
3- "FRACAREA" option

- no local conservation of flux
- realistic flux values

$$A_o =$$



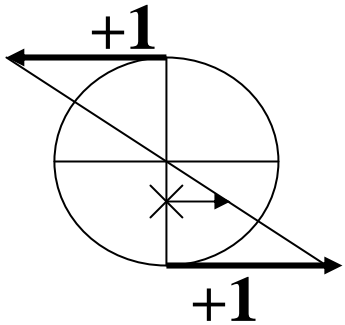
+ nearest non-masked value for ocean cells covered only with masked atmospheric cells



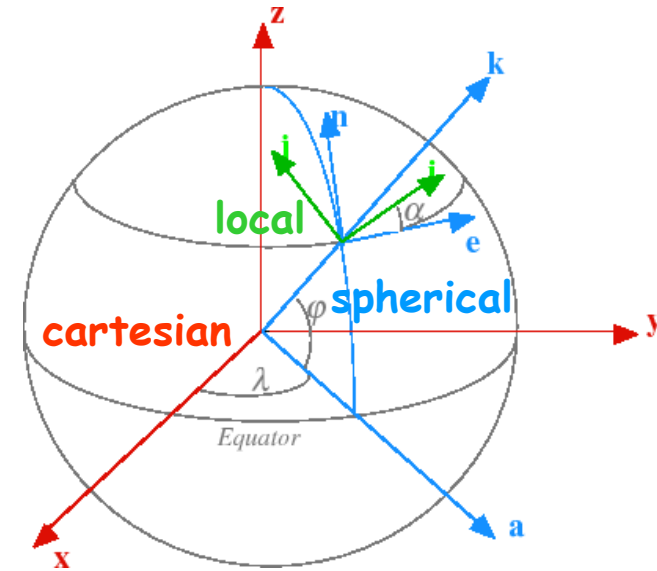
II.6 Vector interpolation (winds, currents, ...)

- ❖ interpolation of vectors component per component is not accurate, especially where the referential changes rapidly

Example interpolation of a zonal wind in the spherical referential near the pole



- At x , one would expect a zonal wind between 0 and 1.
- Interpolation comp. per comp. \rightarrow zonal wind of 1.



Solution (proposed by O. Marti, LSCE, implemented in OASIS):

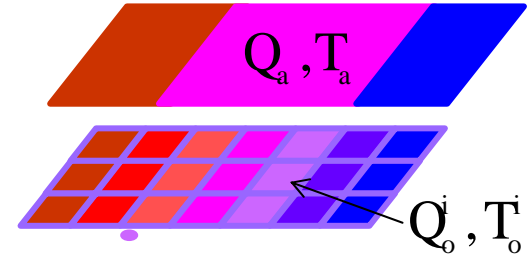
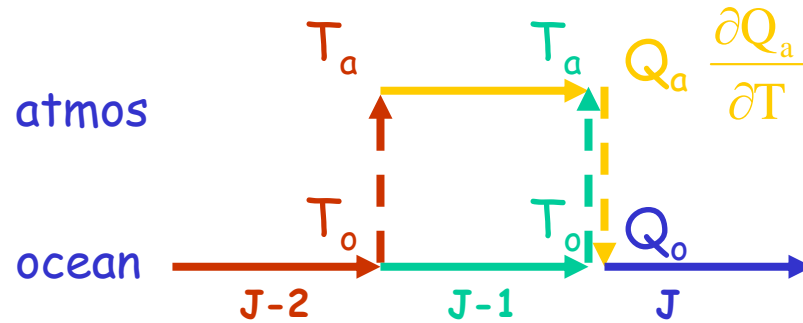
- "turn" the vector in the spherical referential and project the resulting vector in a cartesian referential
- interpolate the 3 components in the cartesian referential
- project back in the spherical referential; check that k component is zero
- possibly "turn" the resulting vector in the target local referential

Conclusions

- Different technical solutions to assemble model codes:
 - Coupling framework (e.g. ESMF):
 - best solution in a controlled development environment
 - Coupler (e.g. OASIS):
 - best solution to couple independently developed codes
- The OASIS coupler :
 - Coarse to fine grid remapping: subgrid variability with 2nd order remapping or SUBGRID (1st order Taylor expansion)
 - Non-matching sea-land masks:
 - DESTAREA: local flux conservation but unrealistic flux values
 - FRACAREA: no local flux conservation but realistic flux values
 - Global conservation can be artificially imposed
 - Vector interpolation: need to project components in a cartesian referential before interpolation.

The end

Use of SUBGRID transformation in practice:



Method 1:

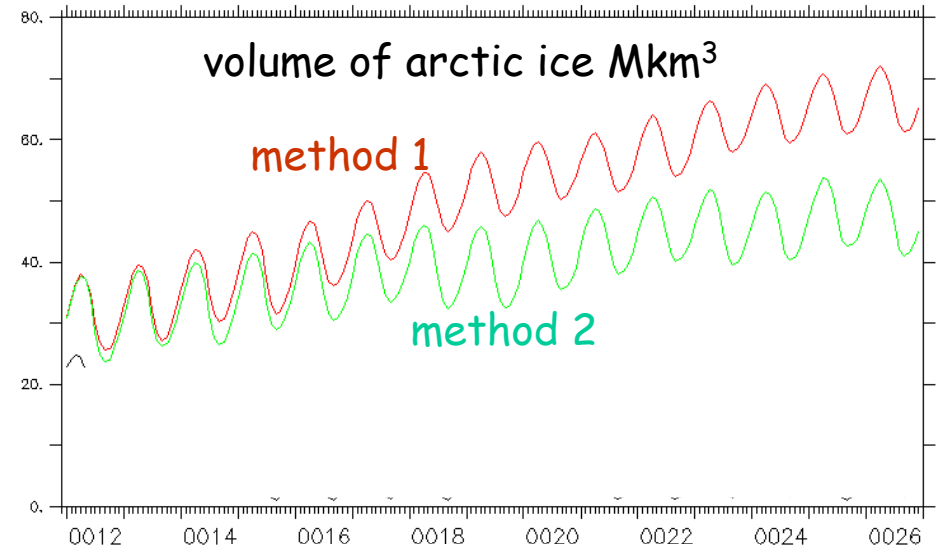
keep and use $T_o^i - T_a$ from J-2

$$Q_o^i = Q_a(T_a) + \delta Q_a / \delta T |_{T_a} \times (T_o^i - T_a)$$

Method 2:

send back T_a and use $T_o^i - T_a$ from J-1

$$Q_o^i = Q_a(T_a) + \delta Q_a / \delta T |_{T_a} \times (T_o^i - T_a)$$



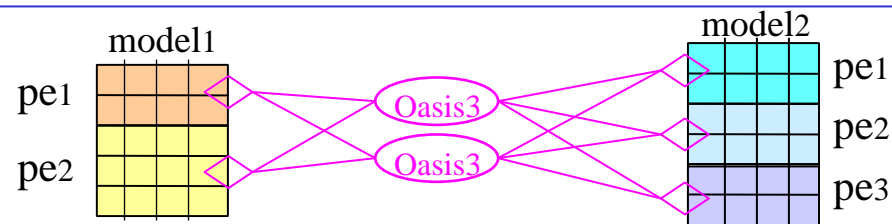
Code	Field name	Units	Definition & Conventions
1: atmosphere to Ocean-surface module exchange			
1.1	Rainfall + int. energy	kg/m ² /s + W/m ²	Mass flux and associated internal energy, positive downwards, includes all liquid precip.
1.2	Snowfall + int. energy	kg/m ² /s + W/m ²	Mass flux and associated internal energy, positive downwards, includes all solid precip.
1.3	Incoming solar radiation	W/m ²	Energy flux, positive downwards
1.4	Solar zenith angle	radians	
1.5	Diffuse solar radiation	W/m ²	Energy flux, positive downwards
1.6	Downward infrared radiation	W/m ²	Positive downward
1.7	Sensitivity of atmos. T and q to surface fluxes		$\partial T / \partial Q_s$ and $\partial q / \partial Q_s$
2: ocean-surface module to atmosphere exchange			
2.1	Sensible heat flux	W/m ²	Energy flux, positive upwards
2.2	Surface emissivity		
2.3	Albedo, direct	-	
2.4	Albedo, diffuse	-	
2.5	Surface radiative temperature	K	
2.6	Evaporation + int. energy	kg/m ² /s	Mass flux, positive upwards
2.7	Wind stress	N/m ²	Momentum flux, vector
2.8	Subgrid fractions	array of [0-1]	If multiple surfaces and tiling scheme
3: atmosphere to surface layer turbulence exchange			
3.1	Mean sea level surface pressure	hPa	
3.2	Air temperature at lowest level	K	
3.3	Air humidity at lowest level	g/g	
3.4	Wind at lowest level	m/s	Vector
3.5	Wind module at lowest level	m/s	Possibly including gustiness effects
3.6	Lowest level height	m	
4: surface layer turbulence to ocean-surface module exchange			
4.1	ρC_d drag coefficient	kg/m ² /s	Surface layer exchange coeff. for momentum
4.2	ρC_e exch. coeff.	kg/m ² /s	Surface layer exchange coeff. for sensible heat
4.3	ρC_h exch. coeff.	kg/m ² /s	Surface layer exchange coeff. for moisture
5: ocean-surface module to surface layer turbulence exchange			
5.1	Surface temperature	K	
5.2	Surface roughness		
5.3	Displacement height		
5.4	Surface velocity	m/s	

Code	Field name	Units	Definition & Conventions
6: ocean-surface module to ocean exchange			
6.1	Non solar heat flux	W/m ²	Energy flux, positive upwards
6.2	Solar radiation	W/m ²	Energy flux, positive downwards
6.3	Fresh water flux	kg/m ² /s	Mass flux, positive downwards
6.4	Salt flux	kg/m ² /s	Mass flux, positive downwards
6.5	Wind stress	N/m ²	Momentum flux, vector
6.6	Wind work	(m/s) ³	U^3
6.7	Mass of snow and ice	kg	
7: ocean to ocean-surface module exchange			
7.1	Temperature at sea-ice base	C	SST or more complex
7.2	Salinity at sea-ice base	PSU	
7.3	Highest level temperature	C	
7.4	Surface radiative temperature	C	
7.5	Surface current	m/s	Vector
7.6	Sea surface salinity	PSU	
7.7	Sea surface height	m	
7.8	Absorbed solar radiation in first oceanic layer	W/m ²	
8: land surface scheme to ocean exchange			
8.1	Continental runoff + int. energy	m ³ /s	Volume flux, positive towards ocean

II.2.2 The OASIS coupler: data exchange

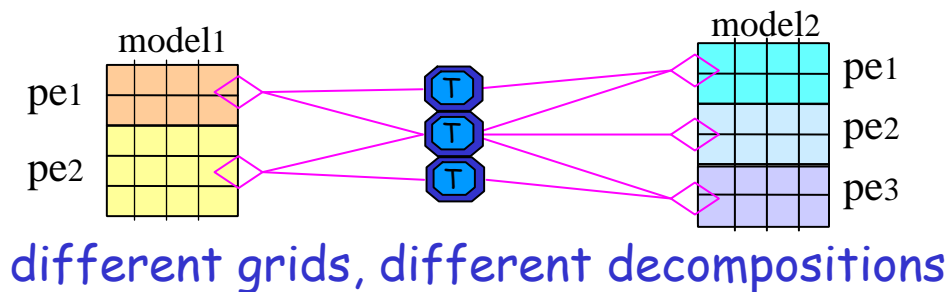
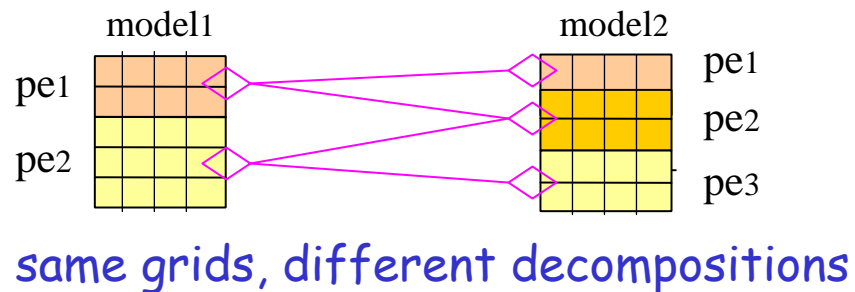
- communication library (MPI message passing) linked to the models

OASIS3: Parallel communication between parallel models and mono-process interpolation instance(s)



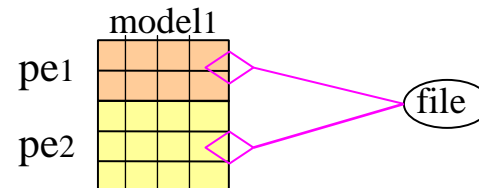
OASIS4: Parallel communication including repartitioning and

parallel interpolation

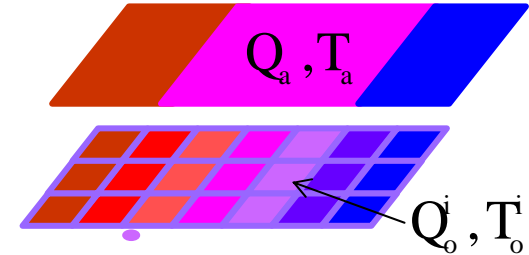
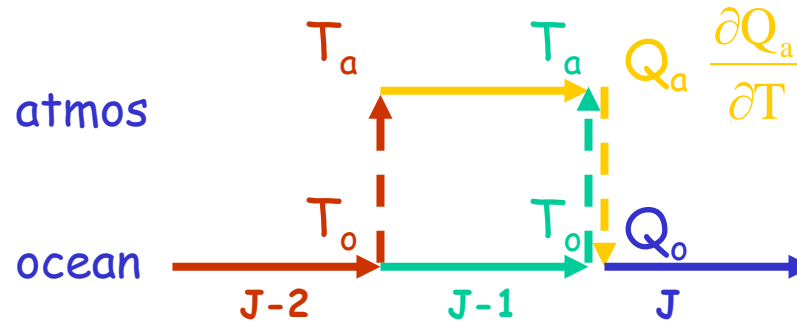


+ I/O functionality (GFDL mpp_io library)

- switch between coupled and forced mode



Use of SUBGRID transformation in practice:



Method 1:

send back T_a and use T_o^i from $J-1$, T_a from $J-2$

$$Q_o^i = Q_a(T_a) + \delta Q_a / \delta T |_{T_a} \times (T_o^i - T_a)$$

Method 2:

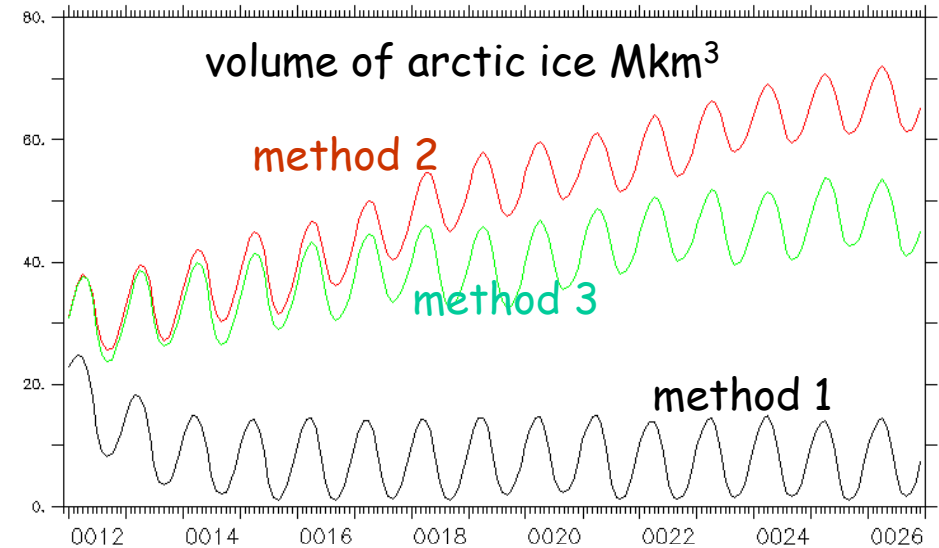
keep and use $T_o - T_a$ from $J-2$

$$Q_o^i = Q_a(T_a) + \delta Q_a / \delta T |_{T_a} \times (T_o - T_a)$$

Method 3:

send back T_a and use $T_o - T_a$ from $J-1$

$$Q_o^i = Q_a(T_a) + \delta Q_a / \delta T |_{T_a} \times (T_o - T_a)$$

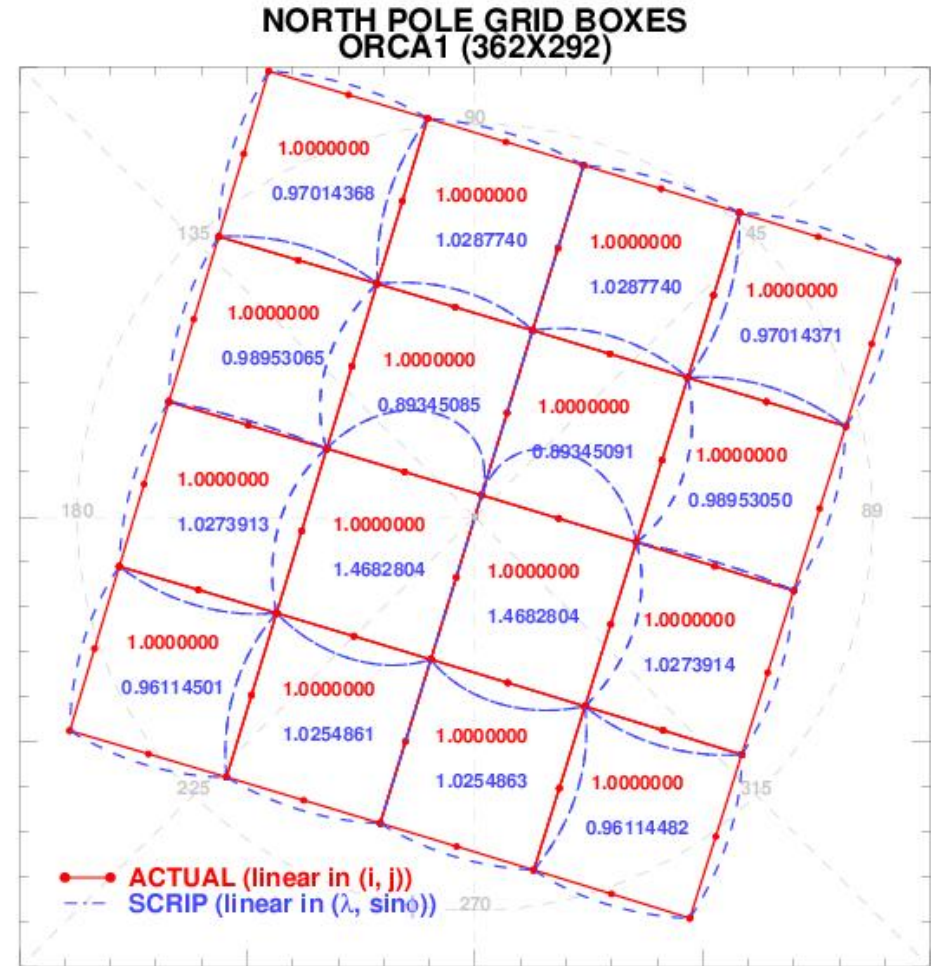


Remapping algorithms available in OASIS

(Los Alamos SCRIP library)

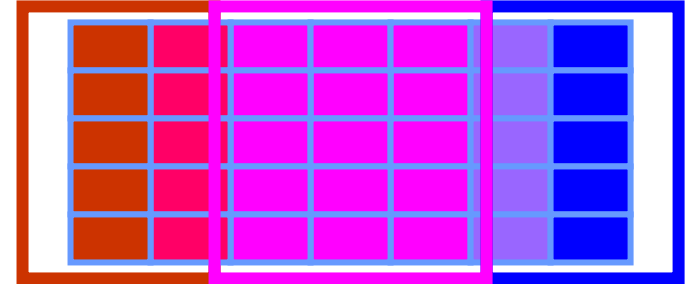
Actual limitations:

- borders are linear in (lat,lon) between corners (for intersection calculation)
 - uses Lambert equivalent azimuthal projection near the pole
- $\sin(\text{lat})$ linear function of \ln (for line intersection calculation); fractional error is $< .001$ further than 10 deg from the pole, and only ~ 0.1 with about 1 deg of it, for the ORCA1 example (for most gridcells the two measures of gridcell area agree to $< 5\%$, but for two gridcells they're off by 10%, and for another two they're off by 5%)
 - need to use a projection for line intersection calculation too
- exact calculation is not possible as "real shape of the borders are not known"
 - to ensure conservation, need to normalize by true area of the cells (i.e. as considered by the models themselves)



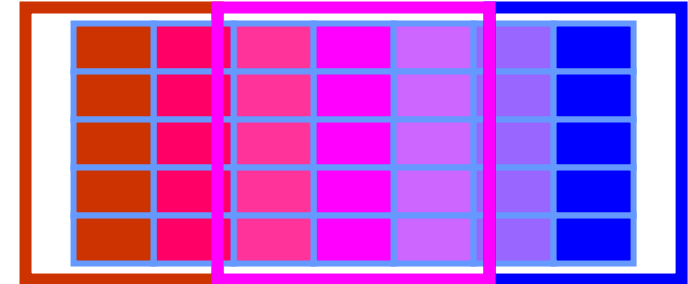
Part B - about ocean-atmosphere technical coupling software

- Problem with 1st order conservative remapping (low to high resolution)



- Solution 1: use 2nd order conservative remapping:

$$F_k = \sum_{n=1}^N \left[f_k w_{1nk} + \left(\frac{\partial f}{\partial lat} \right)_n w_{2nk} + \left(\frac{1}{\cos(lat)} \frac{\partial f}{\partial lon} \right)_n w_{3nk} \right]$$



- Solution 2: use SUBGRID transformation:

Solar type

Non-solar type



Part I - On a revised ocean-atmosphere physical coupling interface

Stand alone (forced AGCM)

