

Assimilation of Satellite Data for Atmospheric Composition

Hendrik Elbern,

**Rhenish Institute for Environmental Research at the
University of Cologne and IEK-8, FZ Jülich**

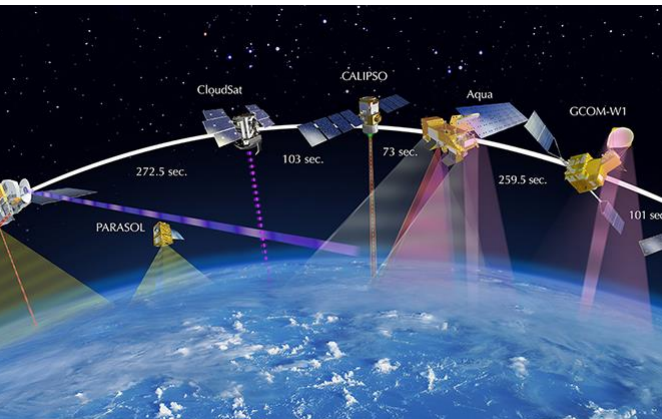
with substantial contributions from

**Angela Benedetti, ECMWF, Frederic Chevallier, CEA-CNRS-
UVSQ, IPSL, Richard Engelen, ECMWF, Johannes Flemming,
ECMWF, Antje Inness, ECMWF, Johannes Kaiser, MPI-C,
Sebastien Massart, ECMWF, and coworkers**

and many others

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5. Tropospheric aerosol assimilation
6. Greenhouse gas assimilation
7. Fire and unexpected event data assimilation
8. Look ahead



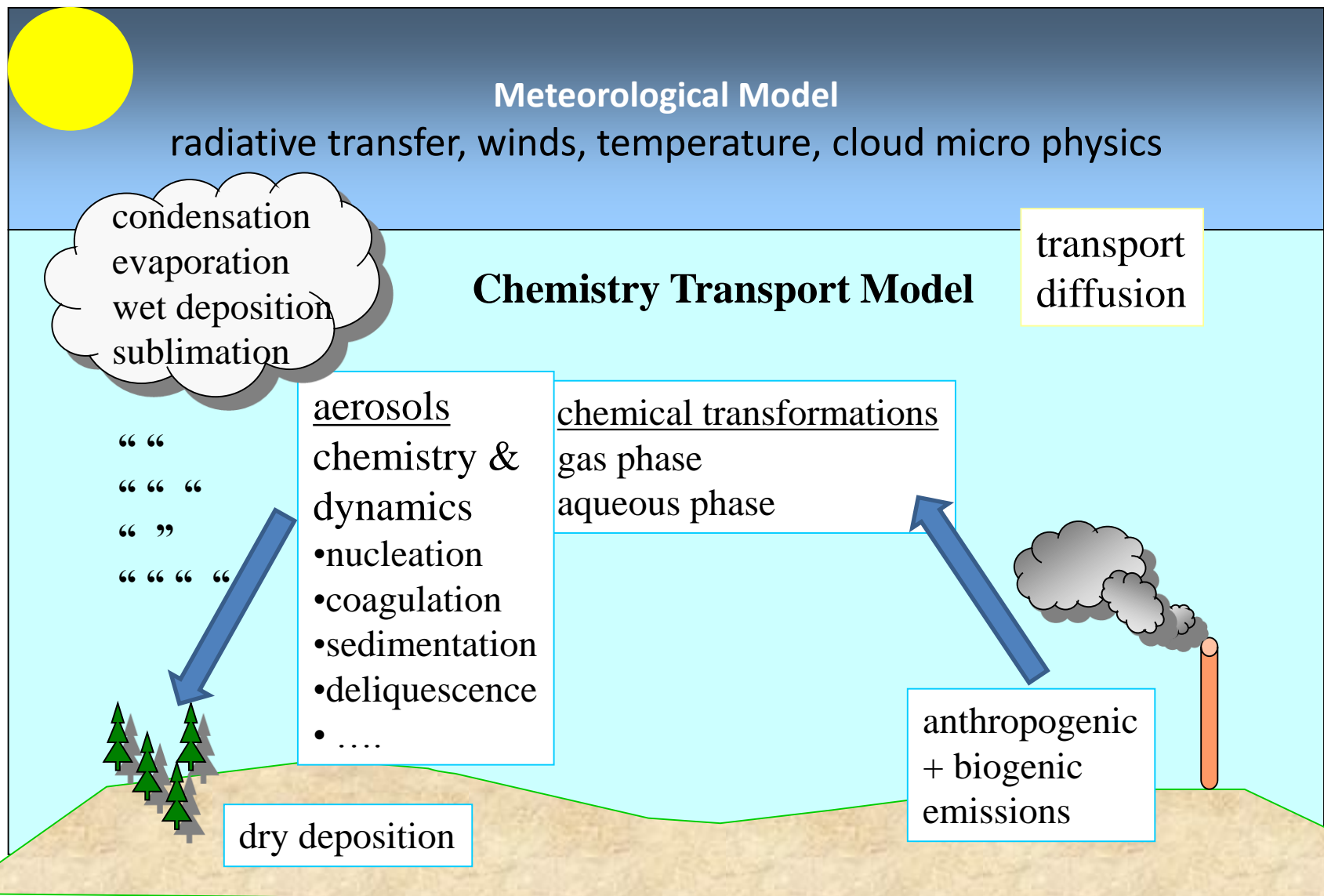
1. INTRODUCTION

What do we expect from composition data assimilation?

- ❑ daily „**chemical weather**“ forecasts = air quality prediction, the analog to NWP
 - ❑ exposure to polluted air and UV-B
- ❑ improve „classical“ NWP better calculation of the radiative transfer equation diabatic processes (aerosols, O₃,...)
- ❑ budget calculations of various constituents
- ❑ optimal chemical state analyses (= **monitoring**) or reanalyses:
 - ❑ earlier detection/attribution of climate change signals
 - ❑ (re-)assessment of radiative forcing

Processes in a complex chemistry-transport model

dim $\sim O(10^7)$



Characteristics of chemistry data assimilation (1)

physical viewpoint

Main sources of uncertainty:

- direct parameters
 - Initial values,
 - emission rates (in tropospheric data assimilation),
 - deposition and sedimentation velocities
 - reaction rates, J-values

- indirect parameters (in trop. data assimilation),
 - boundary layer height
 - vertical exchange mechanisms: convection

Characteristics of tropospheric chemistry data assimilation (2), mathematical viewpoints

- highly underdetermined system - on 2 levels
 - variables/gridpoint: ~ 60 -200
 - satellite data: scalar column value → profile vector
- regionally/locally highly nonlinear chemical dynamics (photo chemistry)
- constraints by physical laws/models are insufficient, however central manifolds variable (“initialisation” problem, chemical balance not guaranteed)
- assimilation or inversion problem to be solved?

Transport-diffusion-reaction equation and its adjoint

Tendency Equations

direct chemistry transport equation

$$\frac{\partial c_i}{\partial t} + \nabla \cdot (\mathbf{v}c_i) - \nabla \cdot (\rho \mathbf{K} \nabla \frac{c_i}{\rho}) - \sum_{r=1}^R \left(k(r) (s_i(r_+) - s_i(r_-)) \prod_{j=1}^U c_j^{s_j(r_-)} \right) = E_i + D_i$$

c_i concentration of species i

\mathbf{v} wind velocity

$k(r)$ reaction rate of reaction r

U number of species in the mechanism

E_i emission rate of species i (source)

c_i^* adjoint of concentration of species i

s stoichiometric coefficient

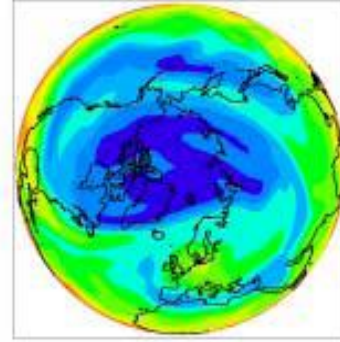
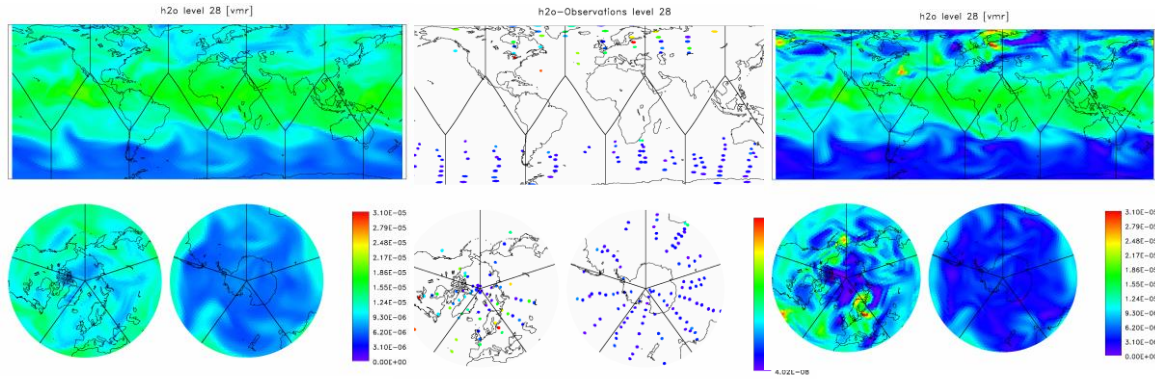
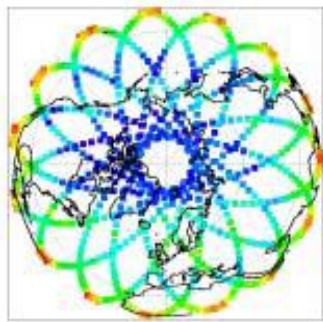
\mathbf{K} diffusion coefficient

R number of reactions in the mechanism

D_i deposition rate of species i (sink)

adjoint chemistry transport equation

$$-\frac{\partial \delta c_i^*}{\partial t} - \mathbf{v} \nabla \delta c_i^* - \frac{1}{\rho} \nabla \cdot (\rho \mathbf{K} \nabla \delta c_i^*) + \sum_{r=1}^R \left(k(r) \frac{s_i(r_-)}{c_i} \prod_{j=1}^U \bar{c}_j^{s_j(r_-)} \sum_{n=1}^U (s_n(r_+) - s_n(r_-)) \delta c_n^* \right) = 0$$



2. STRATOSPHERIC CHEMISTRY DATA ASSIMILATION

Table 1. Photolysis reactions included in the SA(

represents constituents that are not con

| Reaction |
|---|
| (R1) $O_2 + h\nu \rightarrow O(^3P) + O(^3P)$ |
| (R2) $O_3 + h\nu \rightarrow O(^3P) + O_2$ |
| (R3) $O_3 + h\nu \rightarrow O(^1D) + O_2$ |
| (R4) $H_2O + h\nu \rightarrow H + OH$ |
| (R5) $H_2O_2 + h\nu \rightarrow OH + OH$ |
| (R6) $NO_2 + h\nu \rightarrow O(^3P) + NO$ |
| (R7) $NO_3 + h\nu \rightarrow NO + O(^3P)$ |
| (R8) $NO_3 + h\nu \rightarrow NO_2 + O(^3P)$ |
| (R9) $N_2O + h\nu \rightarrow N_2 + O(^3P)$ |
| (R10) $N_2O_5 + h\nu \rightarrow NO_2 + NO_3$ |
| (R11) $HNO_3 + h\nu \rightarrow NO_2 + OH$ |
| (R12) $HNO_4 + h\nu \rightarrow NO_2 + OH$ |
| (R13) $Cl_2O_2 + h\nu \rightarrow Cl_2 + O_2$ |
| (R14) $Cl_2 + h\nu \rightarrow 2Cl$ |
| (R15) $OCIO + h\nu \rightarrow Cl + O_2$ |
| (R16) $HCl + h\nu \rightarrow H + Cl$ |
| (R17) $HOCl + h\nu \rightarrow OH + Cl$ |
| (R18) $ClONO + h\nu \rightarrow Cl + NO$ |
| (R19) $CH_3Cl + h\nu \rightarrow CH_3 + Cl$ |
| (R20) $CCl_4 + h\nu \rightarrow CCl_3 + Cl$ |
| (R21) $CFCl_3 + h\nu \rightarrow CFCl_2 + Cl$ |
| (R22) $CF_2Cl_2 + h\nu \rightarrow CF_2Cl + Cl$ |
| (R23) $CHF_2Cl + h\nu \rightarrow CHF_2 + Cl$ |
| (R24) $CF_2Cl_2 + h\nu \rightarrow CF_2 + Cl_2$ |
| (R25) $CH_3CCl_2 + h\nu \rightarrow CH_3CCl + Cl$ |
| (R26) $BrO + h\nu \rightarrow Br + O$ |
| (R27) $BrCl + h\nu \rightarrow Br + Cl$ |
| (R28) $HOBr + h\nu \rightarrow OH + Br$ |
| (R29) $BrONO + h\nu \rightarrow Br + NO$ |
| (R30) $CH_3Br + h\nu \rightarrow CH_3 + Br$ |
| (R31) $CF_2Cl_2 + h\nu \rightarrow CF_2 + Cl_2$ |
| (R32) $CF_3Br + h\nu \rightarrow CF_3 + Br$ |
| (R33) $HNO_4 + h\nu \rightarrow NO_2 + OH$ |
| (R34) $ClONO + h\nu \rightarrow Cl + NO$ |
| (R35) $N_2O_5 + h\nu \rightarrow NO_2 + NO_3$ |
| (R36) $CH_2O + h\nu \rightarrow H + HCO$ |
| (R37) $CH_2O + h\nu \rightarrow H_2 + CO$ |

stratospheric chemistry example

167 gas phase reactions +

10 heterogeneous reactions on polar strat. clouds

Table 2. Gas phase reactions that are included in

| Reaction |
|---|
| (R38) $O(^3P) + O_3 \rightarrow O_2 + O_2$ |
| (R39) $O(^1D) + O_2 \rightarrow O(^3P) + O_2$ |
| (R40) $O(^1D) + O_3 \rightarrow O_2 + O_2$ |
| (R41) $O(^1D) + O_3 \rightarrow O(^3P) + O_2$ |

Table 2. (continued)

Table 3. Heterogeneous reactions included in the SACADA reaction scheme. The notation "(c)"

indicates a species in the condensed (liquid or solid) phase. The term "products" represents constituents which are not considered in the reaction scheme.

| Reaction | Uptake coefficient | | |
|--|-----------------------------|--------|------|
| | liquid/STS | NAT | ice |
| (R168) $BrONO_2 + H_2O(c) \rightarrow HOBr + HNO_3$ | $f(t, p_{H_2O})^a$ | - | 0.26 |
| (R169) $N_2O_5 + H_2O(c) \rightarrow HNO_3 + HNO_3$ | $f(t, p_{H_2O})^a$ | 0.0004 | 0.02 |
| (R170) $ClONO_2 + H_2O(c) \rightarrow HNO_3 + HOCl$ | $f(t, p_{H_2O}, p_{HCl})^b$ | 0.004 | 0.3 |
| (R171) $ClONO_2 + HCl(c) \rightarrow Cl_2 + HNO_3$ | $f(t, p_{H_2O}, p_{HCl})^b$ | 0.2 | 0.3 |
| (R172) $HOCl + HCl(c) \rightarrow Cl_2 + H_2O$ | $f(t, p_{H_2O}, p_{HCl})^b$ | 0.1 | 0.2 |
| (R173) $N_2O_5 + HCl(c) \rightarrow HNO_3 + \text{products}$ | - | 0.003 | 0.03 |
| (R174) $HOBr + HCl(c) \rightarrow BrCl + H_2O$ | 0.01 | - | 0.3 |
| (R175) $ClONO_2 + HBr(c) \rightarrow BrCl + HNO_3$ | - | 0.3 | 0.3 |
| (R176) $HOCl + HBr(c) \rightarrow BrCl + H_2O$ | - | - | 0.05 |
| (R177) $BrONO_2 + HCl(c) \rightarrow BrCl + HNO_3$ | 0.3 | - | 0.3 |

a: as recommended by Sander et al. [2006]

b: Shi et al. [2001], as recommended by Sander et al. [2006]

| | | | |
|---|--|---|--|
| (R36) $CH_2O + h\nu \rightarrow H + HCO$ | (R71) $OH + HO_2 \rightarrow H_2O + O_2$ | (R105) $ClO + OH \rightarrow Cl + HO_2$ | (R141) $Br + HO_2 \rightarrow HBr + O_2$ |
| (R37) $CH_2O + h\nu \rightarrow H_2 + CO$ | (R72) $OH + H_2O_2 \rightarrow H_2O + O_2$ | (R106) $ClO + OH \rightarrow HCl + O_2$ | (R142) $BrO + HO_2 \rightarrow HOBr + O_2$ |
| | (R73) $HO_2 + O_3 \rightarrow OH + O_2$ | (R107) $OCIO + OH \rightarrow HOCl + O_2$ | (R143) $Br + O_3 \rightarrow BrO + O_2$ |
| | (R74) $HO_2 + HO_2 \rightarrow H_2O + O_2$ | (R108) $HCl + OH \rightarrow Cl + H_2O$ | (R144) $CH_2O + Br \rightarrow HBr + HCO$ |

The early challenge: Polar zone depletion motivated data assimilation with Chemistry-Transport Models

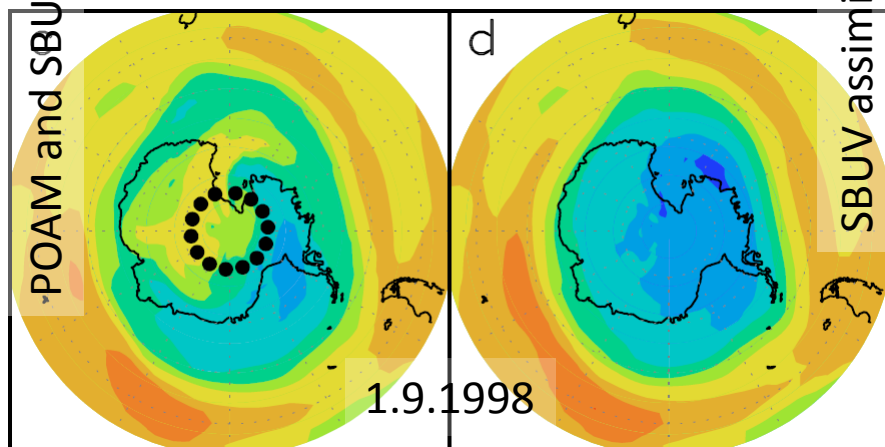
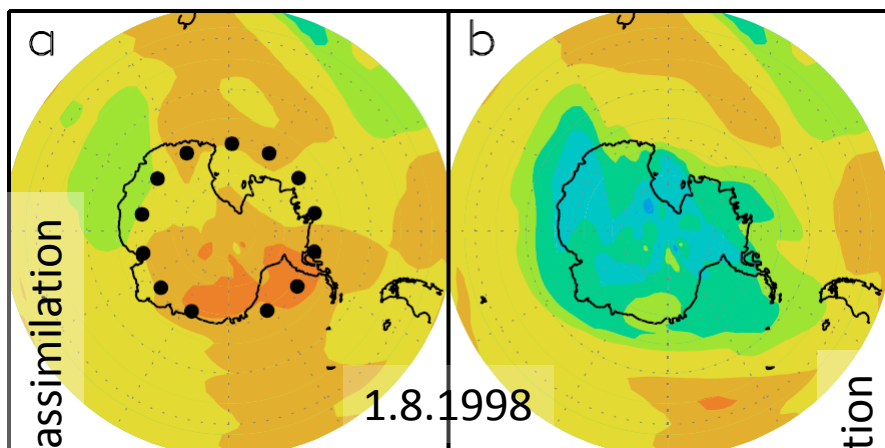
- CTMs driven by off-line winds and temperatures, e.g.:
 - Khattatov *et al.* 1999;
 - Errera and Fonteyn 2001;
 - Stajner *et al.* 2001;
 - Eskes *et al.* 2003,
 - Marchand *et al.* 2004
- assimilation of ozone (profiles and total columns) now operational at a number of institutions making use of CTMs:
 - KNMI <http://www.temis.nl/> **TM5**
 - BIRA-IASB <http://www.bascoe.oma.be/> **BASCOE**
 - DLR-DFD <http://taurus.caf.dlr.de> **SACADA**
 - NASA <http://gmao.gsfc.nasa.gov/operations/> **GEOS**

Early operational ozone analyses (Štajner *et al.*, 2001).

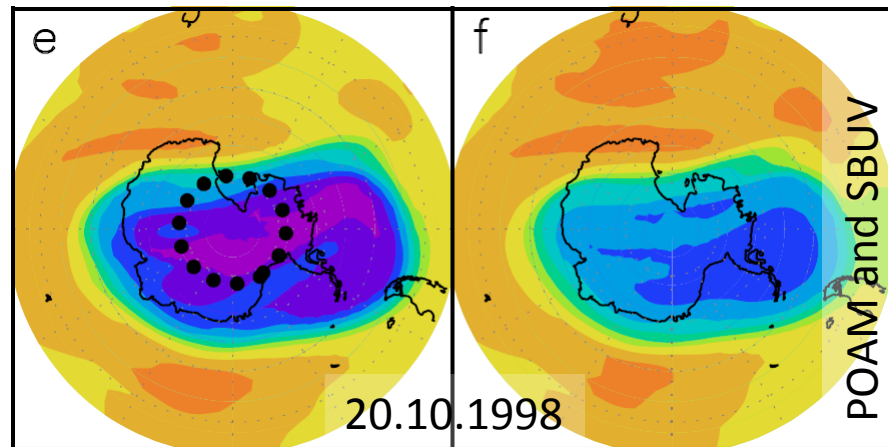
- Goddard Earth Observation System (GEOS) ozone data assimilation system 3-D CTM with parametrized ozone chemistry,
- χ^2 diagnostics to estimate the system parameters,
- operational in 1999,
- stratospheric ozone analyses
- SBUV/2 and TOMS

Example ozone hole focus: SBUV/2 and POAM-III (Štajner and Wargan 2004), combined **occultation**

Ozone mixing ratio (in ppmv) at 70 hPa



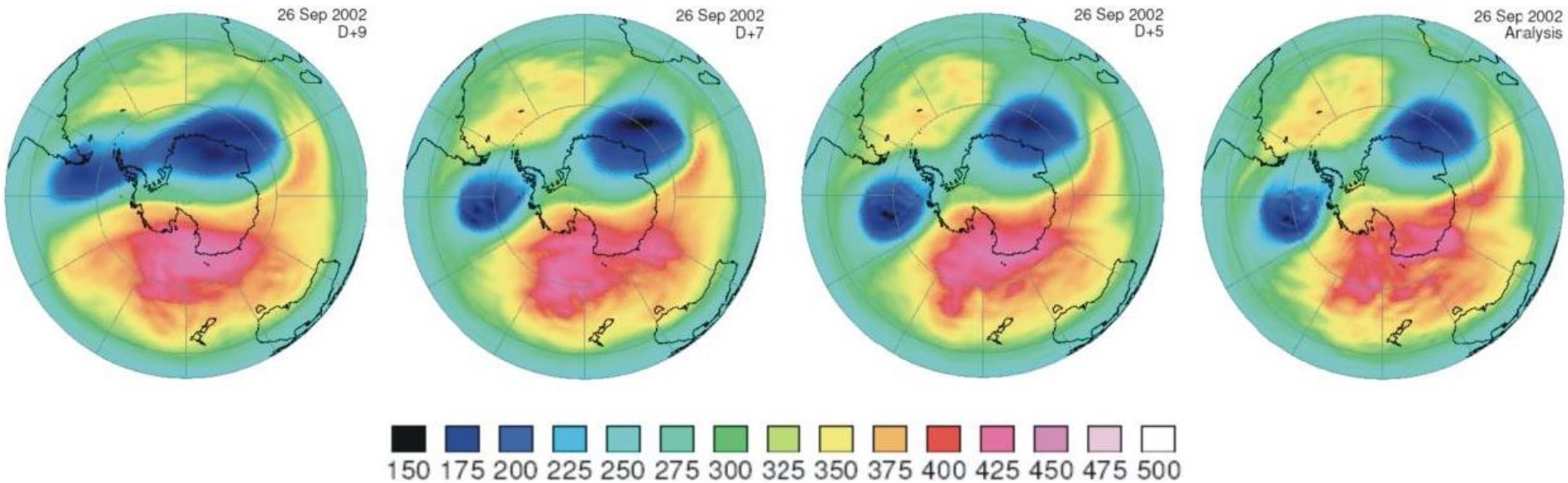
POAM III ozone data solar occultation austral winter and spring in 1998.
sun-synchronous satellite orbit,
14–15 sunsets and sunrises per day,
25.4 longitude apart,
two latitude circles 54N to 71N, 63S to 88S
vertical resolution of v3 ozone 1.1 km
random errors < 5% for $z > 15$ km



● POAM III occultation positions

O3 total column data assimilation (GOME)

Forecasts and analysis of the first southern vortex split event

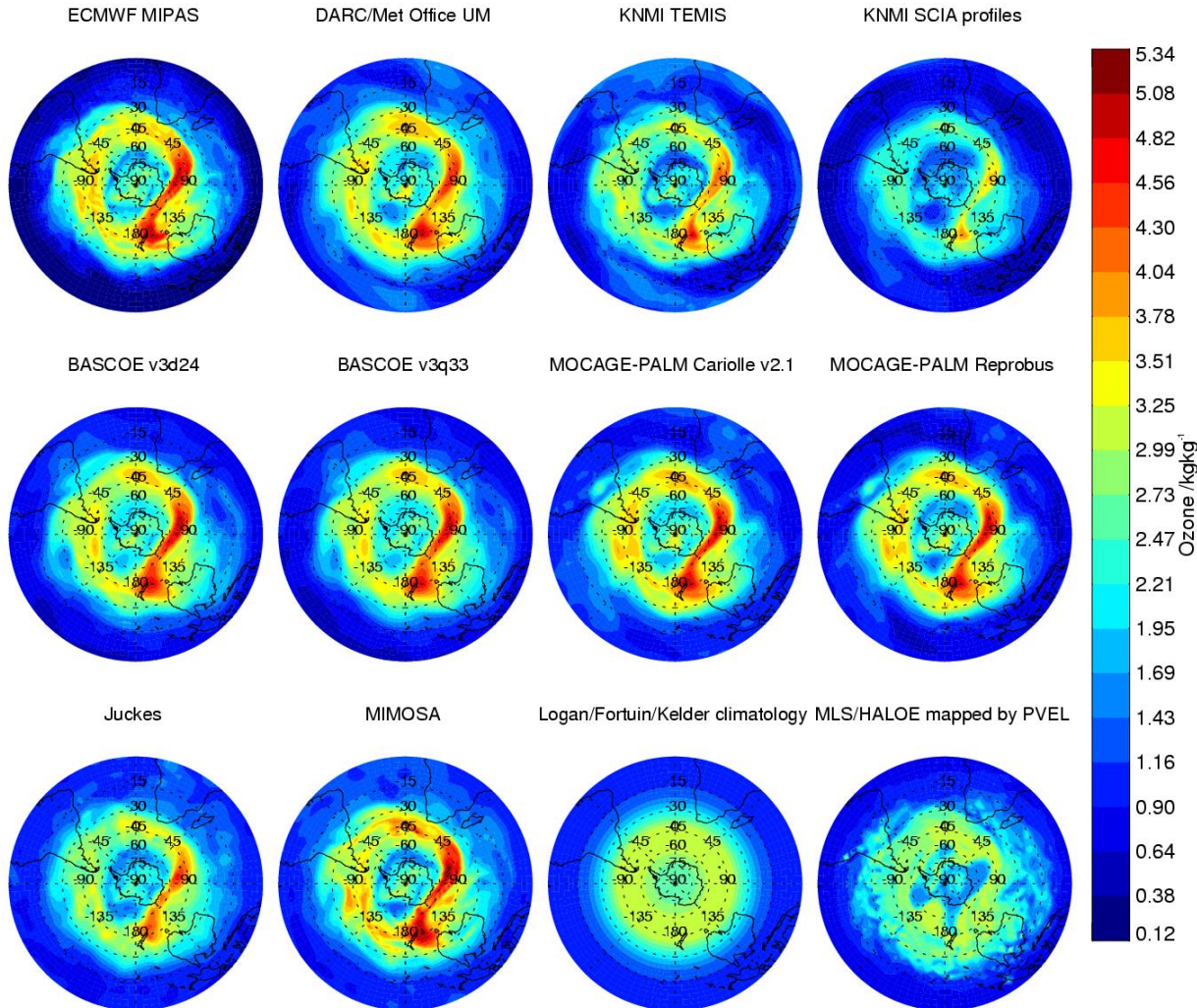


Ozone total column on 26 September 2002, KNMI operational ozone assimilation system. From left to right: 9-, 7-, 5-day forecasts, and the corresponding analysis.

From Eskes *et al.* (2005), KNMI.

ASSET data assimilation analysis comparison:

Ozone at 68.13hPa 12:00:00 31-Aug-2003

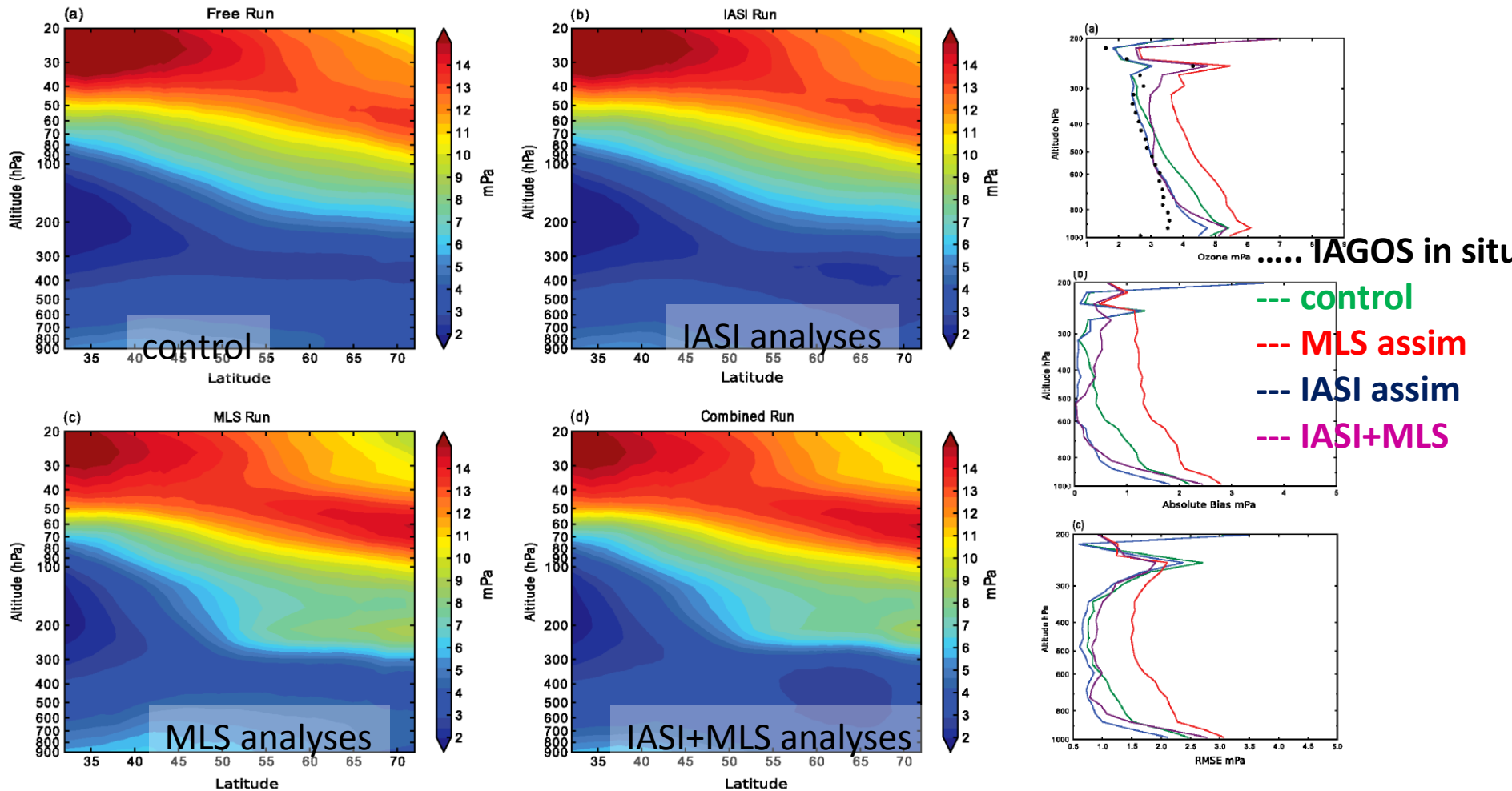


Ozone (ppm) at
68 hPa in the
southern
hemisphere on
31st August 2003.

from Lahoz et al, 2010

Combined assimilation of IASI ozone tropospheric columns and stratospheric MLS profiles

by Barre, J Peuch, VH Lahoz, WA Attie, JL Josse, B Piacentini, Eremenko, M Dufour, Nedelec, P; von Clarmann, T El Amraoui, (2014)



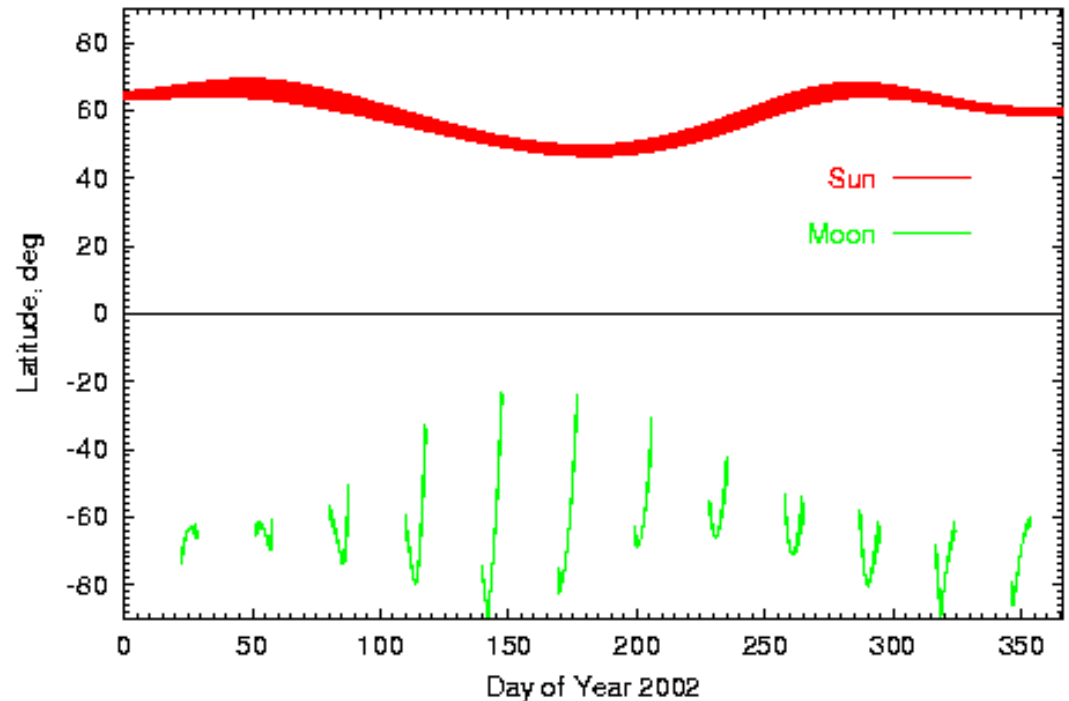
zonal mean of O3 field July 2009 over Europe 15W-35E

ENVISAT MIPAS and SCIAMACHY by SACADA 4D-var-assimilation system

Data Availability 21. Oct. – 14. Nov. 2003

| | MIPAS-IMK | | SCIA-Limb | | SCIA-Occ | |
|--------|-----------|-------------|--|-------------|-----------|-------------|
| | Available | Assimilated | Available | Assimilated | Available | Assimilated |
| O3 | X | X | X | | X | |
| NO2 | X | X | X | | X | |
| N2O | X | X | | | | |
| HNO3 | X | X | | | | |
| HNO4 | X | | | | | |
| NO | X | | | | | |
| N2O5 | X | X | | | | |
| H2O | X | X | | | | |
| CH4 | X | X | | | | |
| CFC-11 | X | X | | | | |
| CFC-12 | X | X | | | | |
| ClO | X | | MIPAS : von Clarman, Stiller et al., KIT Karlsruhe | | | |
| ClONO2 | X | X | SCIAMACHY: Bovensmann and coworkers, Uni Bremen | | | |
| BrO | | | X | | | |

SCIAMACHY Solar Occultation Data Analysis for SACADA



O_3 and NO_2

J. Meyer, A. Bracher, L. Amekudzi, S. Noel, A. Rozanov, B. Hoffmann, H.
Bovensmann, J. P. Burrows

Institute for Environmental Physics, University of Bremen, Germany

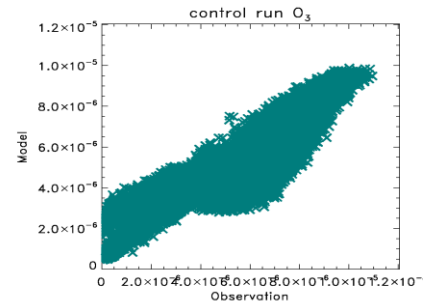
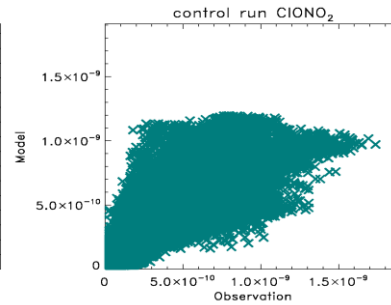
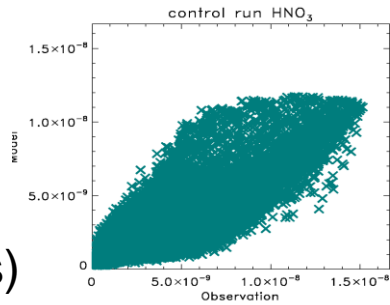
Scatter plots for Nov. 13, 2003

HNO₃

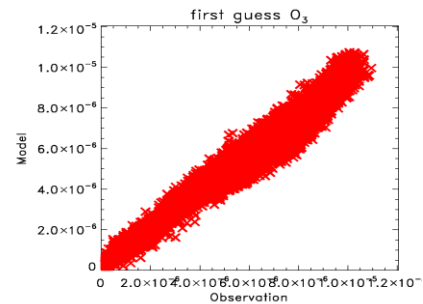
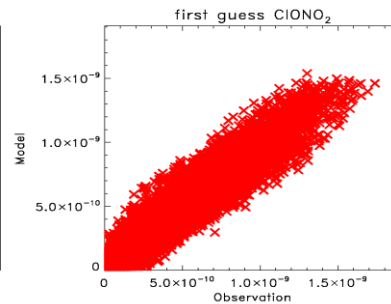
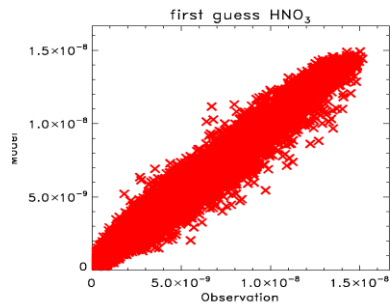
ClONO₂

O₃

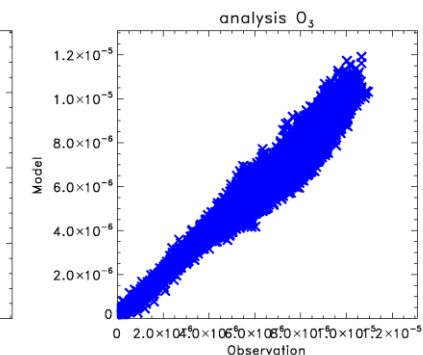
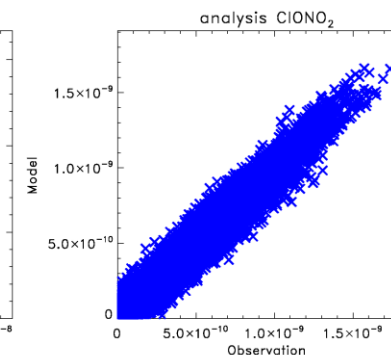
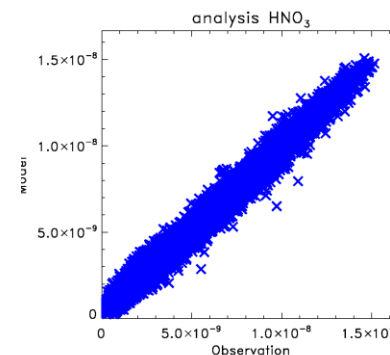
Control Run
(after 23 days of
free integration
with SOKRATES
(2D) initial values)



23 days
consecutive
4D-var:



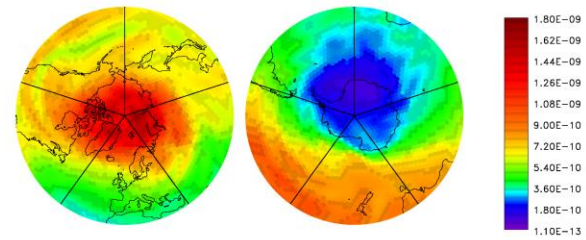
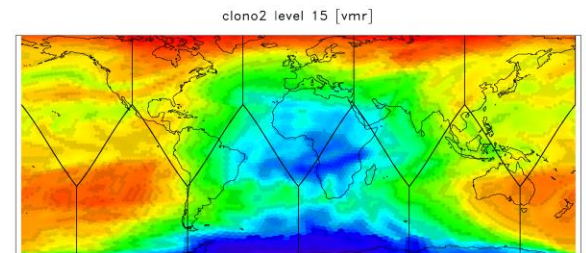
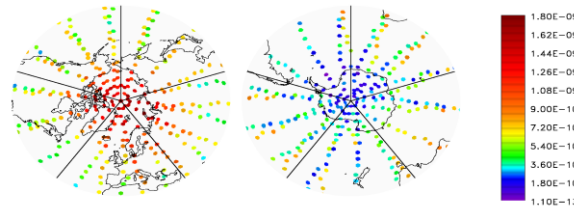
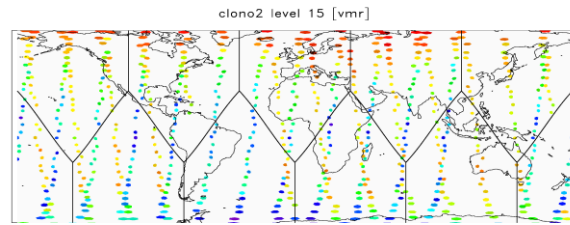
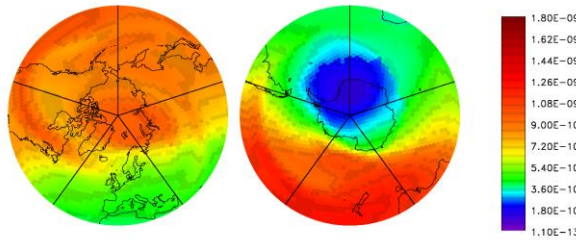
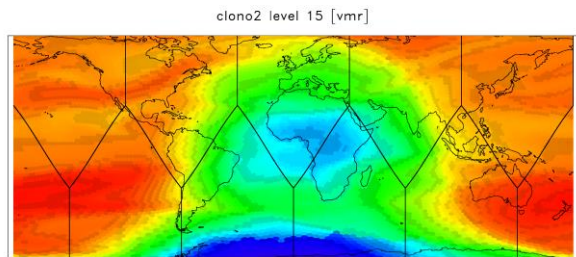
Background



Analysis

SACADA 4D-var, 24 h assimilation window

Results for ClONO₂ at 7.6 hPa (~33 km), Nov. 13, 2003 12:00 UTC



Control Run
(no assimilation)

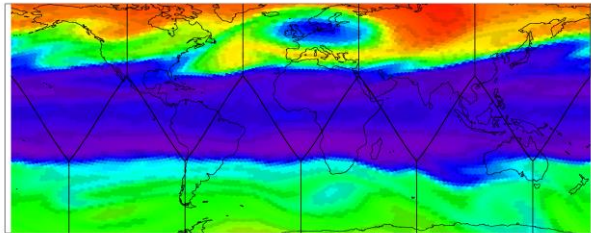
MIPAS Observations

Analysis

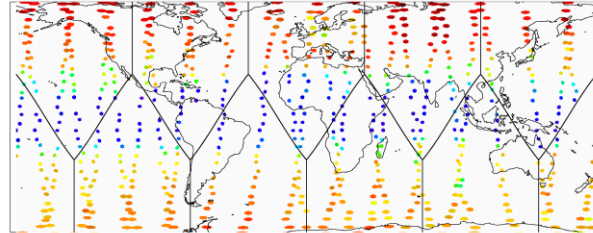
SACADA 4D-var, 24 h assimilation window

HNO_3 at 28 hPa (~ 24 km), Nov. 4, 2003 12:00 UTC

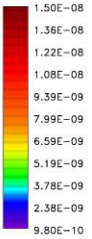
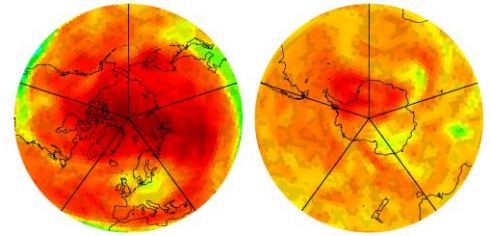
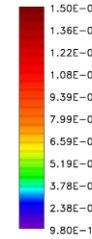
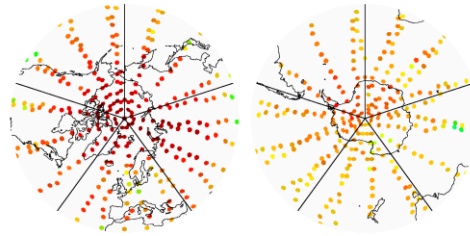
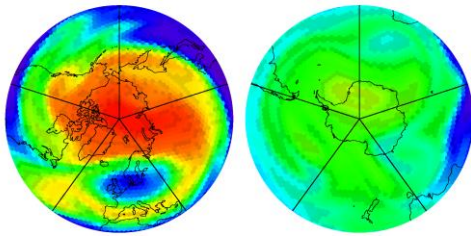
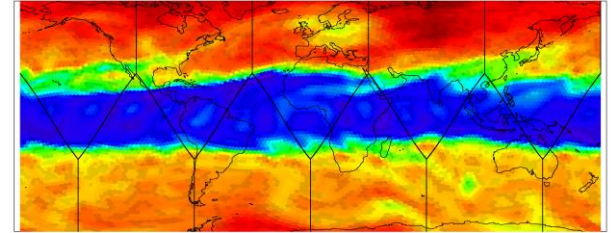
hno3 level 20 [vmr]



hno3 level 20 [vmr]



hno3 level 20 [vmr]

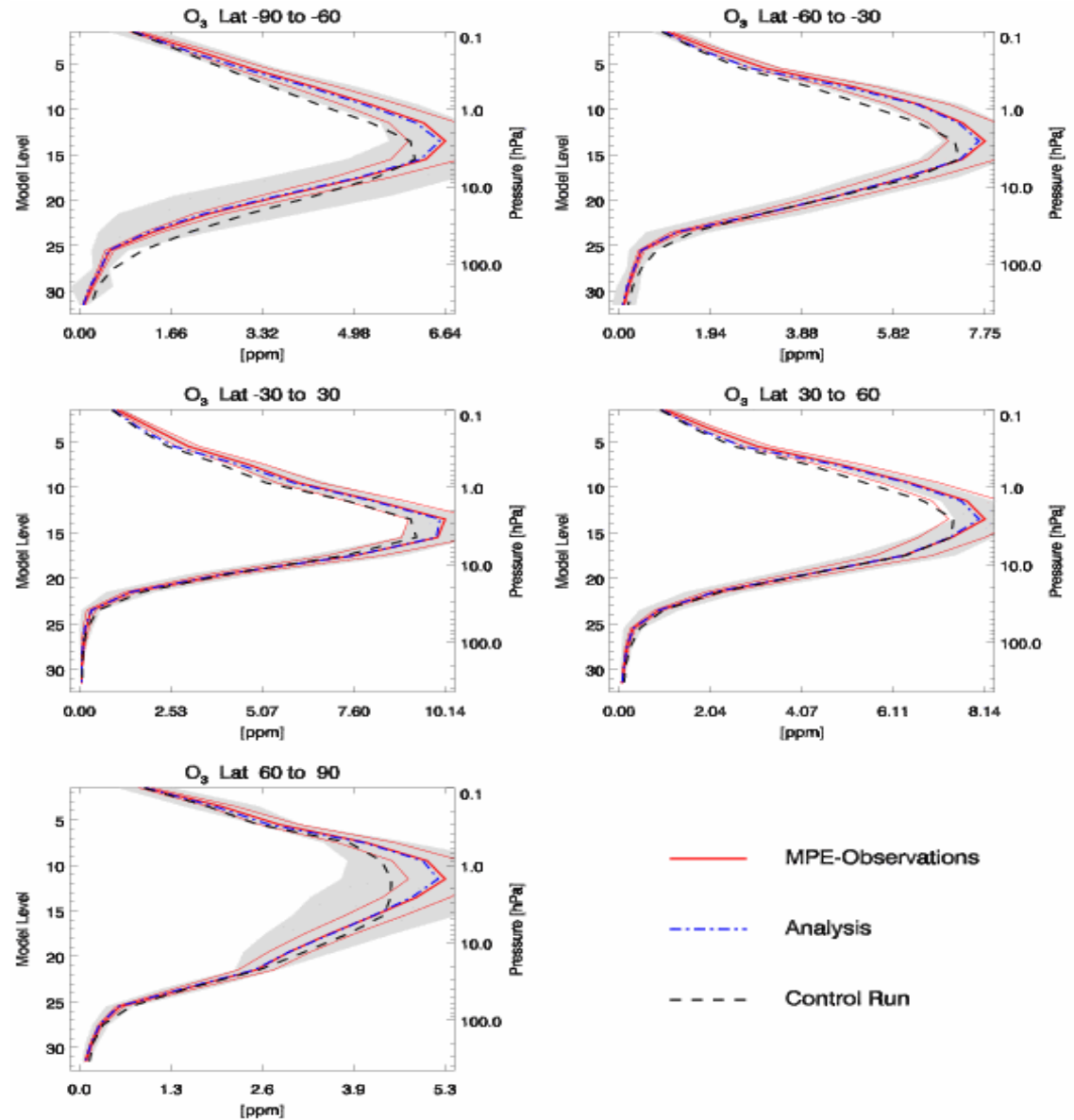


Control Run
(no assimilation)

MIPAS Observations

Analysis

Ozone profiles averaged over the latitude belts indicated and the time span 8.9.-15.10.2002



Full global atmosphere: Data used in MACC NRT system

| Instrument | Satellite | Satellite operator | Data provider | Species | Status |
|-----------------------|-------------|--------------------|---------------|-----------------|----------------------|
| MODIS | Terra | NASA | NASA/NOAA | Aerosol, fires | Active |
| MODIS | Aqua | NASA | NASA/NOAA | Aerosol, fires | Active |
| SEVIRI | Meteosat-9 | EUMETSAT | IM | Fires | Active |
| Imager | GOES-11, 12 | NOAA | NOAA | Fires | Passive |
| Imager | MTSAT-2 | JMA | JMA | Fires | Planned |
| MLS | Aura | NASA | NASA | O ₃ | Active |
| OMI | Aura | NASA | NASA | O ₃ | Active |
| SBUV-2 | NOAA-16,19 | NOAA | NOAA | O ₃ | Active |
| SCIAMACHY | Envisat | ESA | KNMI | O ₃ | Died |
| GOME-2 | Metop-A | EUMETSAT | DLR | O ₃ | Active |
| GOME-2 | Metop-B | EUMETSAT | DLR | O ₃ | Passive |
| IASI | Metop-A | EUMETSAT | LATMOS/ULB | CO | Active |
| IASI | Metop-B | EUMETSAT | LATMOS/ULB | CO | Passive |
| MOPITT | Terra | NASA | NCAR | CO | Active |
| GOME-2 | Metop-A | EUMETSAT | DLR | NO ₂ | Passive/Tests |
| GOME-2 | Metop-B | EUMETSAT | DLR | NO ₂ | Passive/Tests |
| OMI | Aura | NASA | KNMI | NO ₂ | Active |
| OMI | Aura | NASA | NASA | SO ₂ | Active |
| GOME-2 | Metop-A | EUMETSAT | DLR | SO ₂ | Passive/Tests |
| GOME-2 | Metop-A | EUMETSAT | DLR | SO ₂ | Passive/Tests |
| GOME-2 | Metop-B | EUMETSAT | DLR | HCHO | Passive |
| Offline tests: | | | | | |
| IASI | Metop-A | EUMETSAT | LATMOS/ULB | O ₃ | Tests |

Courtesy: A. Benedetti, R. Engelen, J. Flemming, A. Inness, S. Massart, ECMWF, MACC

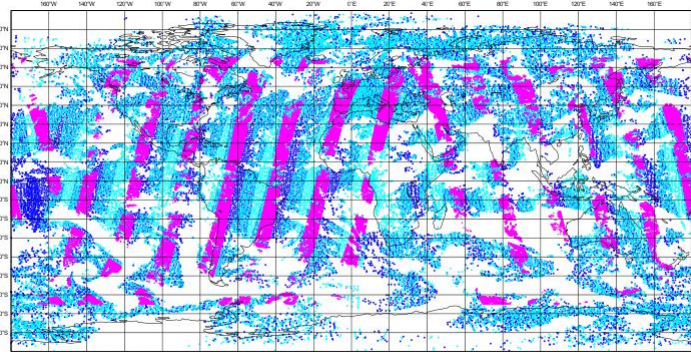
Setup for the reactive gases assimilation

- IFS species: O₃, CO, NO₂, SO₂, HCHO
- More species available from CTM output (and in C-IFS)
- Coupled system or C-IFS
- Background errors calculated with:
 - NMC method (CO, NO_x, HCHO)
 - Analysis ensemble method (O₃)
 - Prescribed profile (SO₂)
- Difficulties assimilating species with short lifetimes (e.g. NO₂): NO_x as control variable and NO₂-NO_x interconversion operator
- Variational bias correction used for reactive gases
- Chemistry included in outer loop (ifstraj) not in minimisation; adjoint of transport only

Reactive gases data usage in MACC NRT system: 20130801, 12z

Tropospheric NO₂

CO

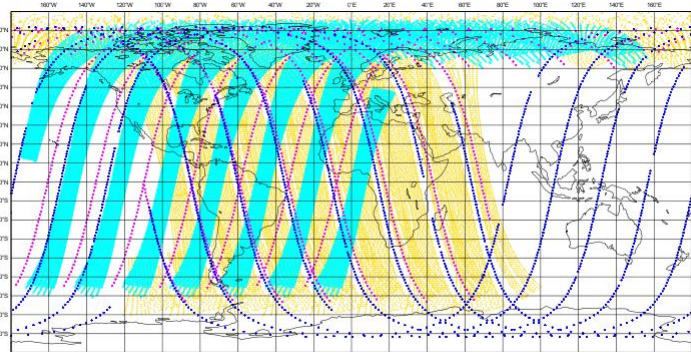


IASI
Metop-A

IASI
Metop-B

MOPITT
TERRA

O₃



GOME-2
Metop-A

OMI
AURA

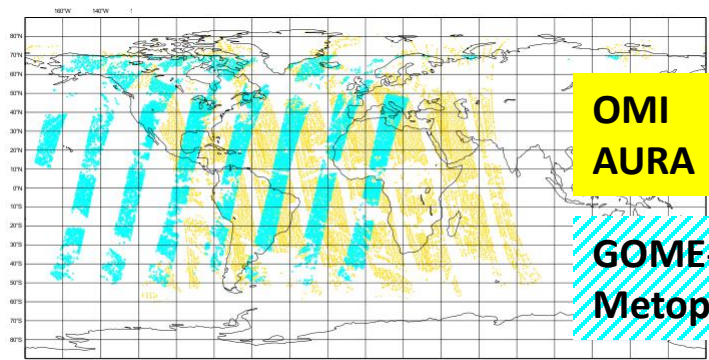
MLS
AURA

SBUV/2
NOAA-16
NOAA-19

assimilated

monitored

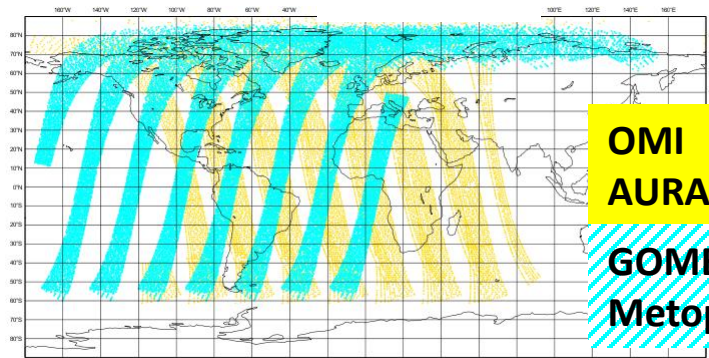
Tropospheric NO₂



OMI
AURA

GOME-2
Metop-A

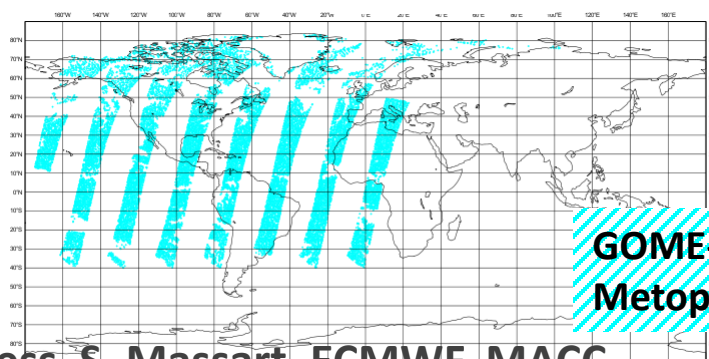
SO₂



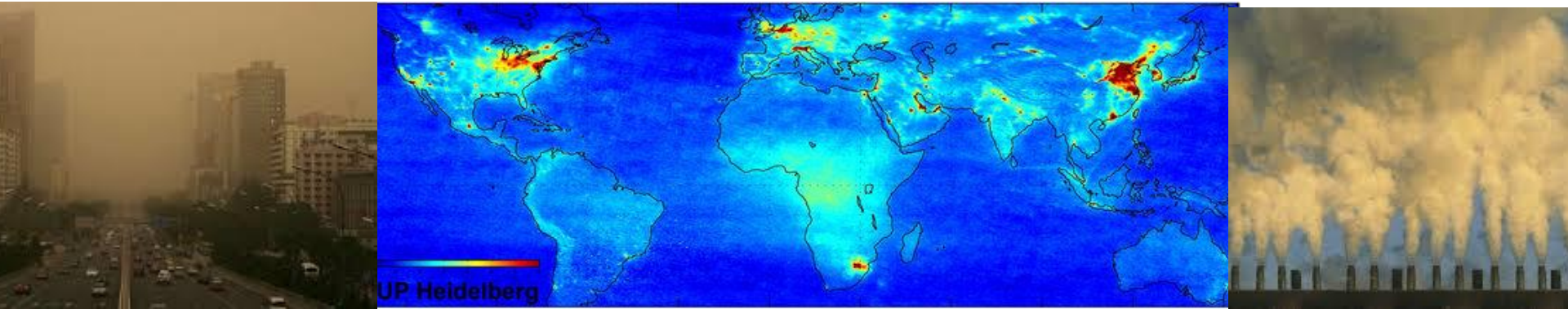
OMI
AURA

GOME-2
Metop-A

HCHO

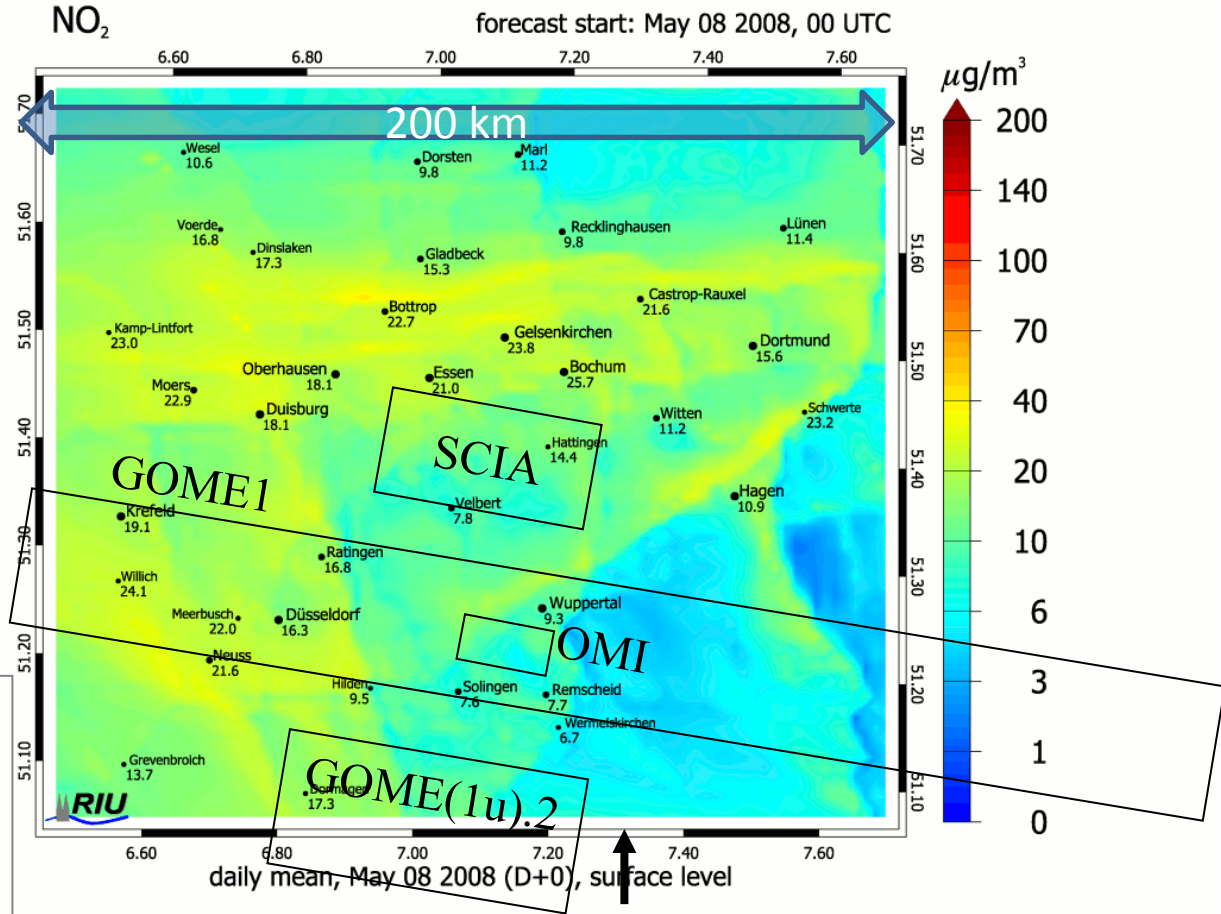
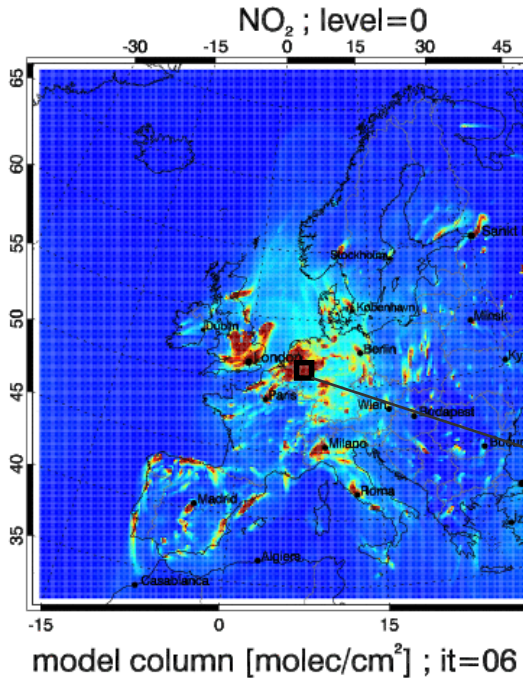


GOME-2
Metop-A



3. SATELLITE DATA ASSIMILATION FOR TROPOSPHERE AND AIR QUALITY

NO₂ tropospheric column satellite information: ESA UV-VIS satellite footprints Ruhr area comparison



nadir areas:

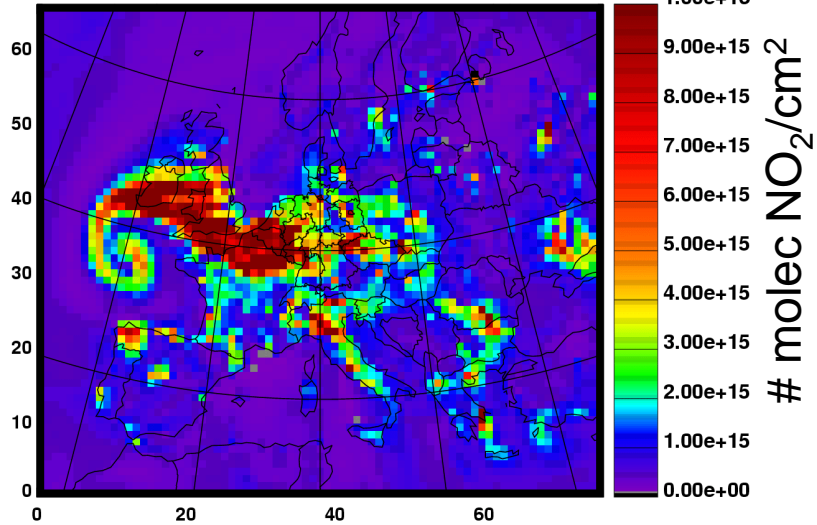
| | |
|----------------|--------------------------|
| GOME 1 | 320 x 40 km ² |
| (special mode) | 80 x 40 “ |
| SCIAMACHY | 60 x 30 “ |
| GOME 2 | 80 x 40 “ |
| OMI | 24 x 13 “ |

Ruhr area domain
(~12 000 000 inhabitants)

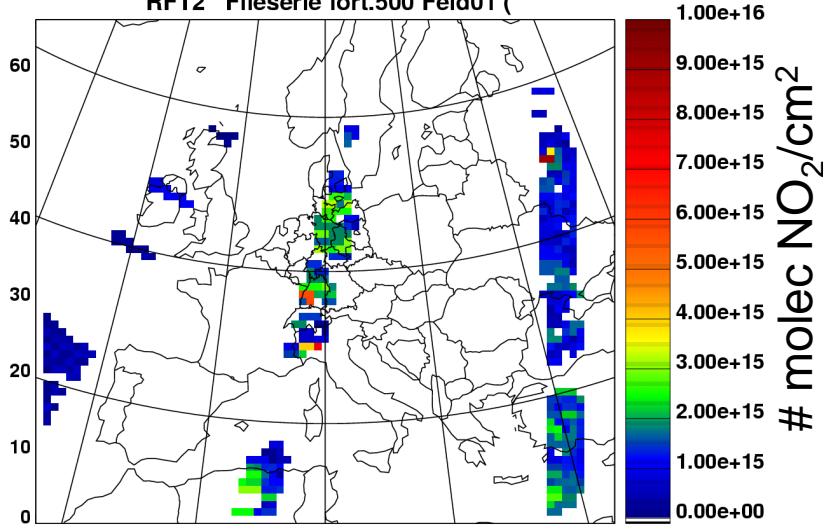
Assimilation of GOME NO₂ tropospheric columns, 4Dvar with EURAD-IM

forecast without assimilation

T2: RNO2 tropospheric column 5.8.1997 forecast for 10:30 UTC (t=0)

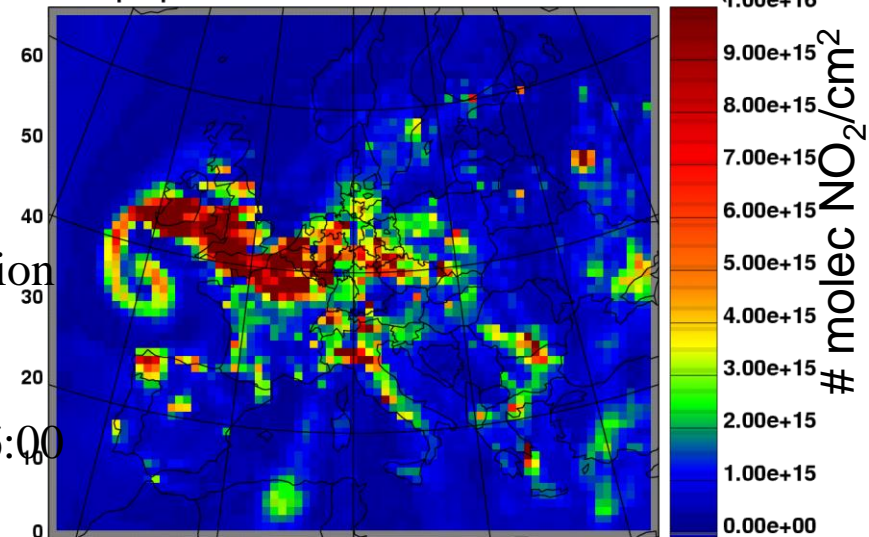


RFT2 Fileserie fort.500 Feld01 (



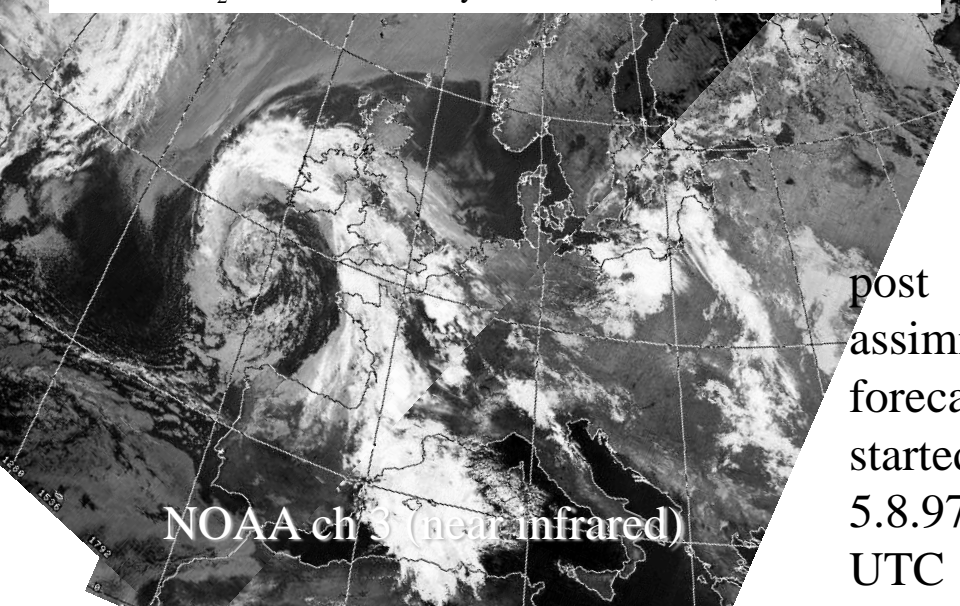
GOME NO₂ columns: Courtesy of A. Richter, IFE, U. Bremen

T2: RNO2 tropospheric column 5.8.1997 forecast for 10:30 UTC (t=8)



post
assimilation
forecast
started:
5.8.97 06:00
UTC

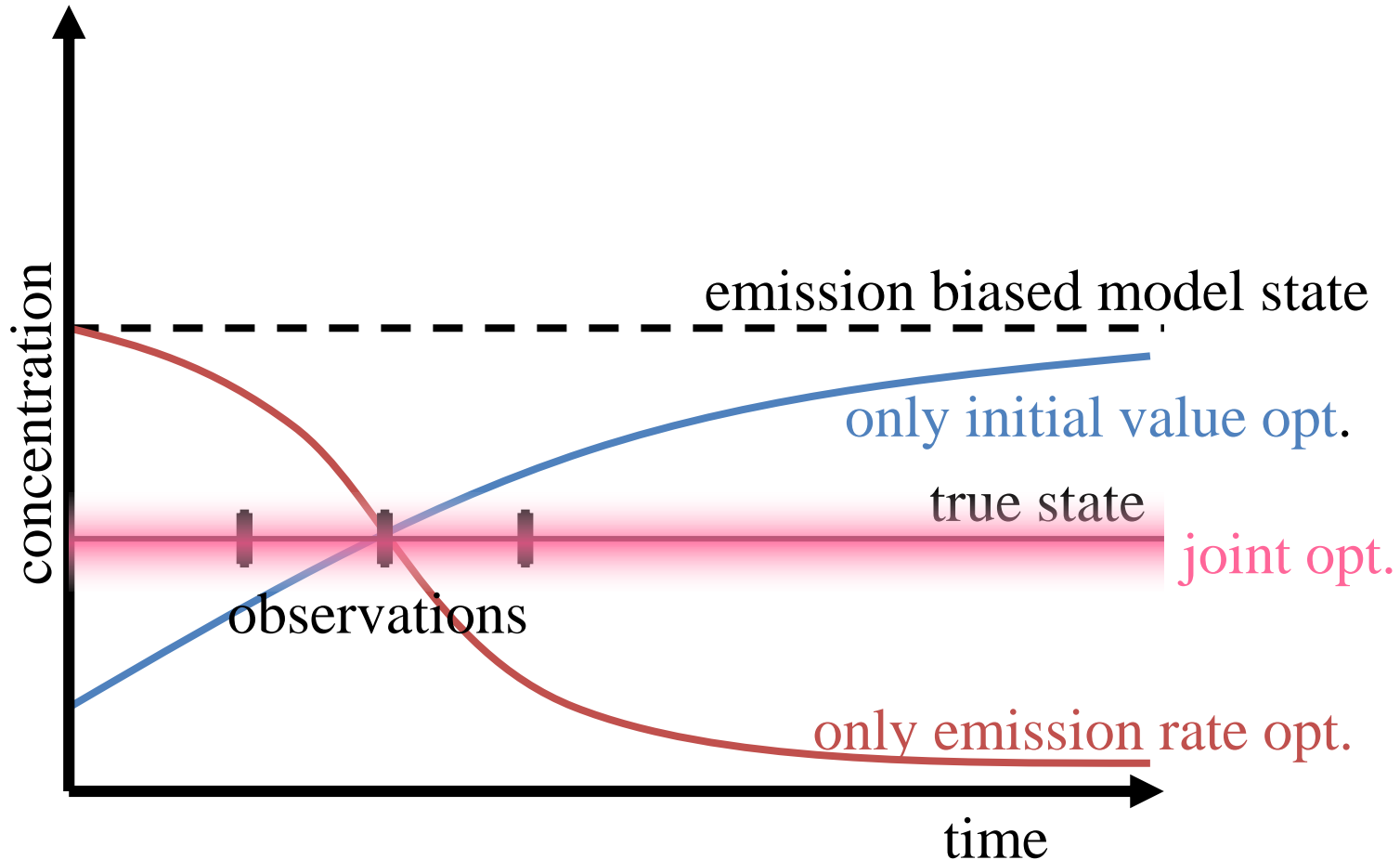
NOAA ch 3 (near infrared)



Question: Which parameter to be optimized?

Hypothesis:

initial state and emission rates are least known



In the troposphere, for **emission rates**, the product (*paucity of knowledge * importance*) is high

Emission Rate Optimization

minimize cost function

$$J(\mathbf{x}(t_0), \mathbf{e}) = \frac{1}{2}(\mathbf{x}^b(t_0) - \mathbf{x}(t_0))^T \mathbf{B}_0^{-1}(\mathbf{x}^b(t_0) - \mathbf{x}(t_0)) + \frac{1}{2} \int_{t_0}^{t_N} (\mathbf{e}_b(t) - \mathbf{e}(t))^T \mathbf{K}^{-1}(\mathbf{e}_b(t) - \mathbf{e}(t)) dt + \frac{1}{2} \int_{t_0}^{t_N} (\mathbf{y}^0(t) - H[\mathbf{x}(t)])^T \mathbf{R}^{-1}(\mathbf{y}^0(t) - H[\mathbf{x}(t)]) dt$$

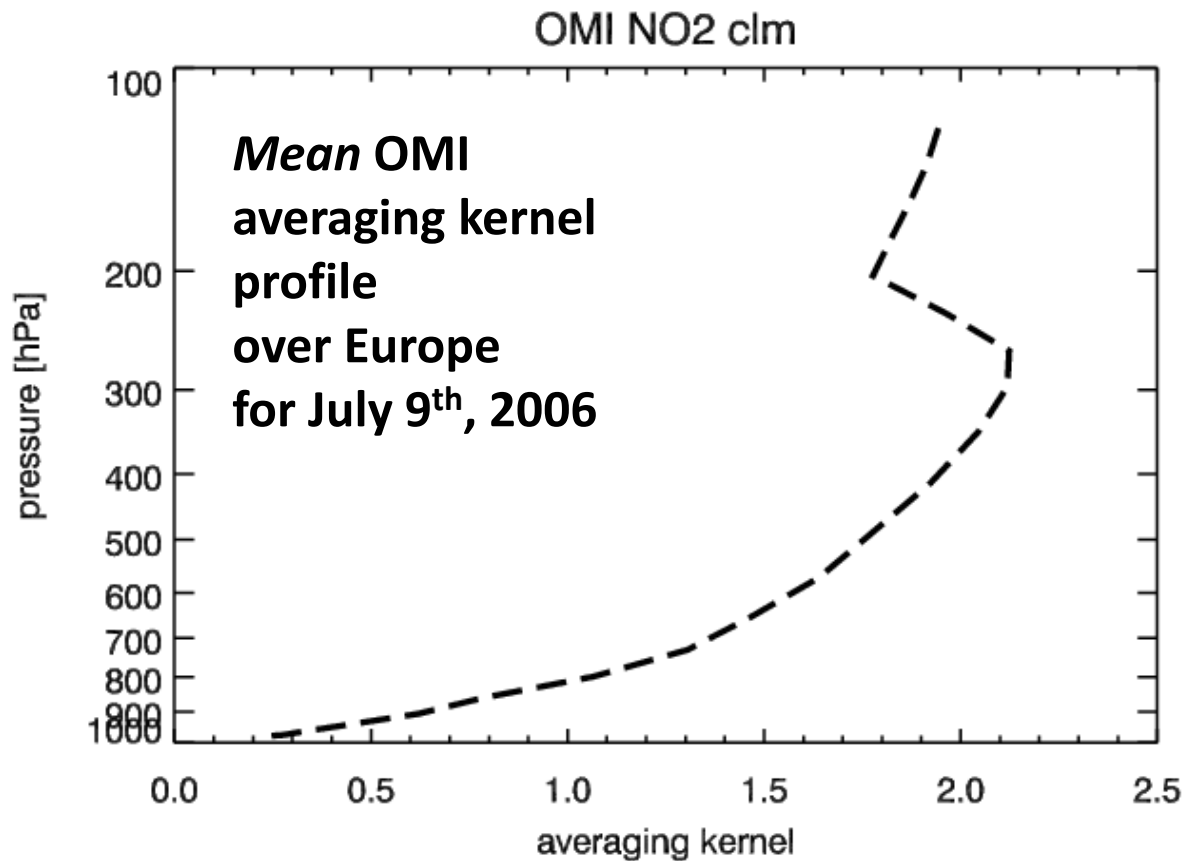
deviations from background initial state

deviations from a priori emission rates

model deviations from observations

- $\mathbf{x}^b(t_0)$ background state at $t = 0$
- $\mathbf{x}(t)$ model state at time t
- $\mathbf{e}_b(t_0)$ background emission rate at $t = 0$
- $\mathbf{e}(t)$ emission rate field at time t
- \mathbf{K} emission rate error covariance matrix
- $H[]$ forward interpolator
- $\mathbf{y}^0(t)$ observation at time t
- \mathbf{B}_0 background error covariance matrix

UV-VIS retrievals: Assimilation by averaging kernels



max. sensitivity of
model domain
mean averaging
kernel
**above boundary
layer!**

How can we still make best use of it?

How to proceed to obtain benefit from trop. column
integral information?

(A typical problem of Inverse Modelling by Integral Equations)

Two more specific questions:

- When is it justified to project averaging kernel information to the surface?
- Can this be done without heuristics, destroying the BLUE property of the assimilation algorithm?

Observation operator \mathbf{H}

Formally an integral equation to be solved for vertical NO_2 molecule density function x (σ vertical coordinate)

$$y = \int_1^0 w(\sigma)x(\sigma)d\sigma$$
$$y = \sum_K h_k x_k$$

At the minimum $\mathbf{x} =: \mathbf{x}_a$

$$d\mathbf{x}_a := \mathbf{x}_a - \mathbf{x}_b = (\mathbf{B}_0^{-1} + \mathbf{H}^T \mathbf{R}^{-1} \mathbf{H})^{-1} \mathbf{H}^T \mathbf{R}^{-1} \{ \mathbf{y}^0 - H[\mathbf{x}_b] \}$$
$$= \mathbf{B} \mathbf{H}^T (\mathbf{R} + \mathbf{H} \mathbf{B} \mathbf{H}^T)^{-1} \{ \mathbf{y}^0 - H[\mathbf{x}_b] \}$$

For scalar column retrieval:

$$d\mathbf{x}_a^c = \underbrace{\mathbf{B} \mathbf{h}^T}_{(r+b)} (r+b)^{-1} \{ y^0 - H[\mathbf{x}_b] \}$$

adjoint representer

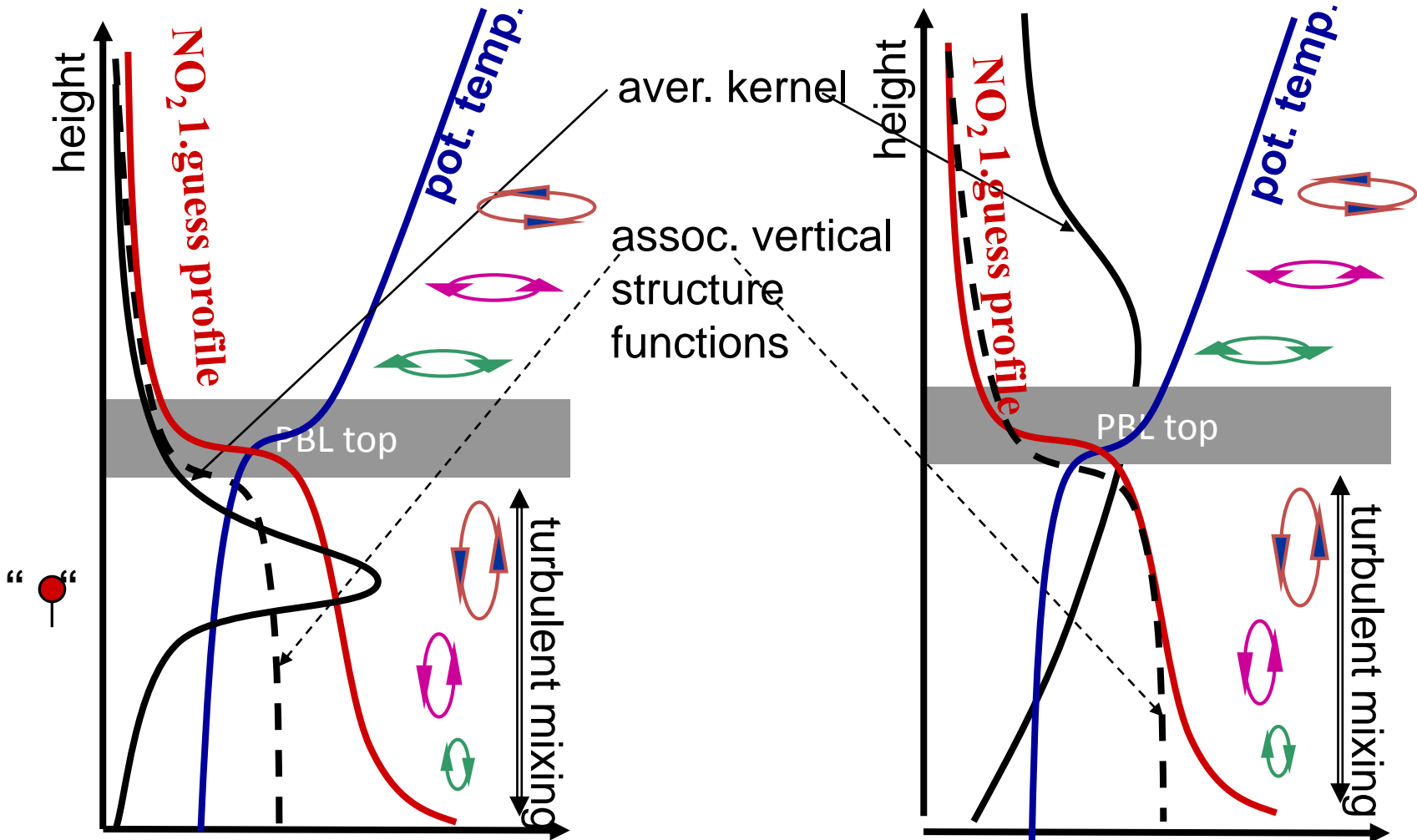
→ vertical structure function in \mathbf{B} essential!

4. Focus: joint emission rate initial value optimisation

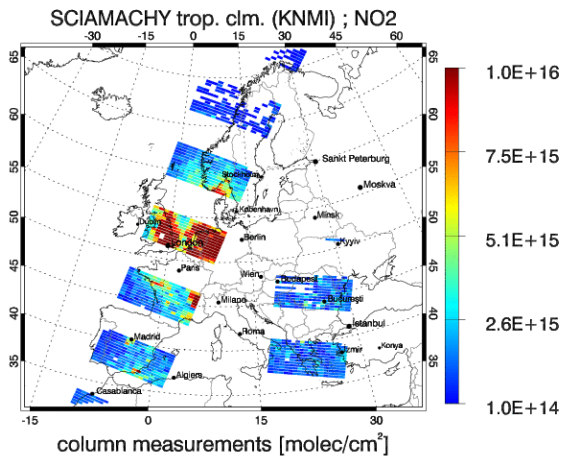
Vertical structure function:

Extending the information from observation location by vertical exchange of pollutants and information

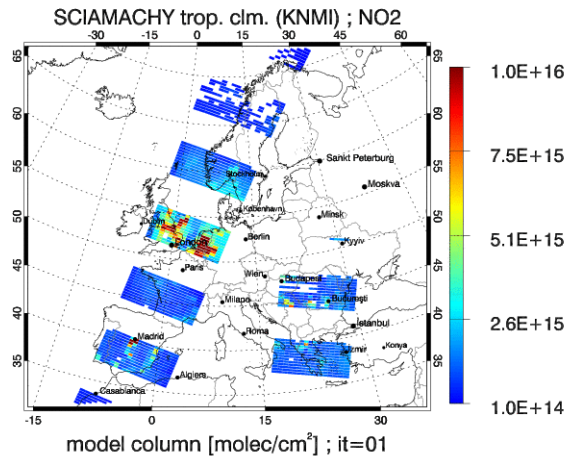
ideal *case* *real*



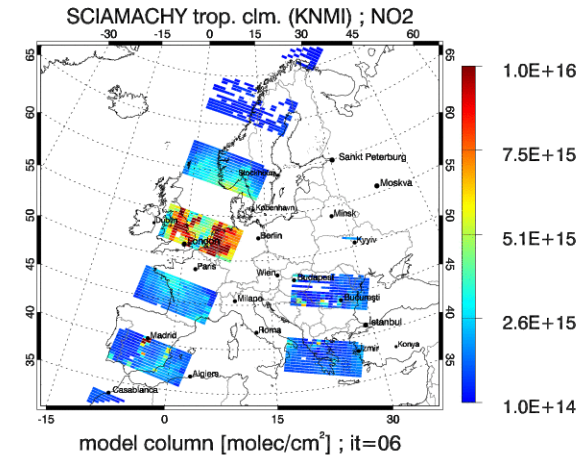
Comparison of NO₂ tropospheric columns in molecules/cm² for July 6th, 2006, 09-12 UTC.



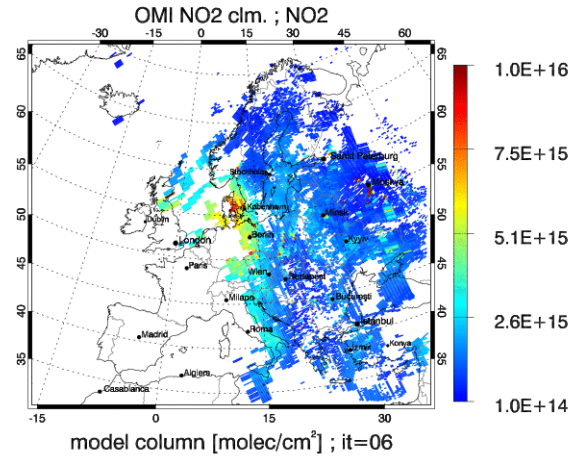
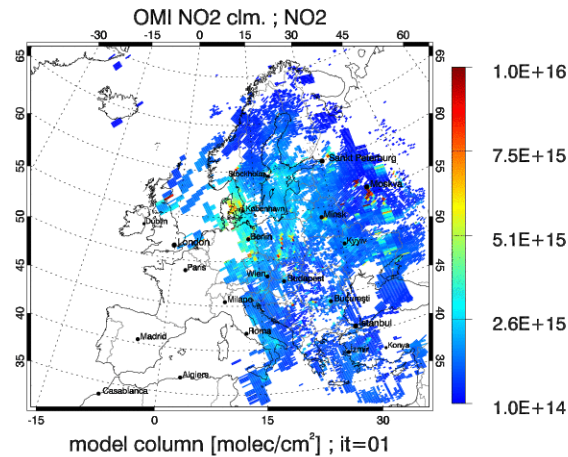
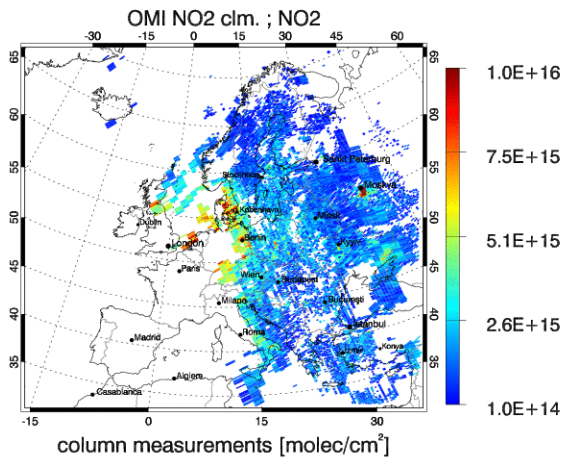
retrievals (\mathbf{y});



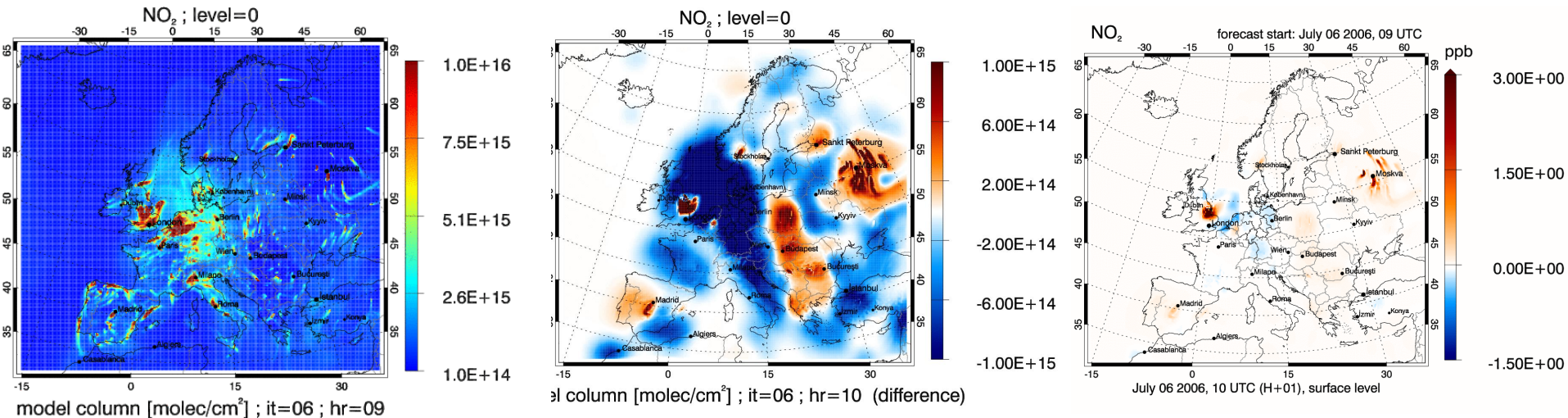
:EURAD forecasted (\mathbf{Hx}_b);



column analyses (\mathbf{Hx}_a).

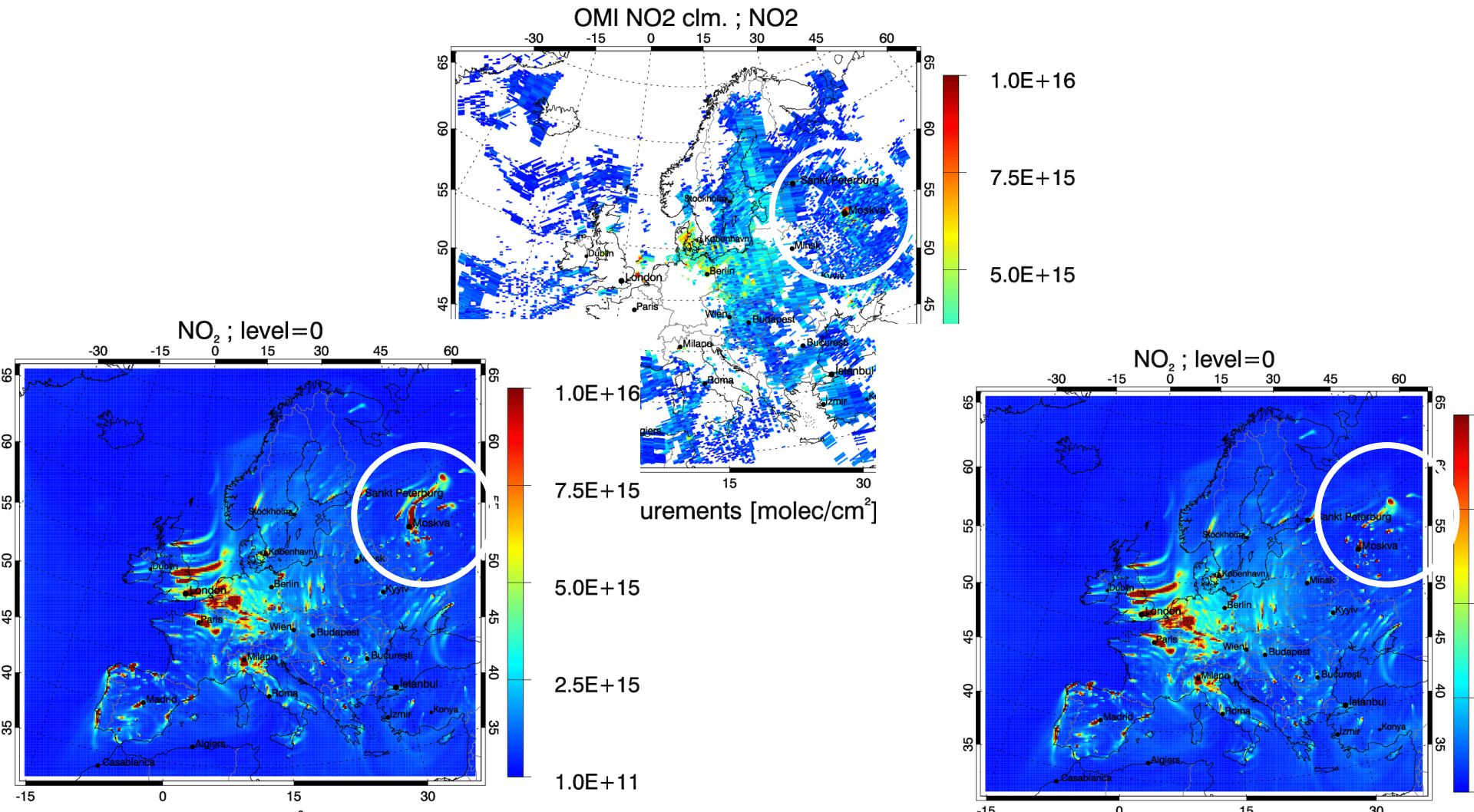


Data assimilation result in terms of tropospheric columns for July 6th, 2006. NO₂ model columns based on OMI and SCIAMACHY assimilation within interval, 09-12 UTC.

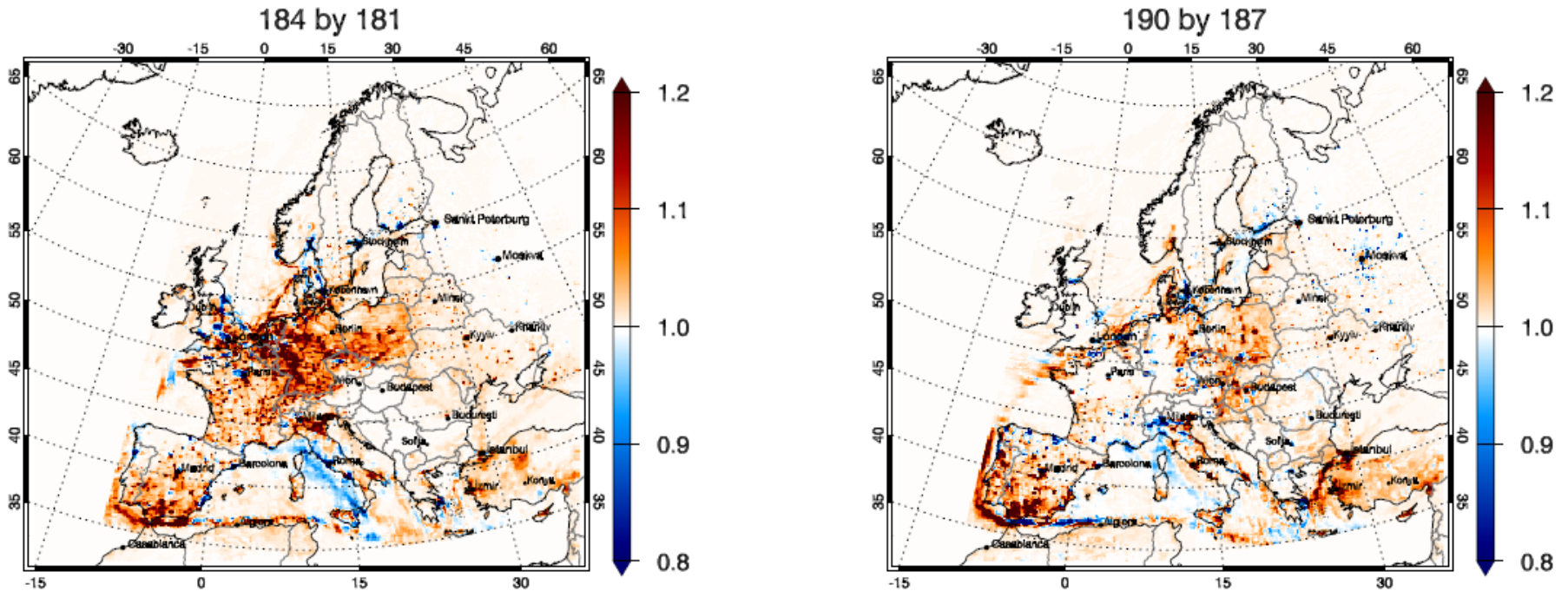


Difference field giving implied changes for tropospheric columns by assimilation (middle), and induced surface concentration changes by NO₂ ppb (right)

Data assimilation result in terms of tropospheric columns for **July 7th, 2006**. NO₂ model columns based on OMI and SCIAMACHY assimilation within the assimilation interval, 09-12 UTC.



Emission rate optimisation factors for NO₂ after assimilation of OMI retrieved NO₂ tropospheric columns

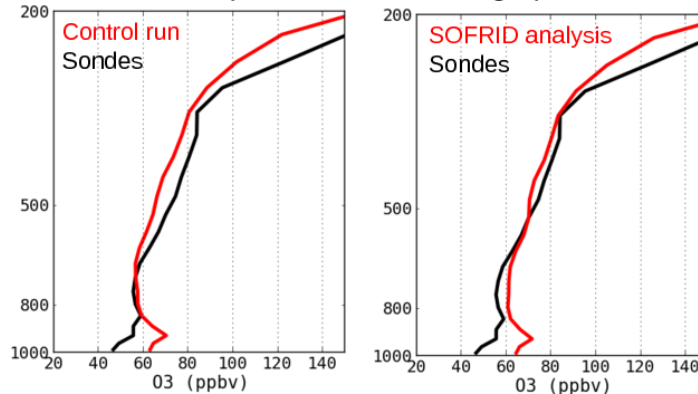


2 x 4 days assimilation sequence. Left panel shows results after assimilation procedures from July 1.-4. 2006, right panel for July 7.-11., 2006. OMI data from KNMI

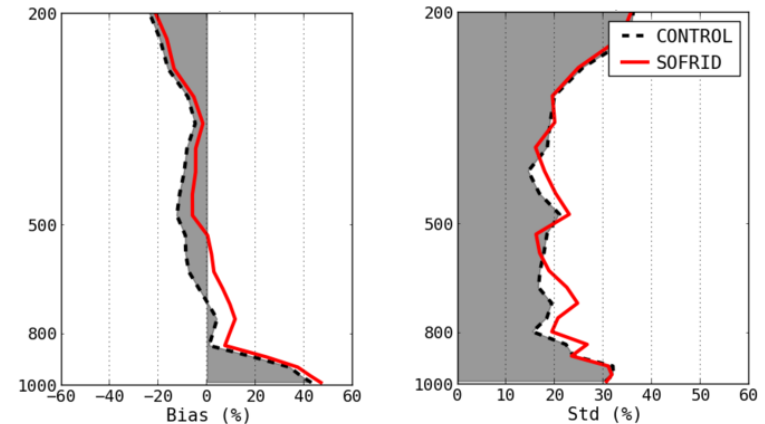
IASI SOFRID O₃ re-analysis (CERFACS)

O₃ profiles in July 2010:

European sondes average profiles

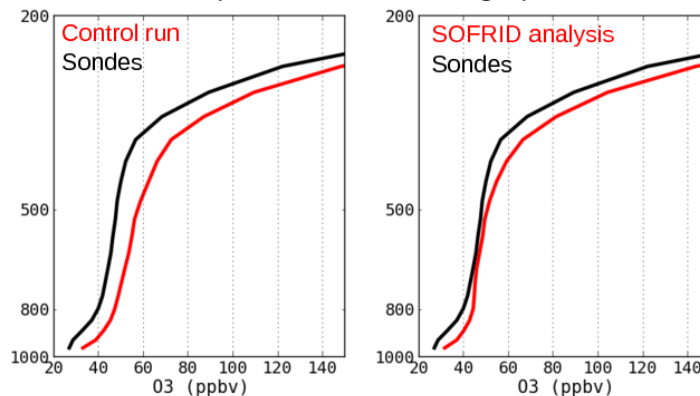


Model-Sondes differences

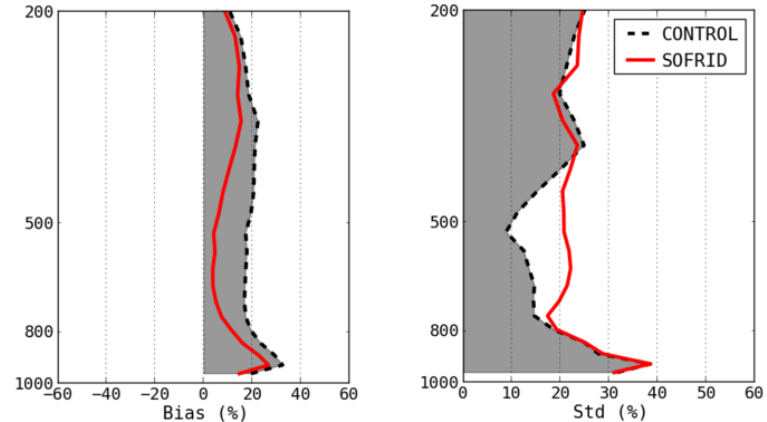


O₃ profiles in Jan 2012:

European sondes average profiles



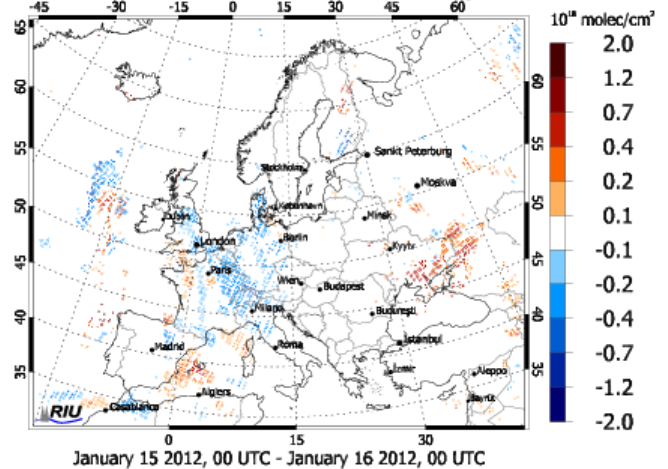
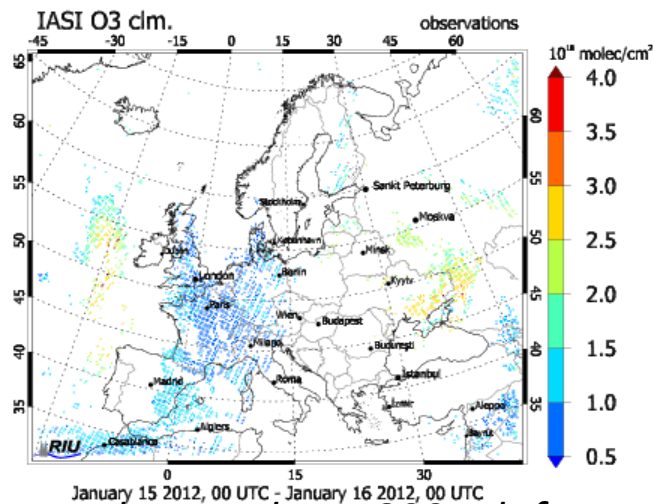
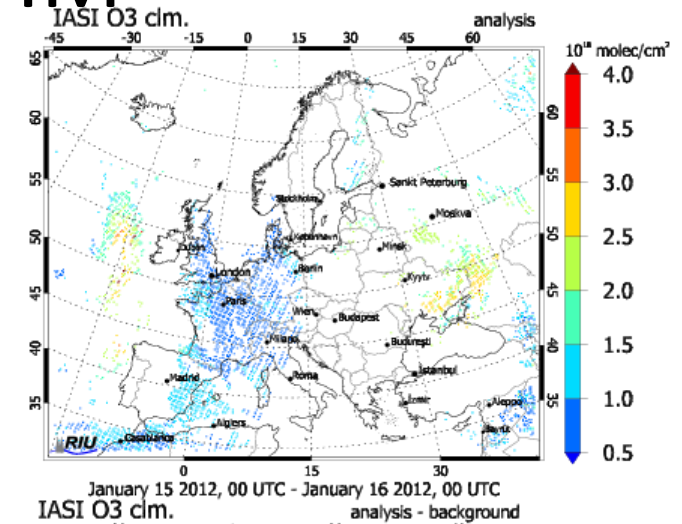
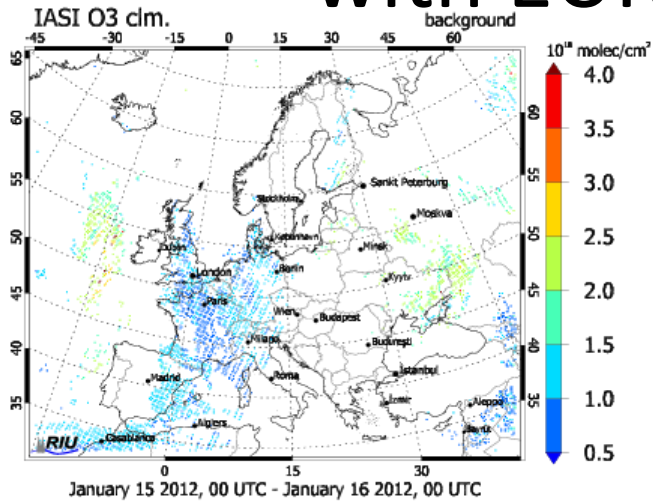
Model-Sondes differences



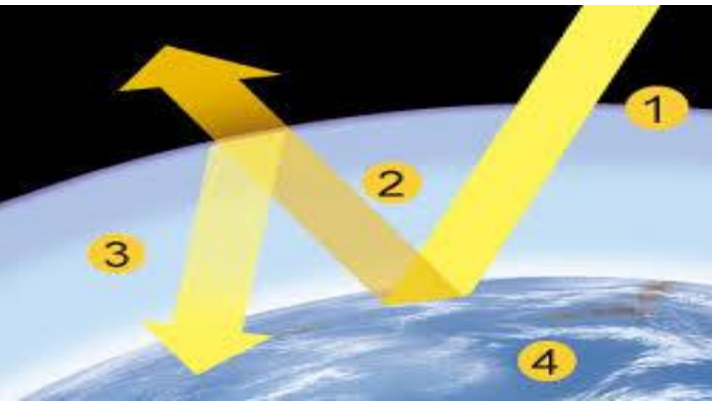
- Bias reduced in the free troposphere
- Surface ozone impact is minor
- MOZAIC-IAGOS as additional validation? (only 2012 available)

Courtesy: E. Emili, CERFACS

IASI partial Ozone column assimilation with EURAD-IM



IASI partial Ozone columns above 800 mb for January 15, 2012 interpolated on the EURAD-IM grid with 15 km resolution.



GREENHOUSE GAS INVERSION

CO2 satellite inversion

- small signals relative to large background values enforce accuracy requirements at the limits of today's space borne spectroscopy
- preferred assimilation method is 4D-var for source/sink inversion

Satellite data types and status resumé

- thermal infra-red spectral domain, with a peak sensitivity in the middle troposphere
 - AIRS, IASI, TES, GOSAT-TANSO
- solar infra-red domain with a more uniform sensitivity to GHGs throughout the atmospheric column, including the boundary layer.
 - SCIAMACHY, GOSAT-TANSO, OCO-2
- of both measurement types in the 12-hour 4D-Var useful,
- but: systematic errors caused by **uncertainties in the spectroscopy** or by **aliasing with other atmospheric signals** like aerosols
- **the inversion of CO₂ surface fluxes from satellite retrievals up to now hampered.**

(from F. Chevallier, MACC report 2014)

GOSAT observations

Characteristics and error estimates

(example from Chevallier et al., 2010)

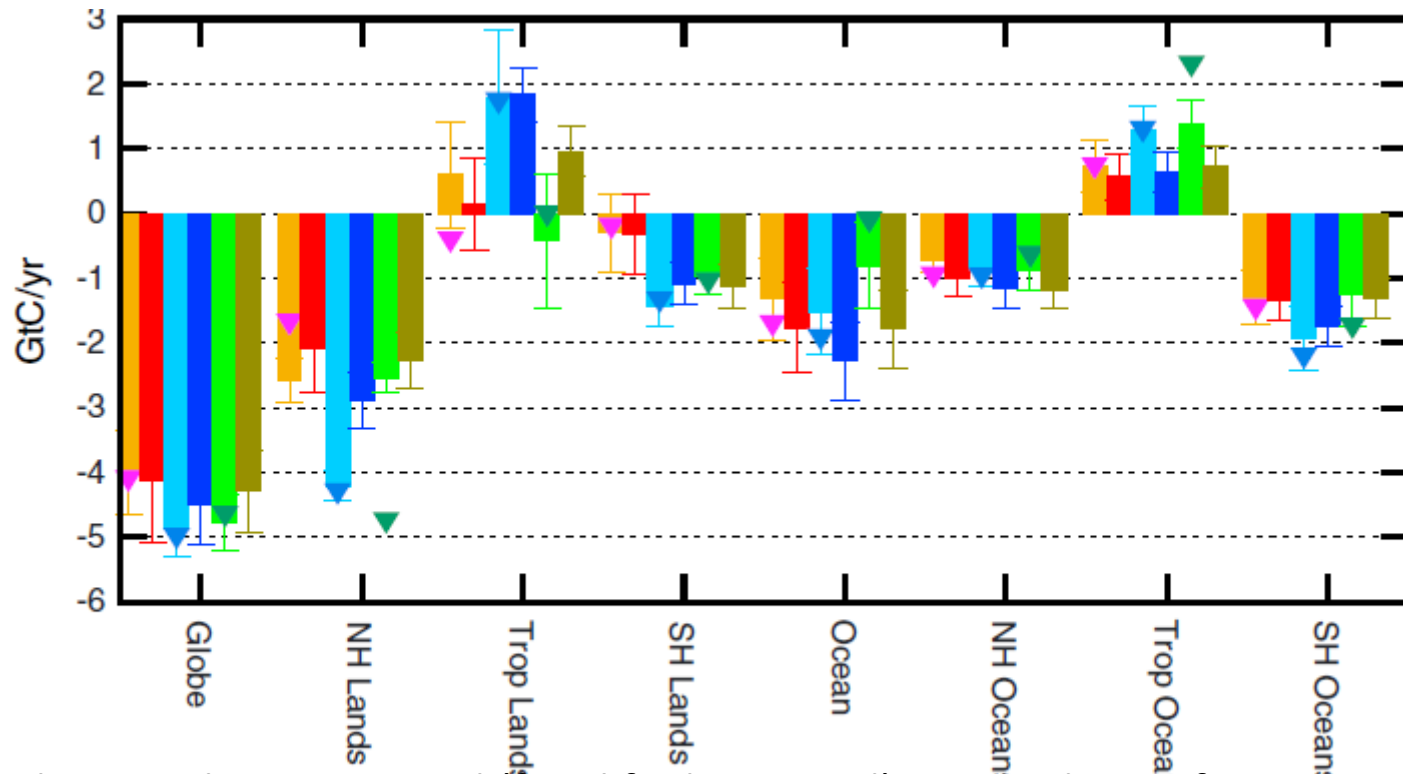
Features

- sun-synchronous GOSAT radiation data in the near infrared
- 12-level averaging kernels for XCO₂ retrieval
- simultaneous fit to 2 CO₂ bands (1.61, 2.06 μm) and O₂-A Band (0.765 μm)
- quality check results in $\sim 300,000$ soundings/a ($\sim 1/20$ of total)

error budget

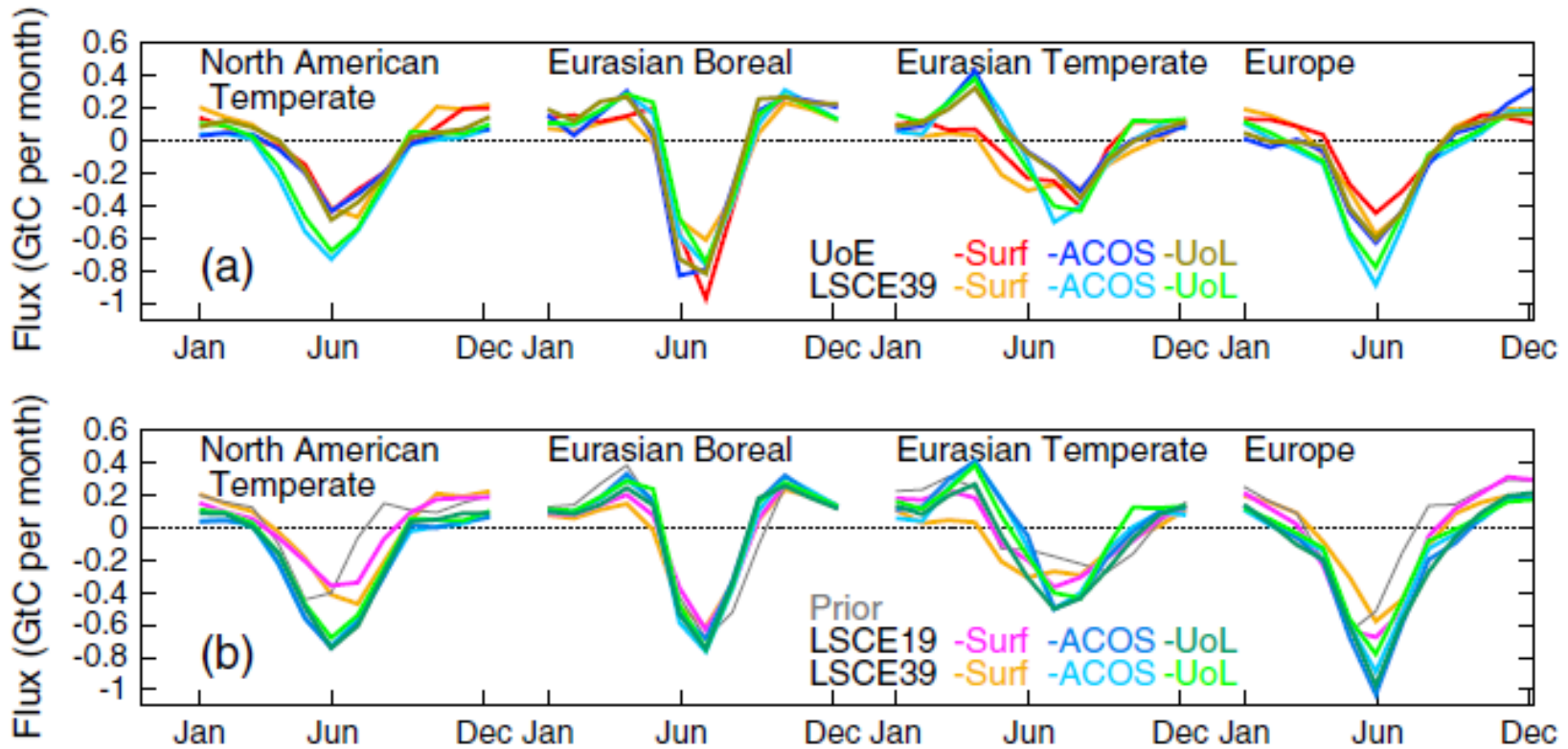
- retrieval error scene-specific
- 1 ppm to errors associated with the radiation model, representativity, and transport model \rightarrow
- **1.8 - 7.2 ppm total error, needed are < 0.5 ppm**

Fluxes derived from GOSAT XCO₂ backscattered sunlight at short-wave IR for 2010



Large-scale annual mean natural (fossil fuel removed) CO₂ budgets of inversions
 Comparison of the UoE (Kalman F. University of Edinburgh)
 LSCE-39 (variational ,Laboratoire des Sciences du Climat et de l'Environnement)
 ACOS (Bayesian; NASA), UoL (Bayesian; Univ. of Leicester) results.

Differences



Regional seasonal CO₂ flux estimates 2010

Chevallier et al, 2014

Comparison of the UoE (Kalman F. University of Edinburgh)

LSCE-39 (variational ,Laboratoire des Sciences du Climat et de l'Environnement)

ACOS (Bayesian; NASA), UoL (Bayesian; Univ. of Leicester) results.

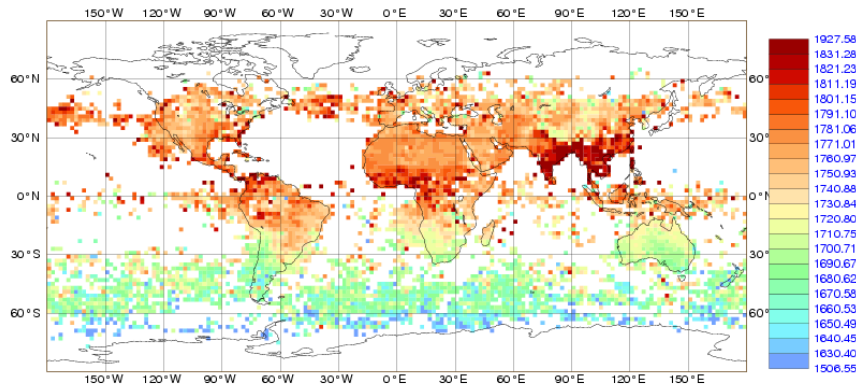
(b) Comparison of the LSCE-19 results, the LSCE-39 results, and the LSCE prior fluxes.

ECMWF-MACC:

Assimilated GHG satellite data

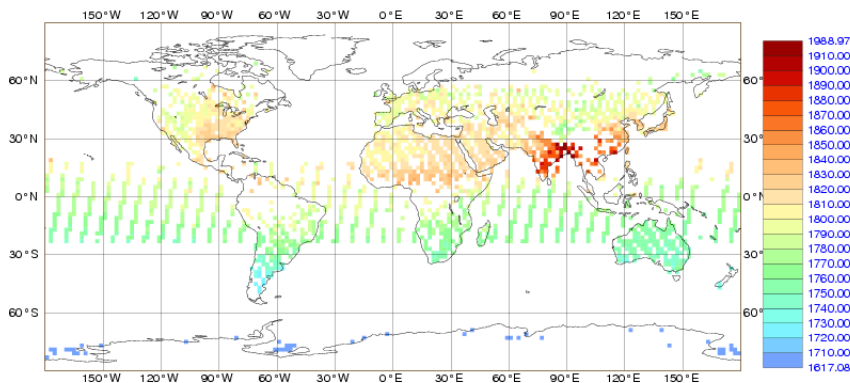
ENVISAT/SCIAMACHY

CH₄ and CO₂ – Lower tropo.



GOSAT/TANSO

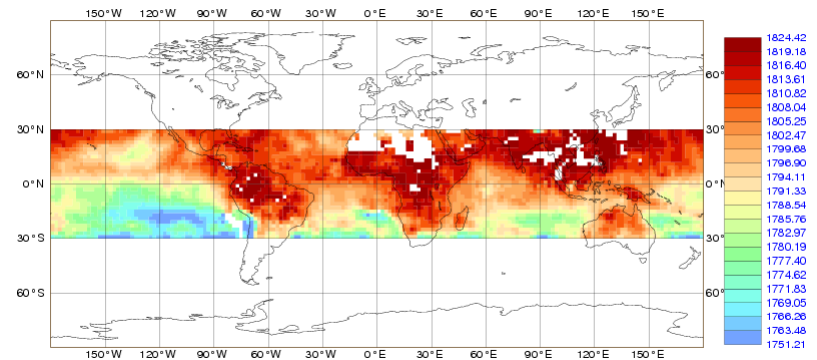
CH₄ and CO₂ – Lower tropo.



Monthly averages of the observed XCH₄ for October 2011

METOP-A/IASI

CH₄ and CO₂ – Middle tropo.



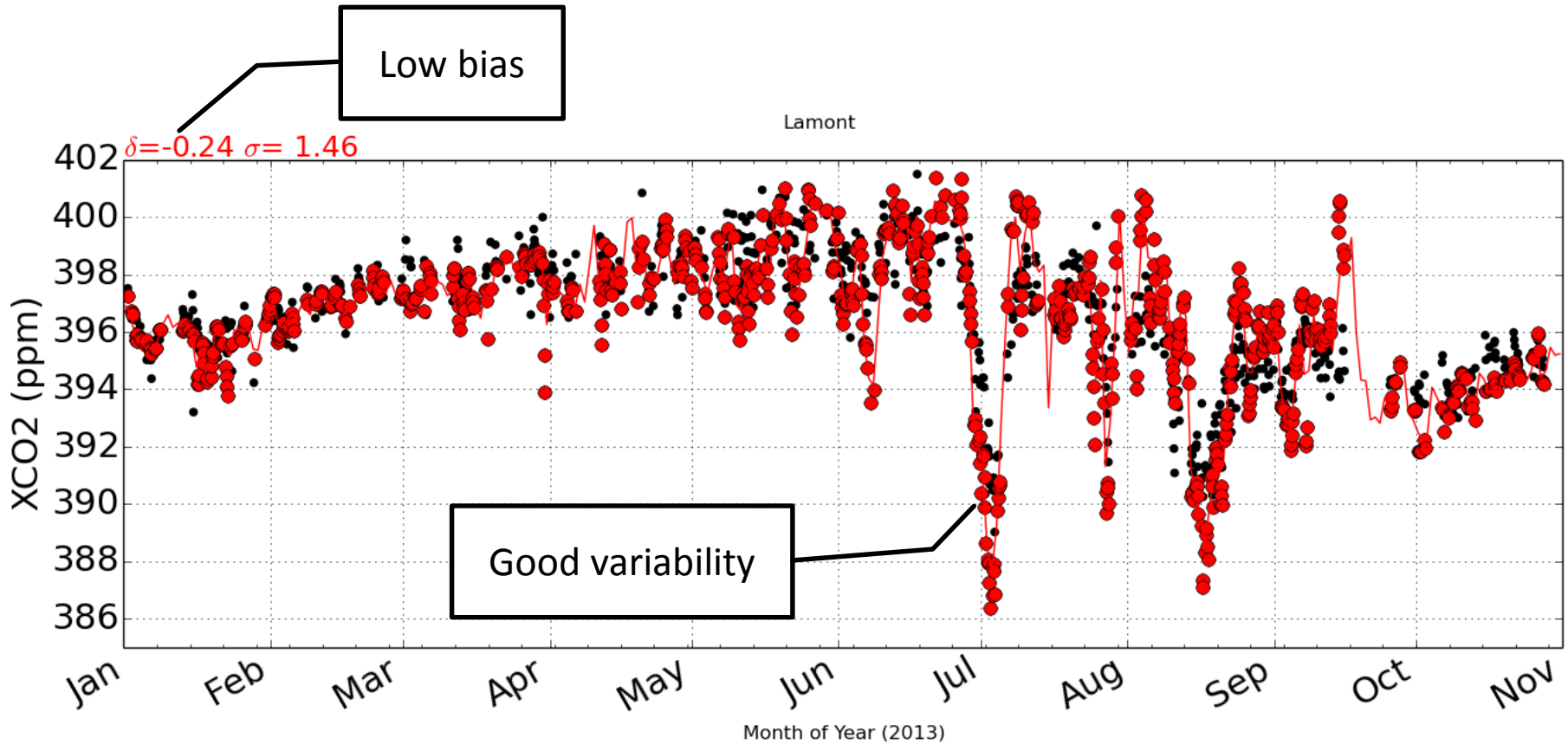
Column-averaged dry-air mole fractions of CO₂ and CH₄ provided by:

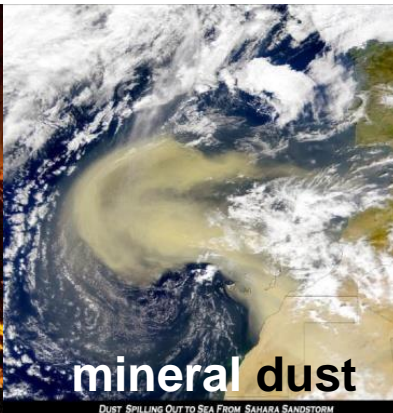


Tests with AIRS and IASI radiances for CO₂

Courtesy S. Massart@ECMWF

Zoom over 2013

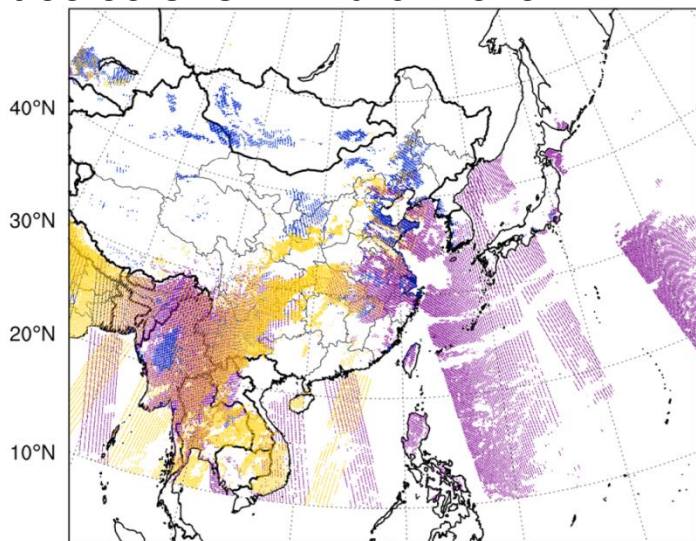




SATELLITE AEROSOL DATA ASSIMILATION

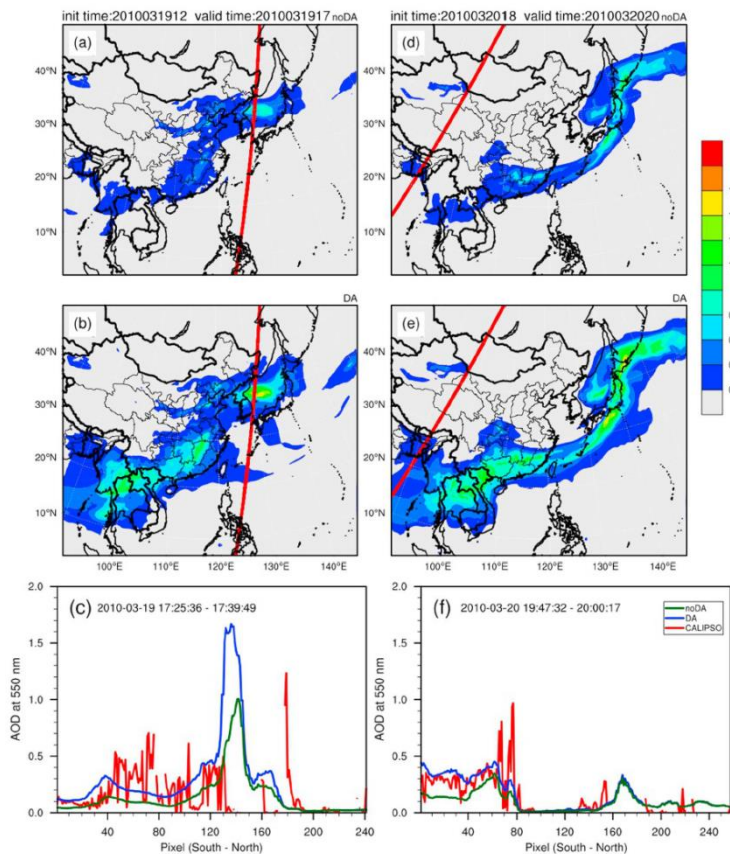
Three-dimensional variational Assimilation of MODIS aerosol optical depth:

(MODIS) AOD coverage from the Aqua and Terra at 06:00 UTC 21 March 2010.



Purple: dark surface retrievals from Aqua;
gold: dark surface Terra;
blue: deep blue produced from Aqua.

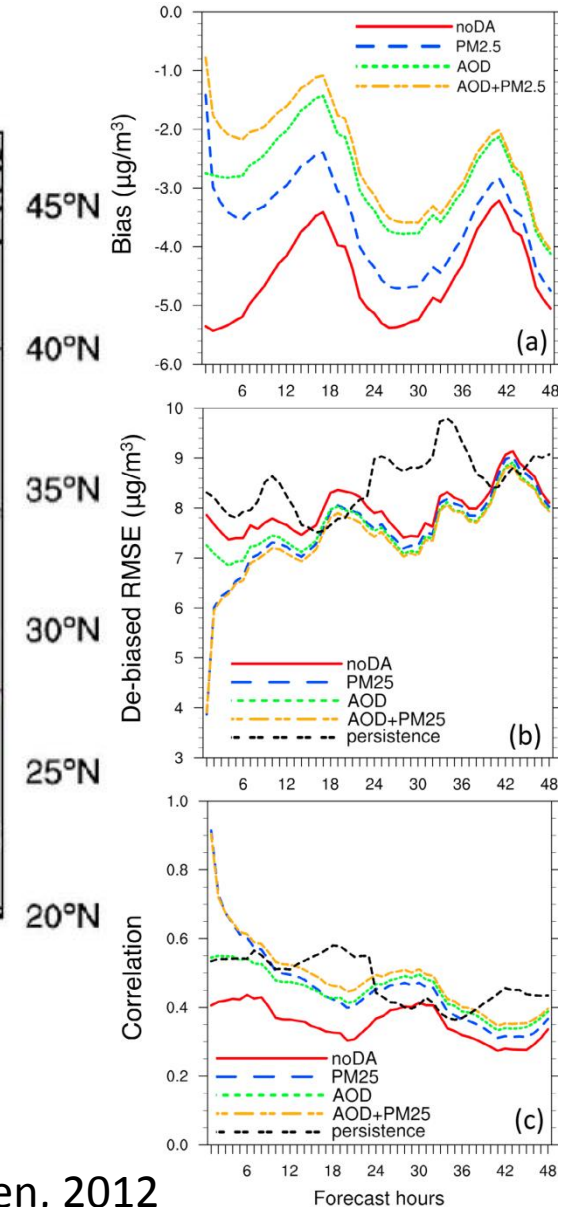
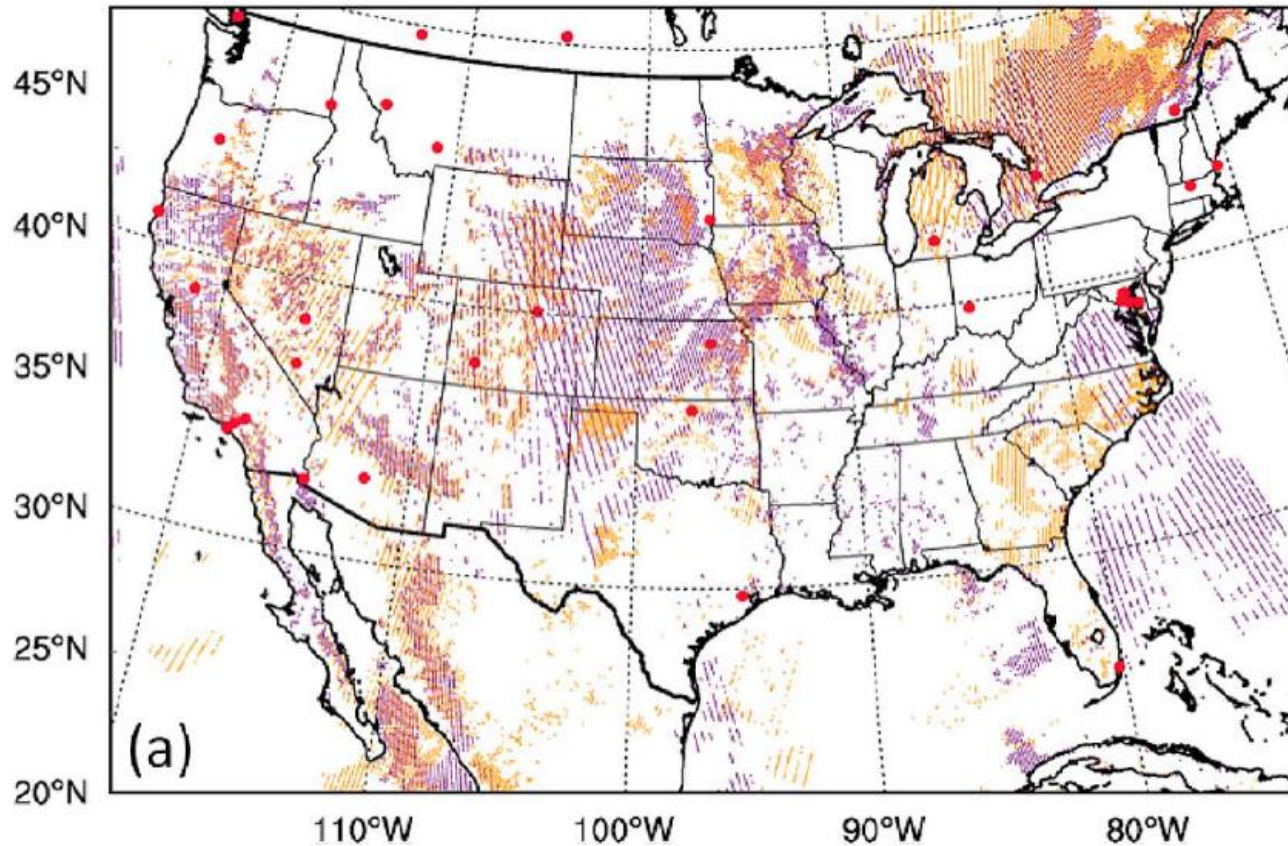
Validation with CALIPSO.



alongtrack assimilation control

Zhiquan Liu, Quanhua Liu, Hui-Chuan Lin,1Craig S. Schwartz, Yen-Huei Lee, and Tijian Wang, 2011

MODIS AOD assimilation



Can we bridge from optical to chemical information?

Example: Aerosol Chemistry in **MADE**

Modal Aerosol Dynamics for
EURAD/Europe
(Ackermann et al., 1998,
Schell 2000)

$$dM_i^k/dt = \text{nuk}_i^k + \text{coag}_{ji}^k + \text{coag}_{ij}^k + \text{cond}_i^k + \text{emi}_i^k$$

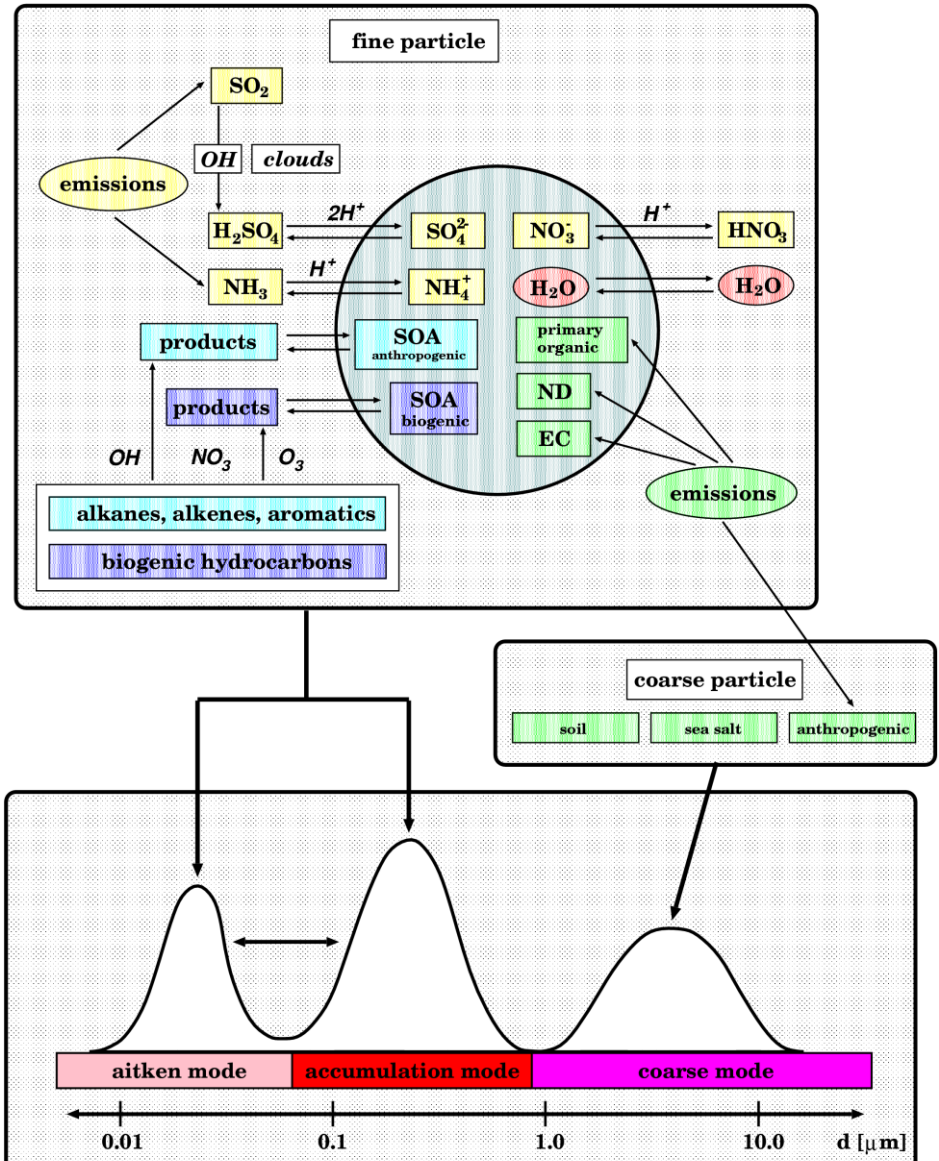
$M_i^k := k^{\text{th}}$ Moment of i^{th} Mode

assimilation of aerosol

By satellite retrievals: e.g.

MERIS MODIS

AATSR+SCIAMACHY, ...



Assimilation of Aerosol observations

- In situ:

EEA Airbase: Database of groundstations of EU member countries & states:

- 450 stations for PM₁₀ (2003)
- No PM_{2.5}. (4 stations in UK only)

- Satellite measurements:

SYNAER (SYNERgetic AErosol Retrieval, DLR-DFD, [Holzer-Popp, 2001])*

- combines GOME&ATSR-2, SCIAMACHY&AATSR measurements aboard ERS-2/ENVISAT

- ATSR-2/AATSR:

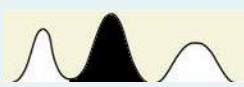




dark field detection, BLAOT (Boundary Layer Aerosol Optical Thickness) and albedo are calculated

- GOME/SCIAMACHY:

Provides PM_{0.5}, PM_{2.5} and PM₁₀ columns and its composition (6 intrinsic species)

SYNAER retrieval algorithm

Species Mapping

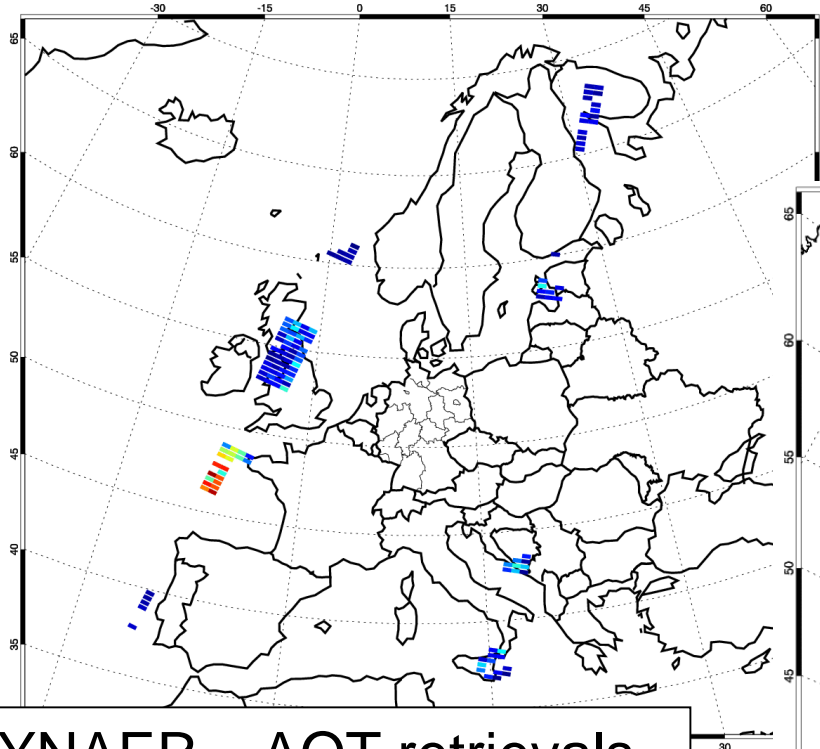
| EURAD-IM [$\mu\text{g}/\text{m}^3$] | SYNAER - AOT |
|---|------------------------|
| SO ₄ , NH ₃ , NO ₃ , H ₂ O, SOA  | WASO (WATER SOLuble) |
| Unidentified PM  | INSO (water INSOLuble) |
| Elemental Carbon  | SOOT |
| Sea Salt  | SEAS |
| Mineral Dust  | DUST |

radiative transfer model



adjoint radiative transfer model

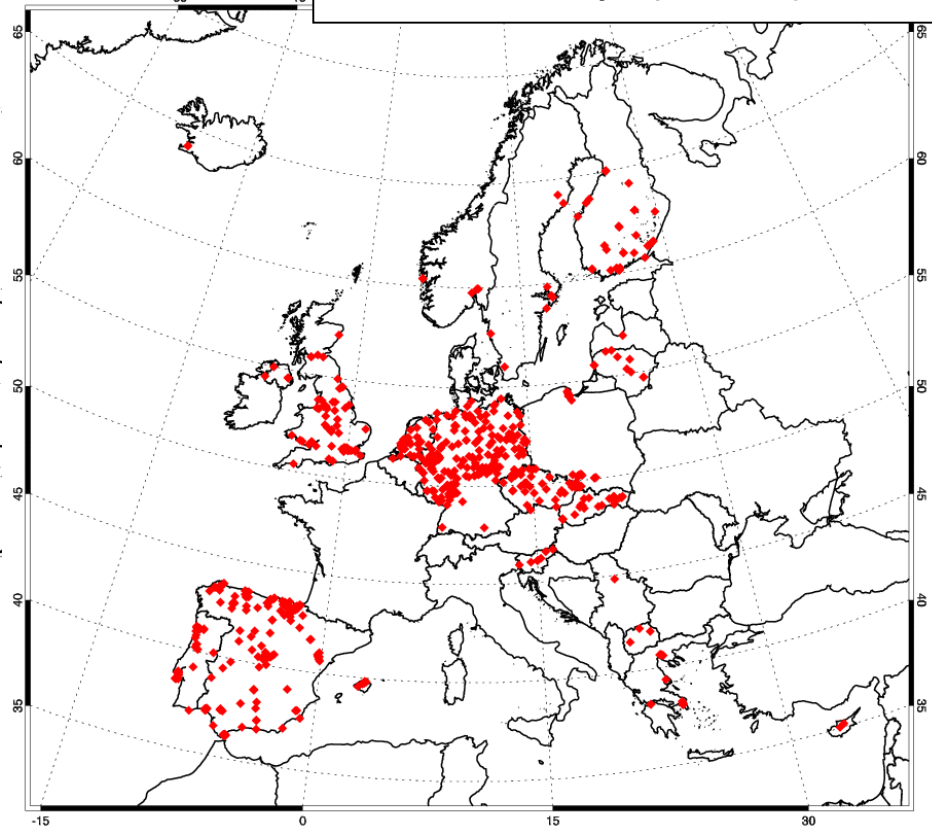
Model Domain and Observations



SYNAER – AOT retrievals
2 – 3 traces per day

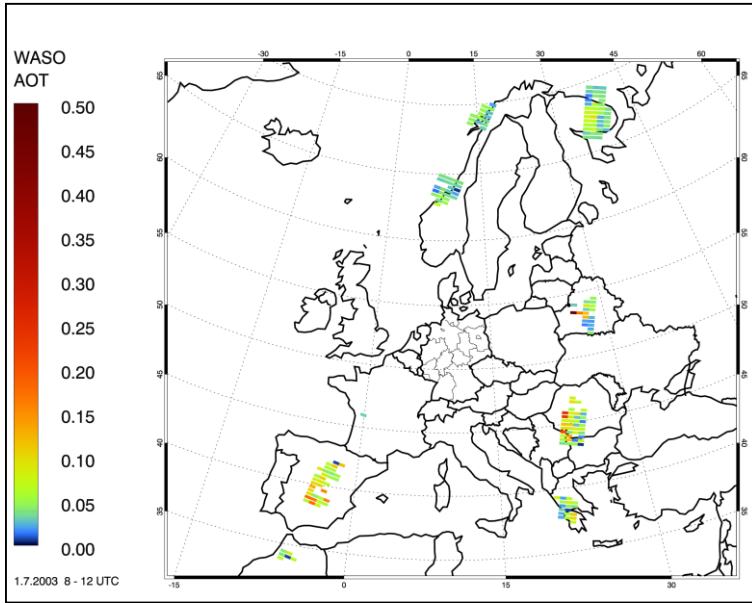
13 July 2003

EEA in-situ PM₁₀
hourly (2003)



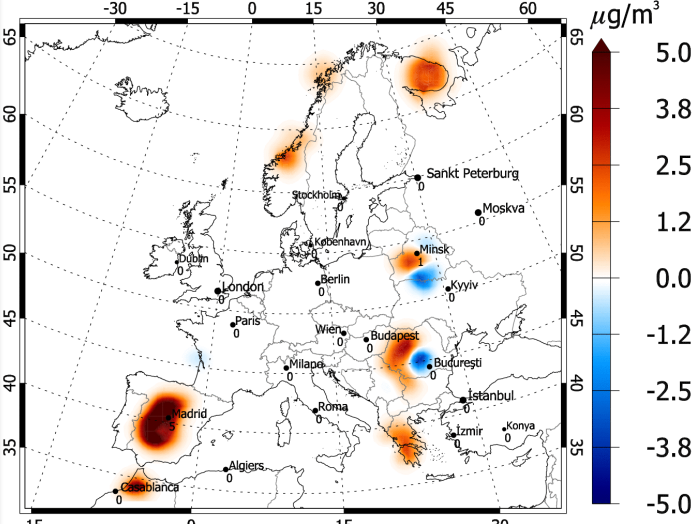
Analysis and increments for acc. SO_4^{2-} July 1, 8 UTC

SYNAER AOT-WASO



VSO4AJ

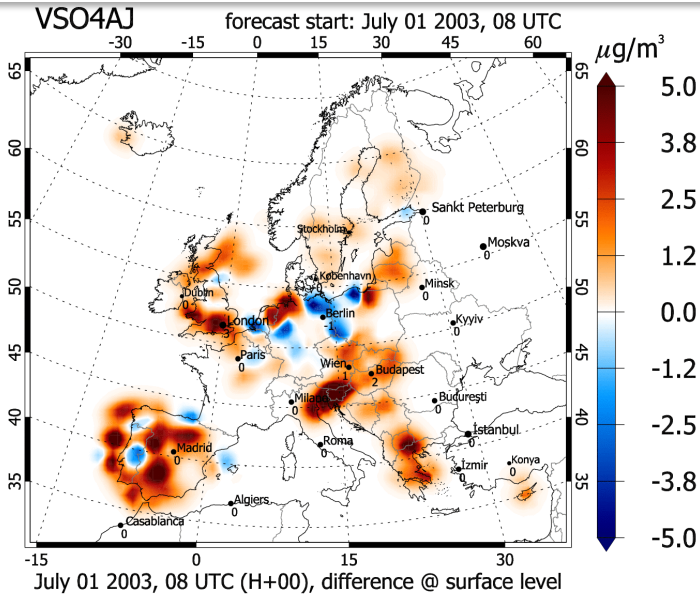
forecast start: July 01 2003, 08 UTC



July 01 2003, 08 UTC (H+00), difference @ surface level

Analysis – Background
SYNAER only

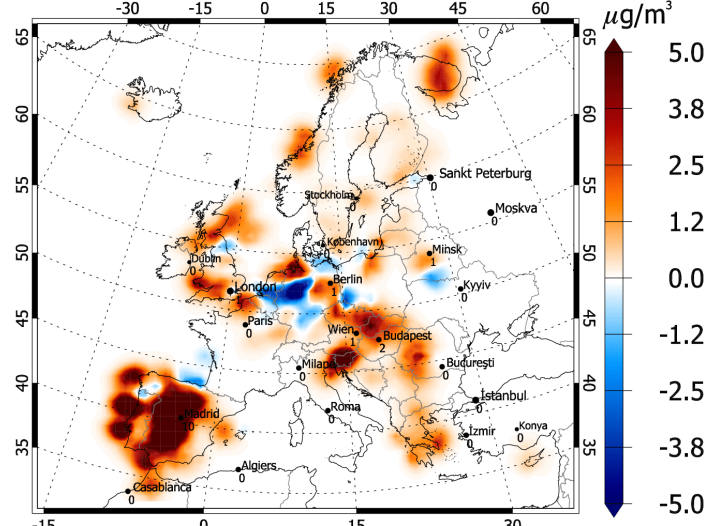
Analysis – Background
EEA only



July 01 2003, 08 UTC (H+00), difference @ surface level

VSO4AJ

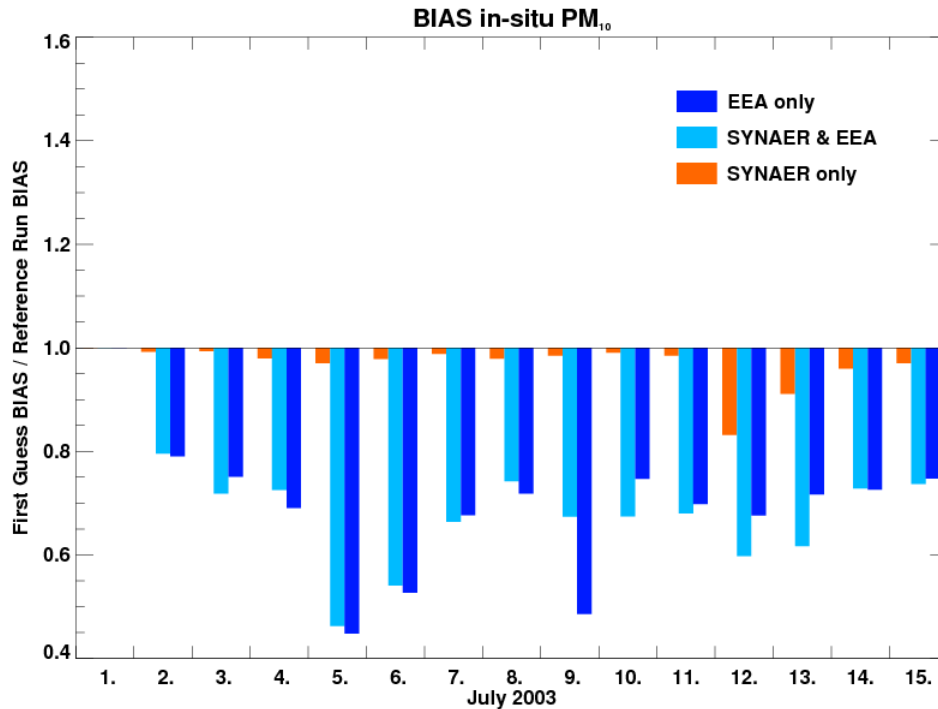
forecast start: July 01 2003, 08 UTC



July 01 2003, 08 UTC (H+00), difference @ surface level

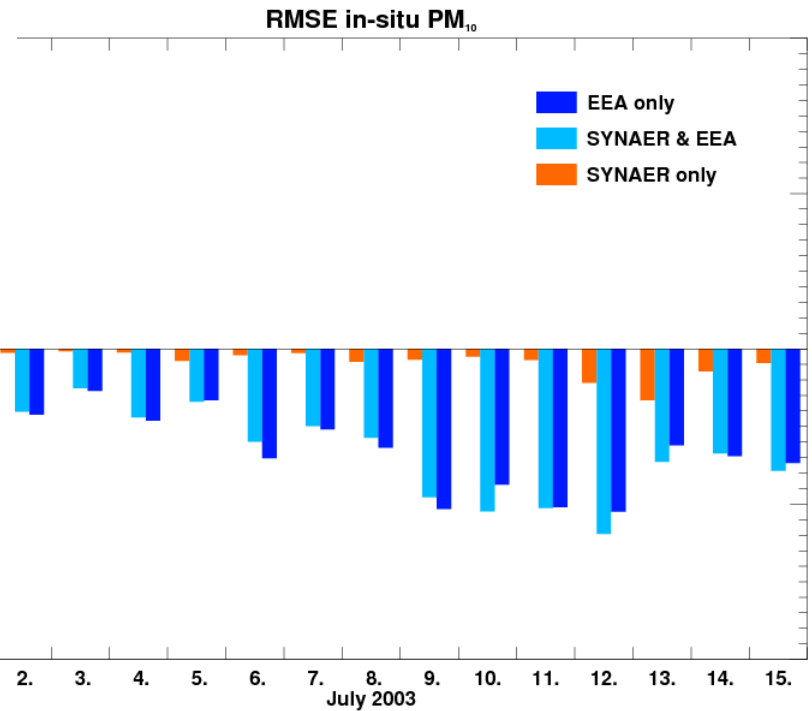
Analysis – Background
SYNAER & EEA

Development of forecast performance



Overall BIAS reduction

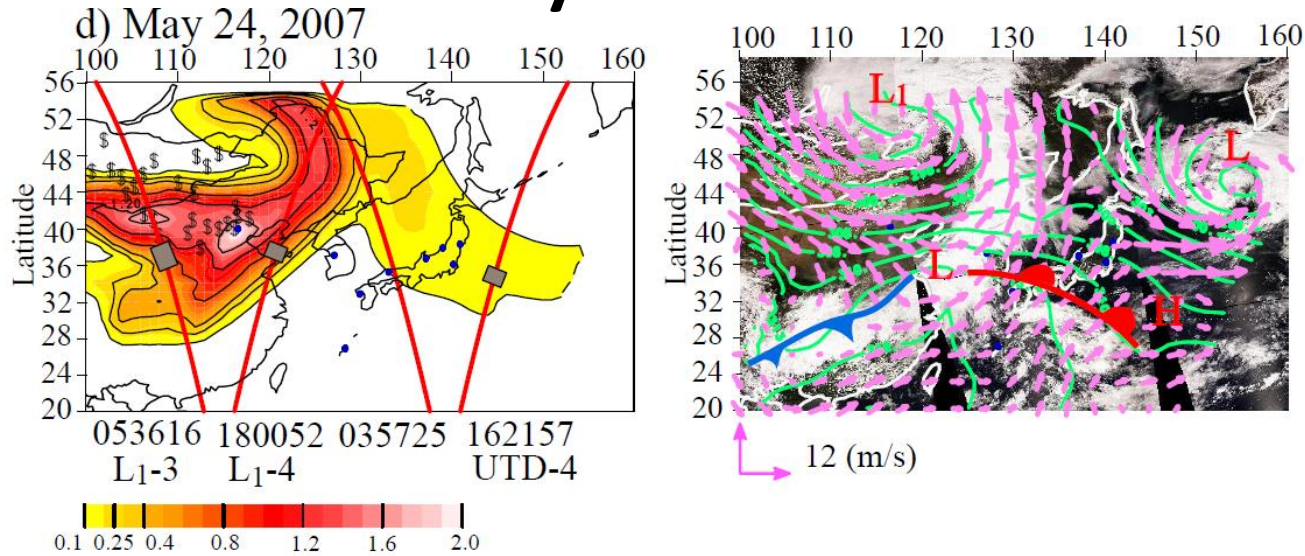
| | |
|-----------|------|
| SYNAER | 0.96 |
| SYN & EEA | 0.67 |
| EEA | 0.67 |



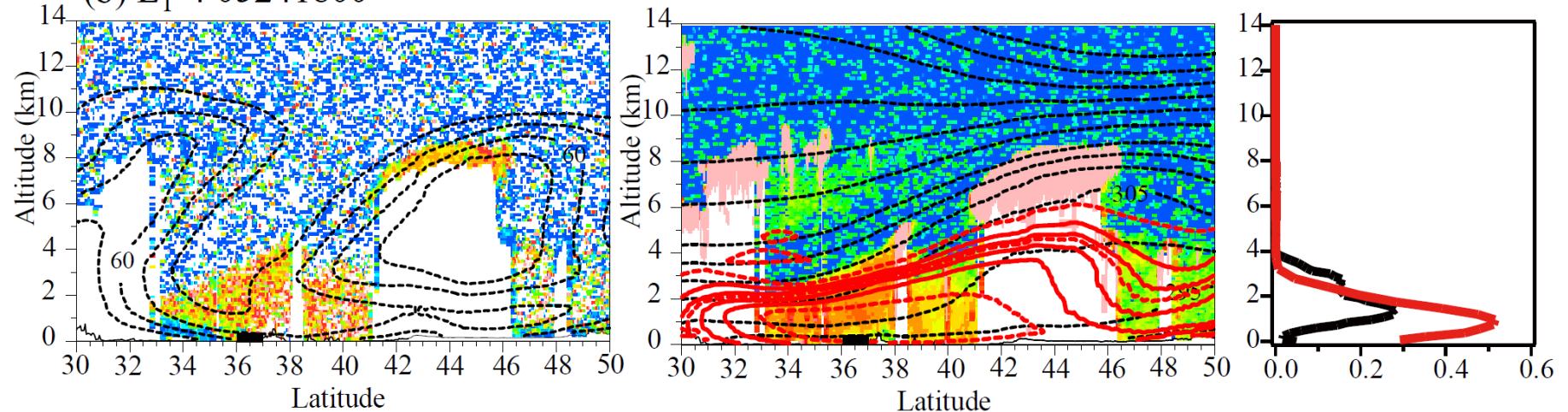
Overall RMSE reduction

| | |
|-----------|------|
| SYNAER | 0.99 |
| SYN & EEA | 0.93 |
| EEA | 0.92 |

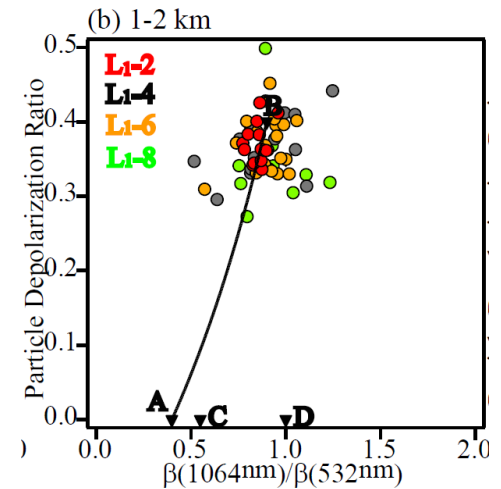
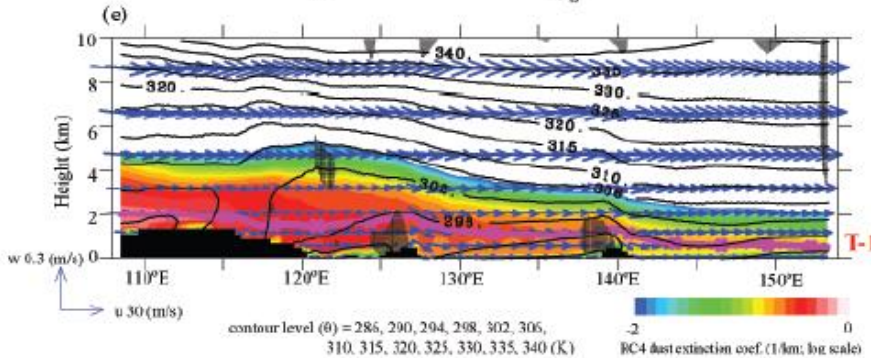
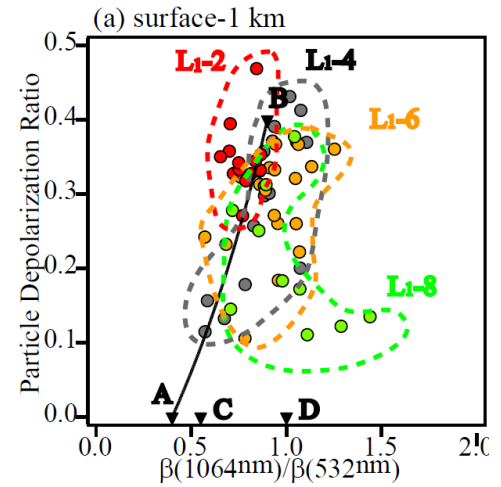
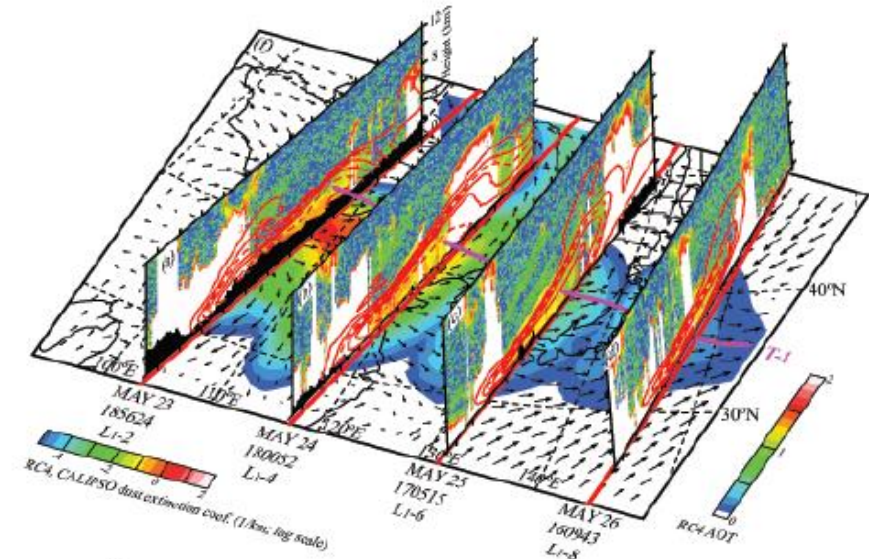
Asian Dust by 4D-var CALIOP DA



(b) L₁-4 05241800



Asian Dust by 4D-var CALIOP DA

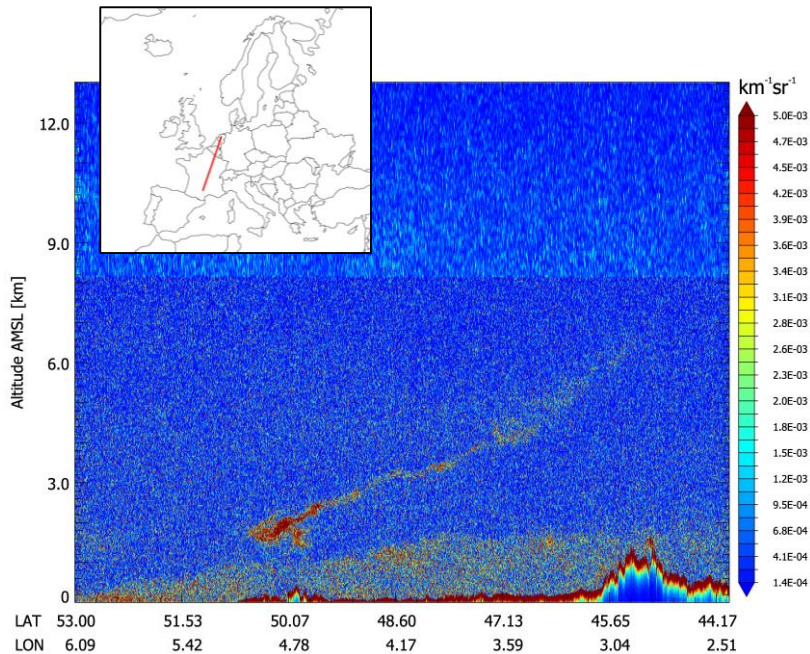


Example CALIOP

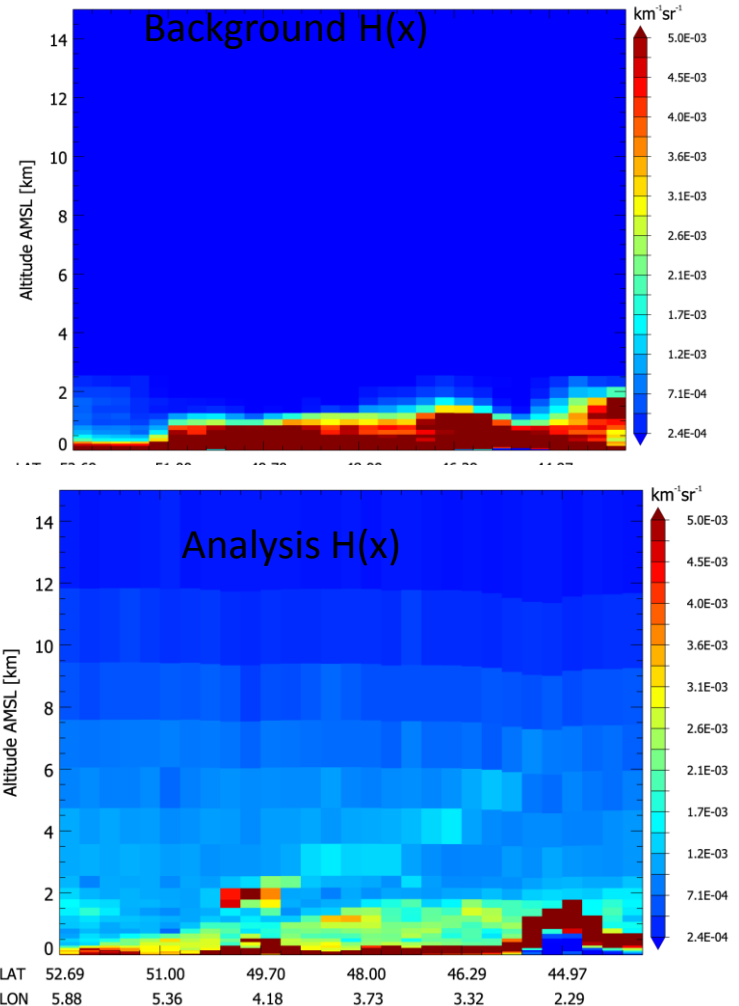
Variational volcanic ash data Assimilation Module with selective background weakening for special events

LiDAR 4D-var data assimilation for improved analysis of unexpected aerosol events
→ automatable online adaptation of background error covariance matrix

CALIOP observation of the Eyjafjallajökull ash cloud
17 April 2010, 02:01:19 - 02:14:53 UTC (Winker et al., 2012)



A. Lange, master thesis

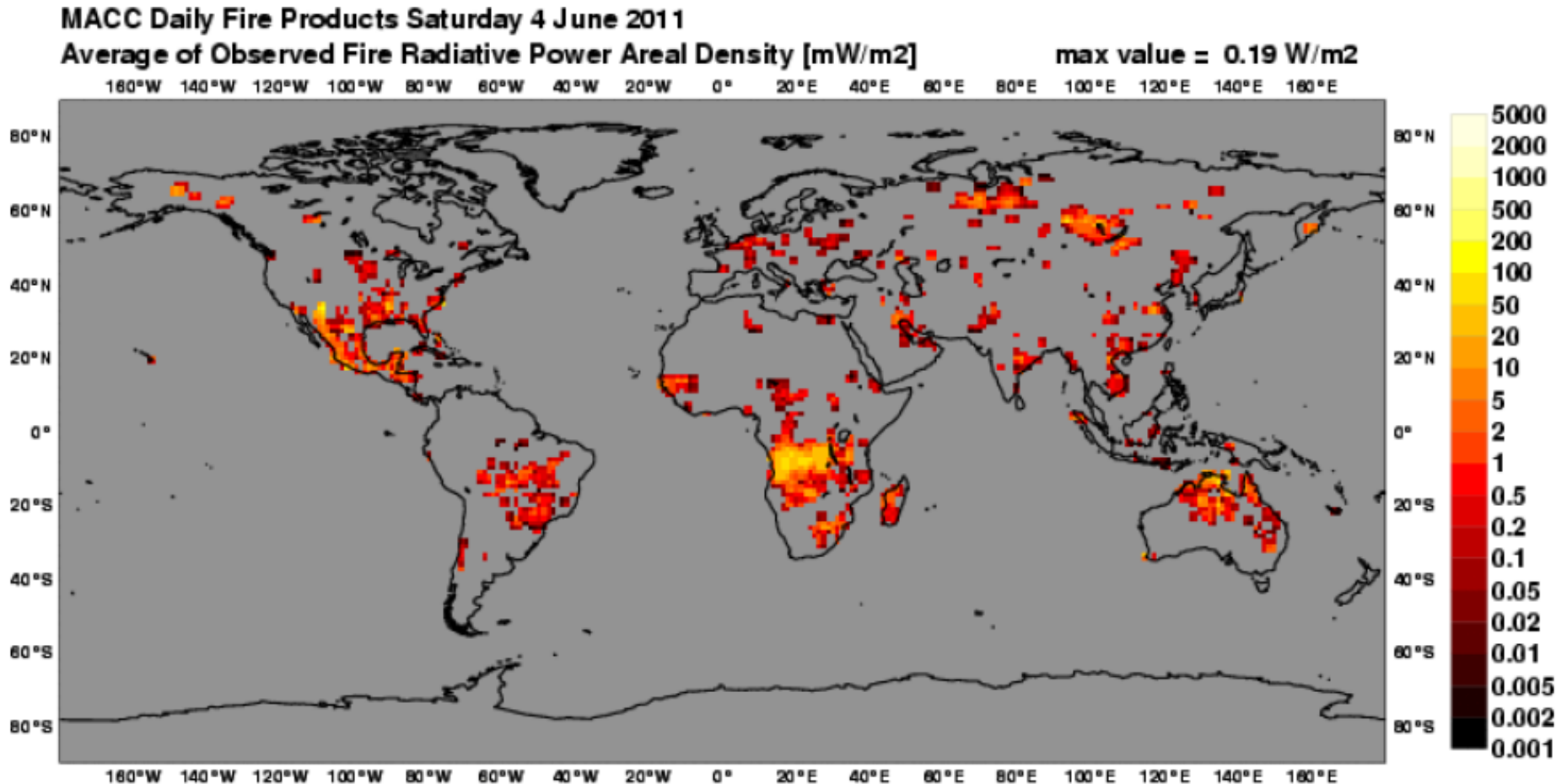




FIRE SATELLITE DATA ASSIMILATION

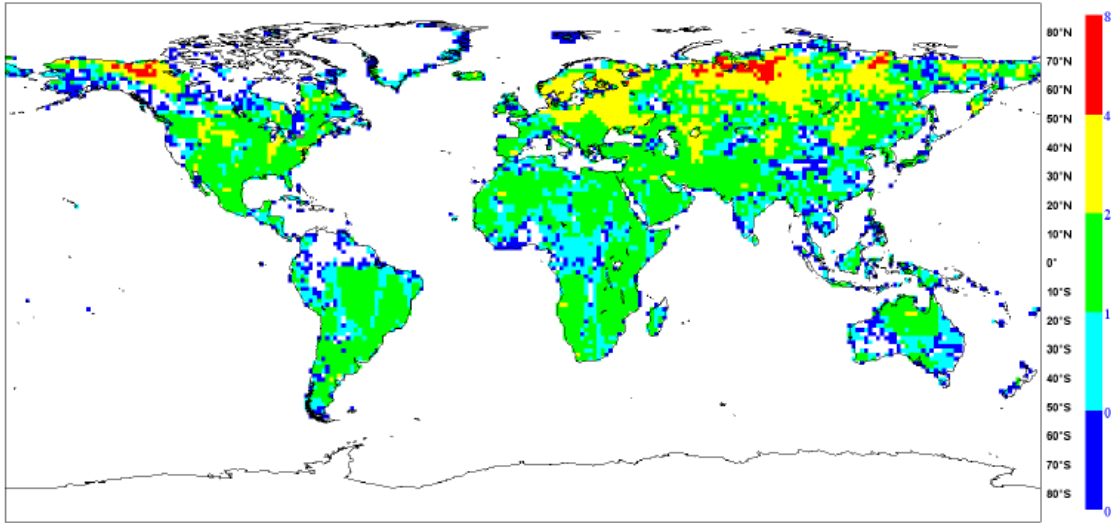
GFAS (Global Fire Assimilation System)

J. Kaiser and coworkers, ECMWF and MPI-C
implemented for CIFS and MACC regional models

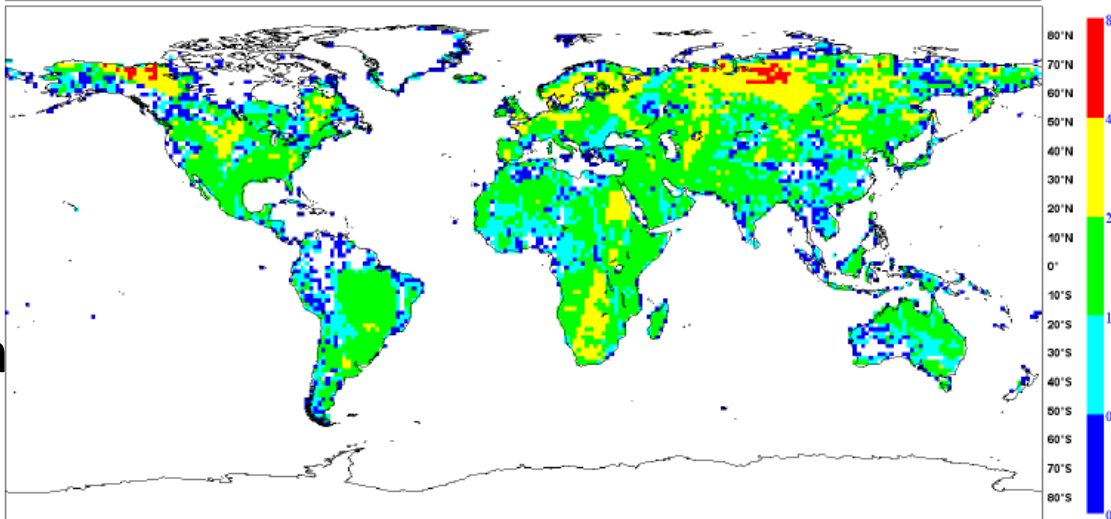


Effective number of satellite observations of grid cells by **MODIS**

MODIS
on TERRA
morning
overpass



MODIS
on AQUA
afternoon
overpass



4 June 2011,
00:00–24:00 UTC

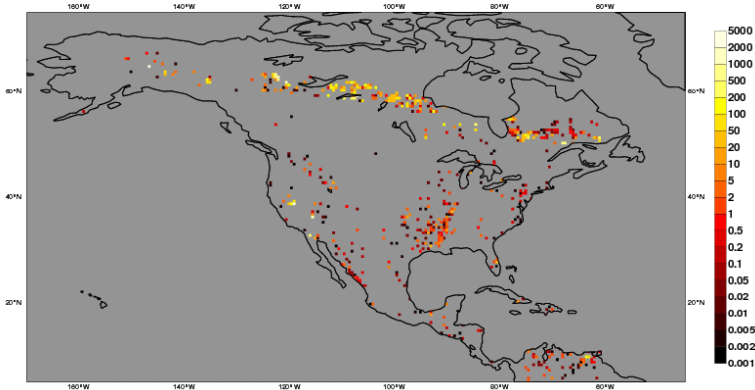
Kaiser, Heil, Andreae, A. Benedetti, N. Chubarova, L. Jones, J.-J. Morcrette, M. Razinger, M. G. Schultz, M. Suttie, and G. R. van der Werf, 2012

Canadian smoke over Europe (July 2013)

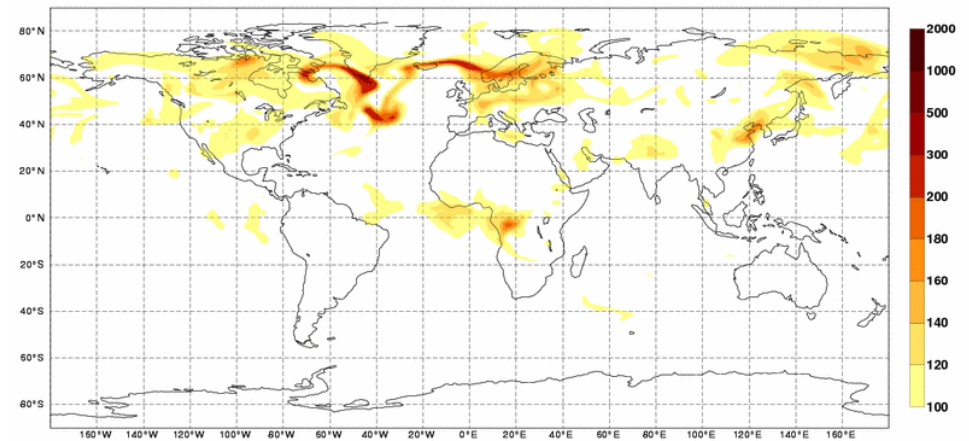
MACC Daily Fire Products Monday 8 July 2013

Average of Observed Fire Radiative Power Areal Density [mW/m²]

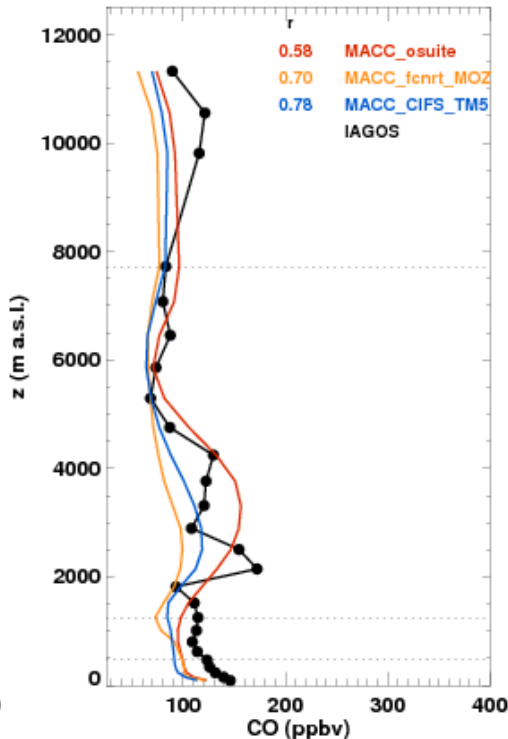
max value = 2.95 W/m²



Monday 8 July 2013 00UTC MACC-II Forecast t+000 VT: Monday 8 July 2013 00UTC
500 mb Carbon Monoxide [ppbv]



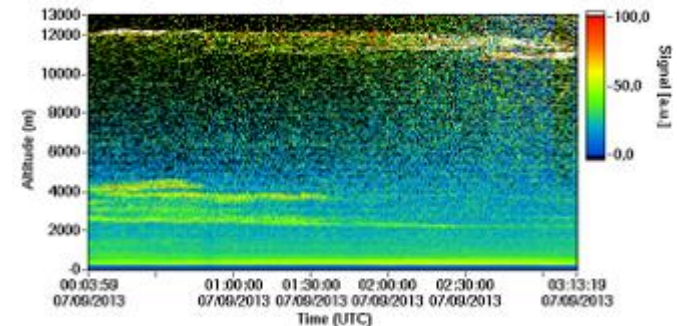
PARIS CO
20130708 (n=1)



The assimilation run (red, MACC o-suite) pick ups increased levels of pollutants, here Carbon Monoxide, between 2 and 4 km. Independent aircraft observations (black) confirm the presence of the plume, which is seen also by European lidars and ceilometers.

J. Kaiser

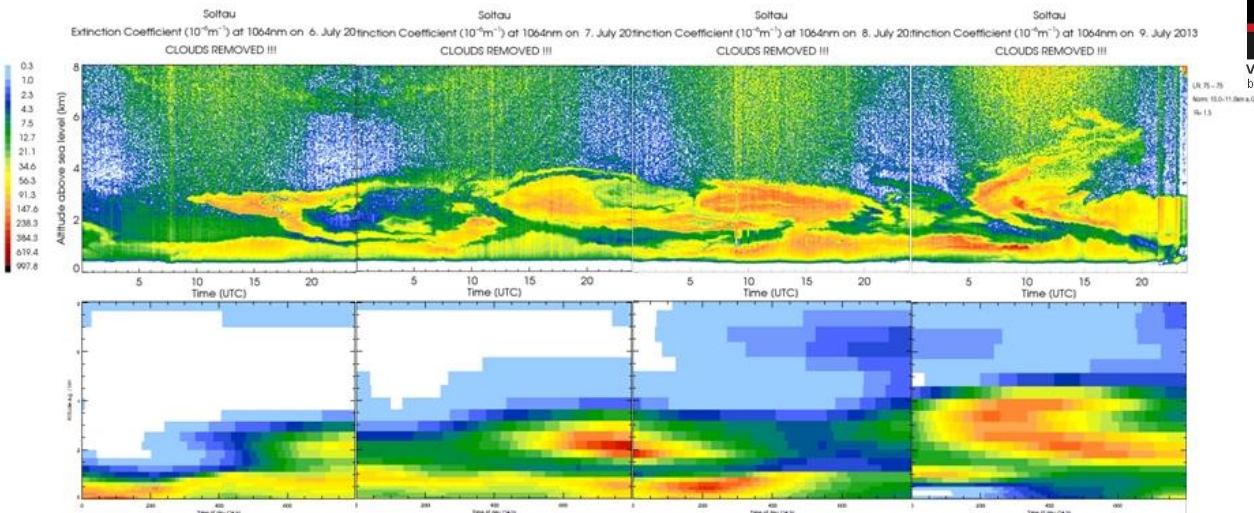
Range corrected signal, 532 nm, Polly 1st Generation (IIT), Stockholm, Sweden



Canadian Smoke in Europe July 2013

Comparison of Canadian forest fire plume seen by
Ceilometers over Soltau, North Germany
6 – 9 July 2013

MACC-2D plot is **QUALITATIVE** and linear scale in contrast to ceiplot!!!
Shall just show the reproduction of the plume structure



Verification of MACC aerosol forecast with ceilometer data
shows good performance for most plume occurrences (plots
courtesy of Harald Flentje, DWD)

A. Benedetti



Video 1. Maxime Duperré, traveling in a truck near Nemiscau, Quebec, took this video of one of the massive fires burning in Quebec this July.

Canada's 2nd largest fire on record
spreading smoke to Europe

Posted by Jim at Monday, July 15, 2013

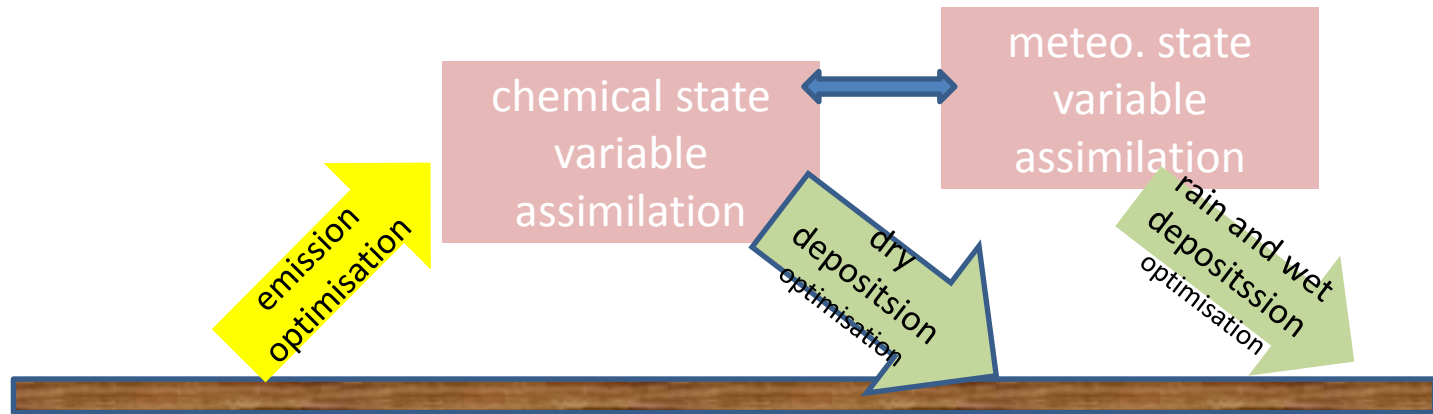


By Dr. Jeff Masters
13 July 2013

FUTURE OBJECTIVES

Future directions

1. boundary layer and air quality data assimilation must include coupling with surface processes (flux inversion):
 - related parameters for biogenic emissions and deposition (LAI, fPAR, ...)



2. aerosols must be classified (lidar colour ratio, lidar ratio, depolarisation, T-matrix,...)
3. joint assimilation/inversion to be extended to further multiple impacting parameters (e.g. radiative)
4. larger ensemble approaches for non-normal errors

Thank You
for Your Attention